

# **In-line Processing in Scanning Atmospheric Radar Networks**

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# Radar Network System

## Radar Nodes



## System Operation & Control Center

Scan Command  
& Control

User  
Requirement

Internet

Real-time Data

Obtained over the Internet

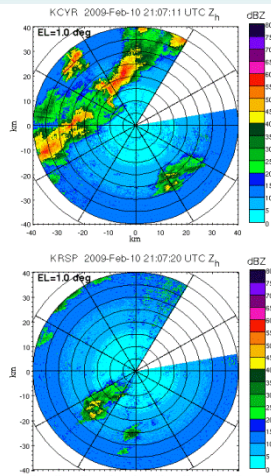
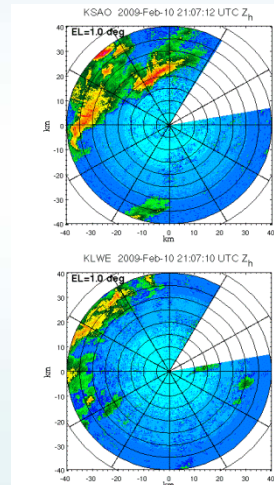
Internet

## End-users



Internet

## Signal Processor



High bandwidth  
Storage

Data Storage

Retrieval algorithm  
and Data Archive

# Real time processing of Radar parameter computation

- Radar Variables
- Ground Clutter Filtering
- Range-Velocity Ambiguity Mitigation
- Overlaid Echo Suppression

# Estimates of Spectral Moment

- ❖ The mean received power provides the intensity of precipitation within the resolution volume

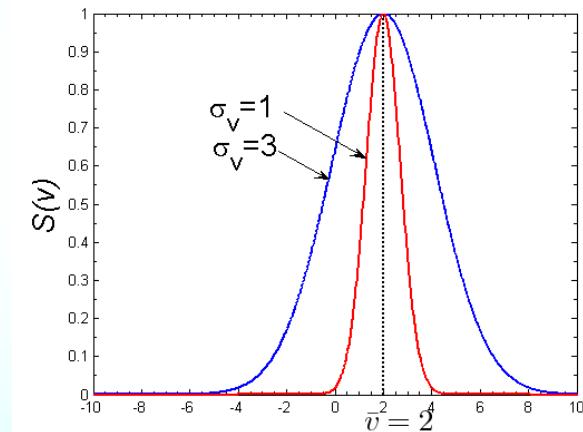
$$\text{time domain : } \hat{P}_h = \frac{1}{N} \sum_{k=0}^{N-1} |v_h[k]|^2 \quad \text{spectral domain : } \hat{P}_h = \frac{1}{N^2} \sum_{k=0}^{N-1} |s_h[k]|^2$$

- ❖ The mean radial velocity of the precipitation is obtained based on the Doppler phase shift (caused by the motion of particles) of the received signal

$$\hat{R}(1) = \frac{1}{N} \sum_{k=0}^{N-2} v_h(k+1) v_h^*(k)$$

$$\text{time domain : } \hat{v} = -\frac{\lambda}{4\pi T_s} \arctan \{ \hat{R}(1) \}$$

$$\text{spectral domain : } \hat{v} = \frac{\sum_{k=0}^{N-1} v[k] s[k]^2}{\sum_{k=0}^{N-1} |s[k]|^2}$$



- ❖ The precipitating volume consist of a large number of hydrometeors with widely varying radial velocities resulting a Doppler velocity spread about the mean velocity. The spectral width gives a measure of turbulence of the medium

$$\text{time domain : } \hat{\sigma}_v = \frac{\lambda}{2\pi T_s \sqrt{2}} \sqrt{\ln \left| \frac{R(0)}{R(1)} \right|}$$

$$\text{spectral domain : } \hat{\sigma}_v = \sqrt{\frac{\sum_{k=0}^{N-1} (v_k - \hat{v})^2 |s_k|^2}{\sum_{k=0}^{N-1} |s_k|^2}}$$

# Estimates of Polarimetric Variables

- ❖ Differential reflectivity between polarization channels provides a measure of mean particle shape

$$\hat{Z}_{dr} = 10 \log_{10} \left( \frac{\hat{\rho}_h}{\hat{\rho}_v} \right)$$

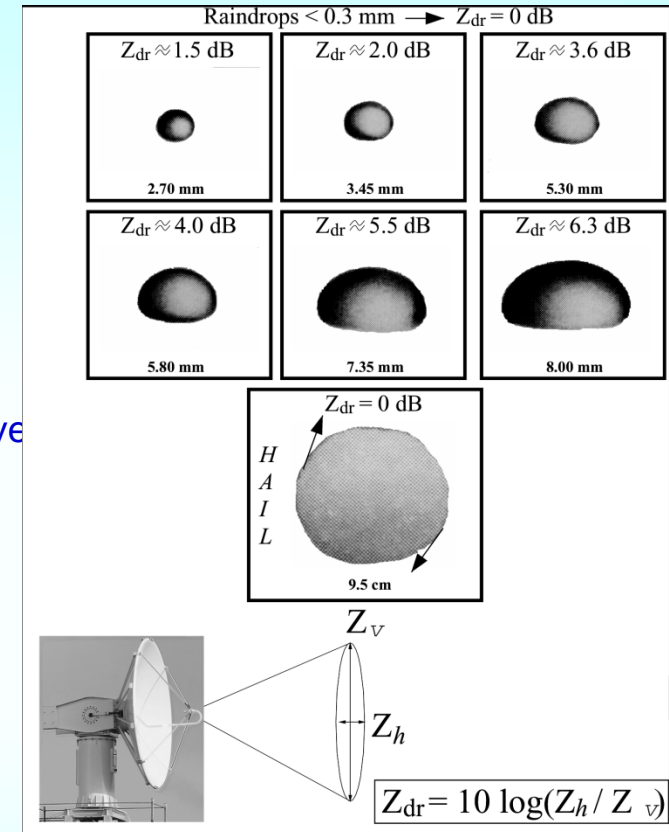
- ❖ The differential propagation phase shift is the phase difference between vertical polarization and horizontal polarization as the wave propagates through rain

$$\text{time domain : } \hat{\psi}_{dp} = \arctan \left\{ \sum_{k=0}^{N-1} v_v[k] v_h^*[k] \right\}$$

- ❖ The correlation between the received signal in the horizontal polarization and vertical polarization gives an indication of similarity in the nature of back scattering from the hydrometeors

$$\text{time domain : } |\hat{\rho}_{hv}(0)| = \frac{\left| \sum_{k=0}^{N-1} v_v[k] v_h^*[k] \right|}{\sqrt{\sum_{k=0}^{N-1} |v_h[k]|^2 \sum_{k=0}^{N-1} |v_v[k]|^2}}$$

$$\text{spectral domain : } |\hat{\rho}_{hv}(0)| = \frac{\left| \sum_{k=0}^{N-1} s_v[k] s_h^*[k] \right|}{\sqrt{\sum_{k=0}^{N-1} |s_h[k]|^2 \sum_{k=0}^{N-1} |s_v[k]|^2}}$$



# Spectral Clutter Filtering Example

- ❖ Obtain spectral coefficients and power spectral density of received signal

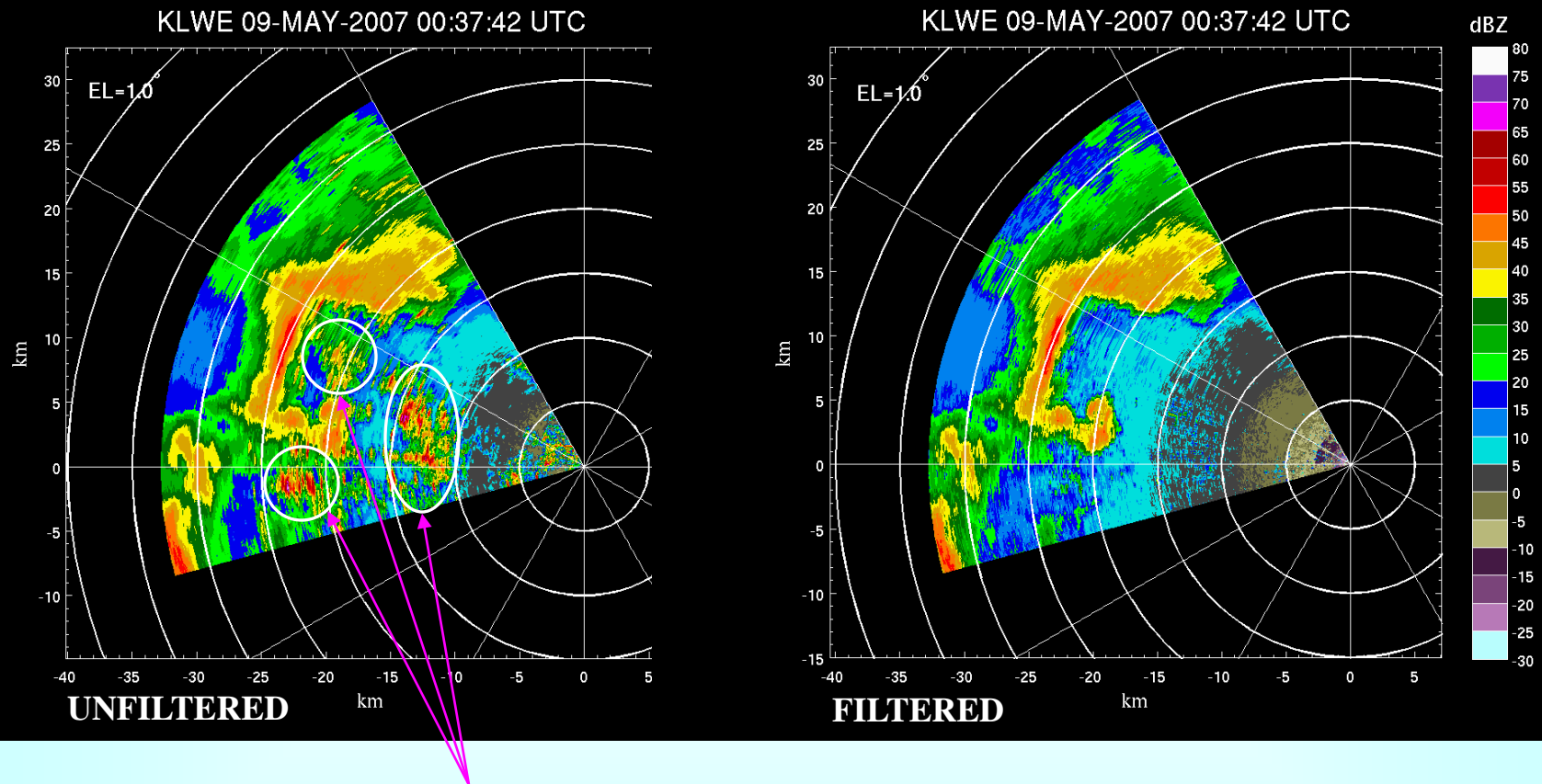
$$S(\mathbf{v}, \theta) = \frac{p_c}{\sqrt{2\pi\sigma_c^2}} \exp\left\{-\frac{\mathbf{v}^2}{2\sigma_c^2}\right\} + \frac{p}{\sqrt{2\pi w^2}} \exp\left\{-\frac{(\mathbf{v} - v)^2}{2w^2}\right\} + \frac{2T_s}{\lambda} p_n$$

- ❖ Obtain adaptive noise floor by sorting spectral coefficients by power
- ❖ Design notch filter in spectral domain
  - Estimate clutter model based on Gaussian model fit to zero Doppler region
  - Estimate notch width based on clutter model and noise

$$n = \left\lceil \frac{4\sigma_c T_s}{\lambda} \sqrt{2 \ln \left[ \sqrt{2\pi\sigma_c} \left( \frac{p_c}{\bar{p}_n} \right) \right]} \right\rceil$$

- ❖ Notch the clutter signal with a spectral clipper
- ❖ Interpolate the notch filtered region by iteratively fitting a Gaussian model to the weather signal
- ❖ Replace the clutter region with model and subtract noise power

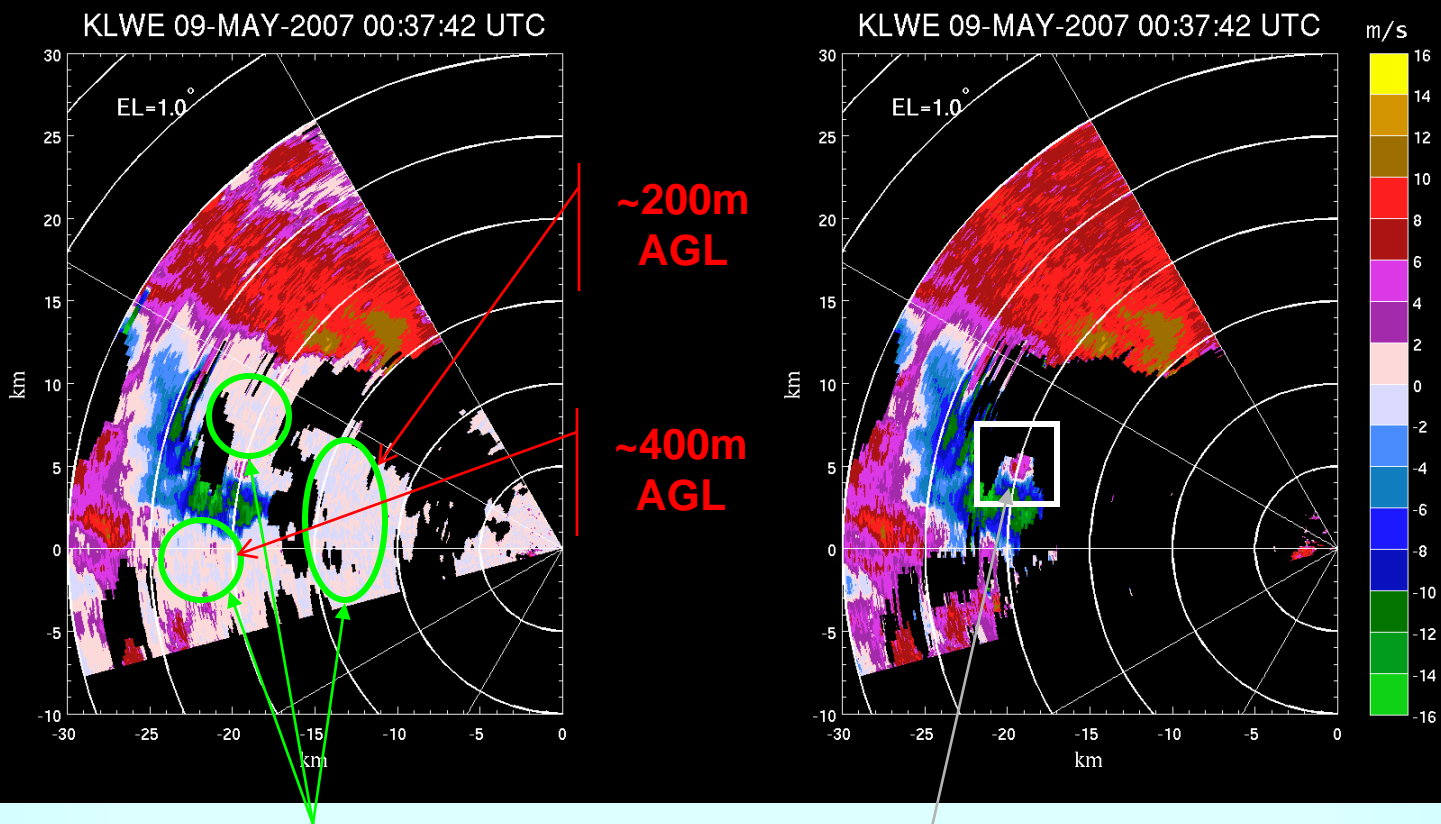
# *Ground Clutter Filtering*



Ground clutter

Reflectivity before and after ground clutter filtering. Data collected on May 09, 2007 at Lawton (EL=1 deg).

# Ground Clutter Filtering



Ground clutter

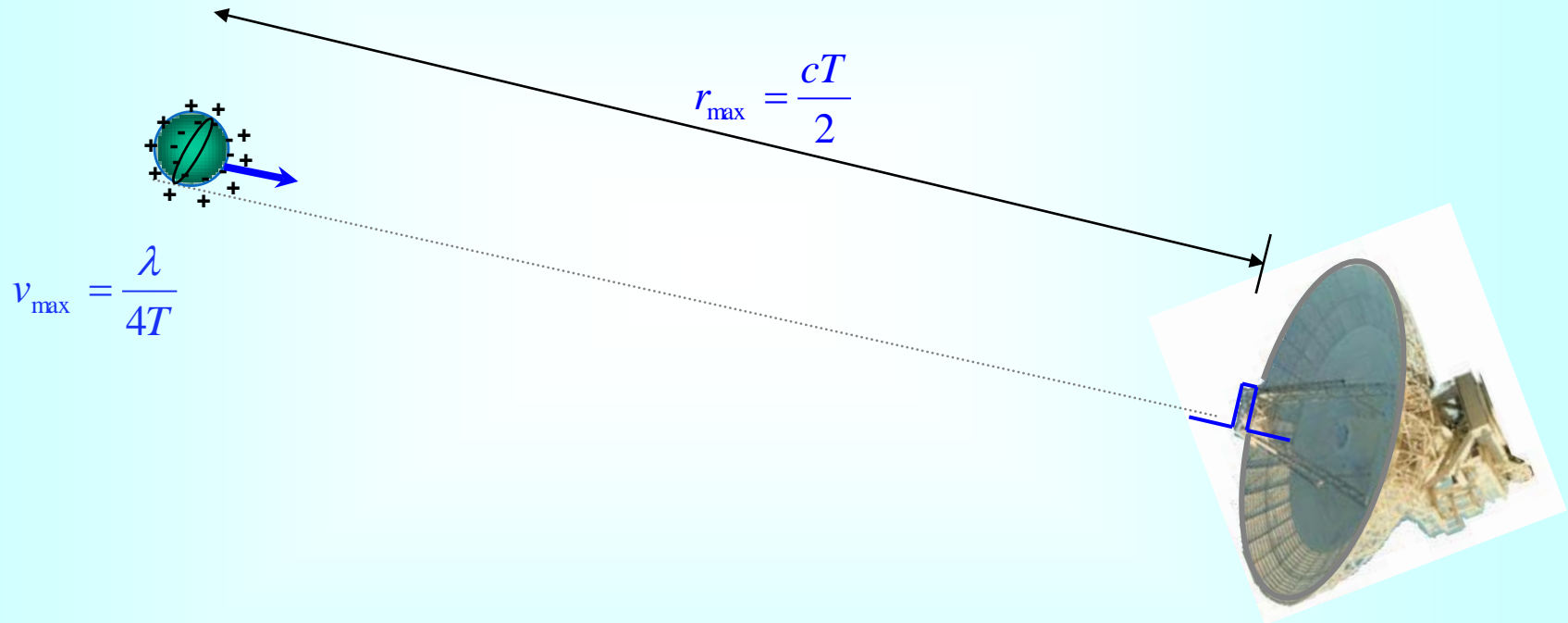
Circulation signature

Velocity before and after ground clutter filtering.  
Data collected on May 09, 2007 at Lawton ( $EL=1$  deg).



# Range-velocity Ambiguity Mitigation Methods

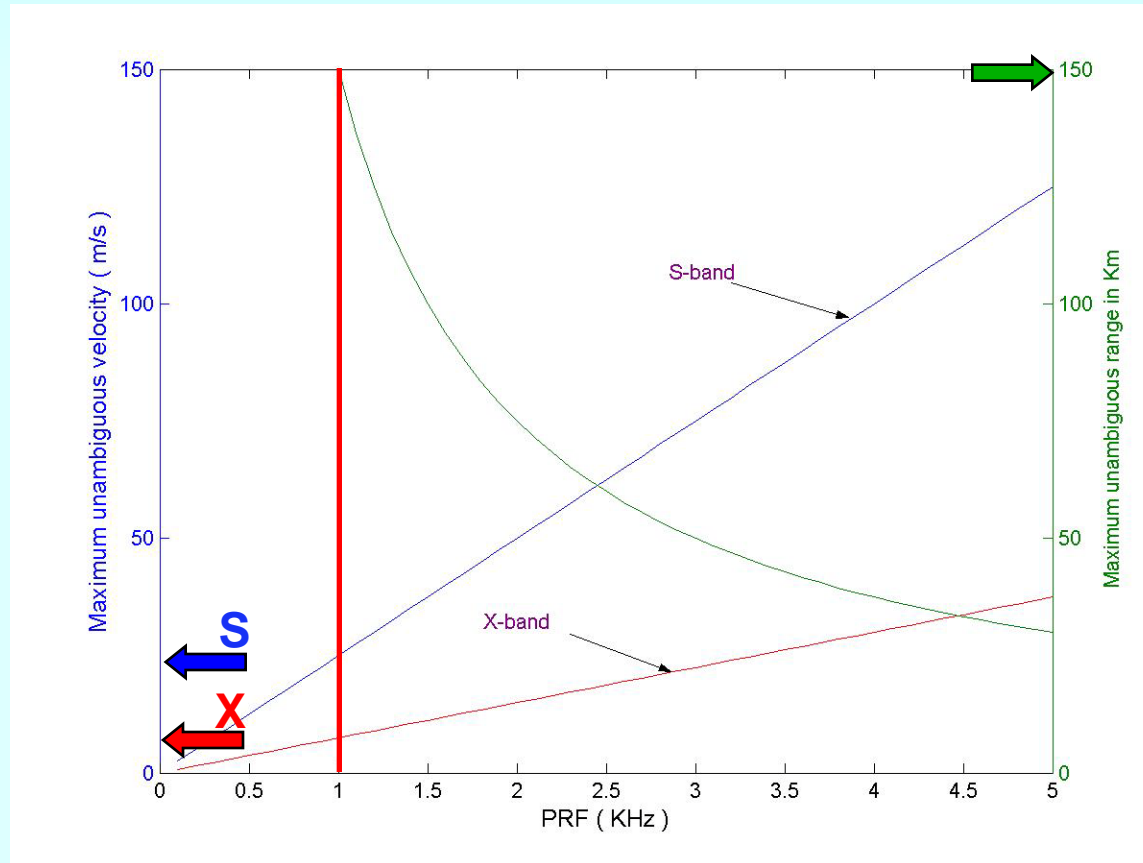
- The maximum unambiguous range and unambiguous velocity have a limitation based on wavelength and pulse repetition time



- Maximum unambiguous range and unambiguous velocity are related to each other as

$$r_{\max} v_{\max} = \frac{c\lambda}{8}$$

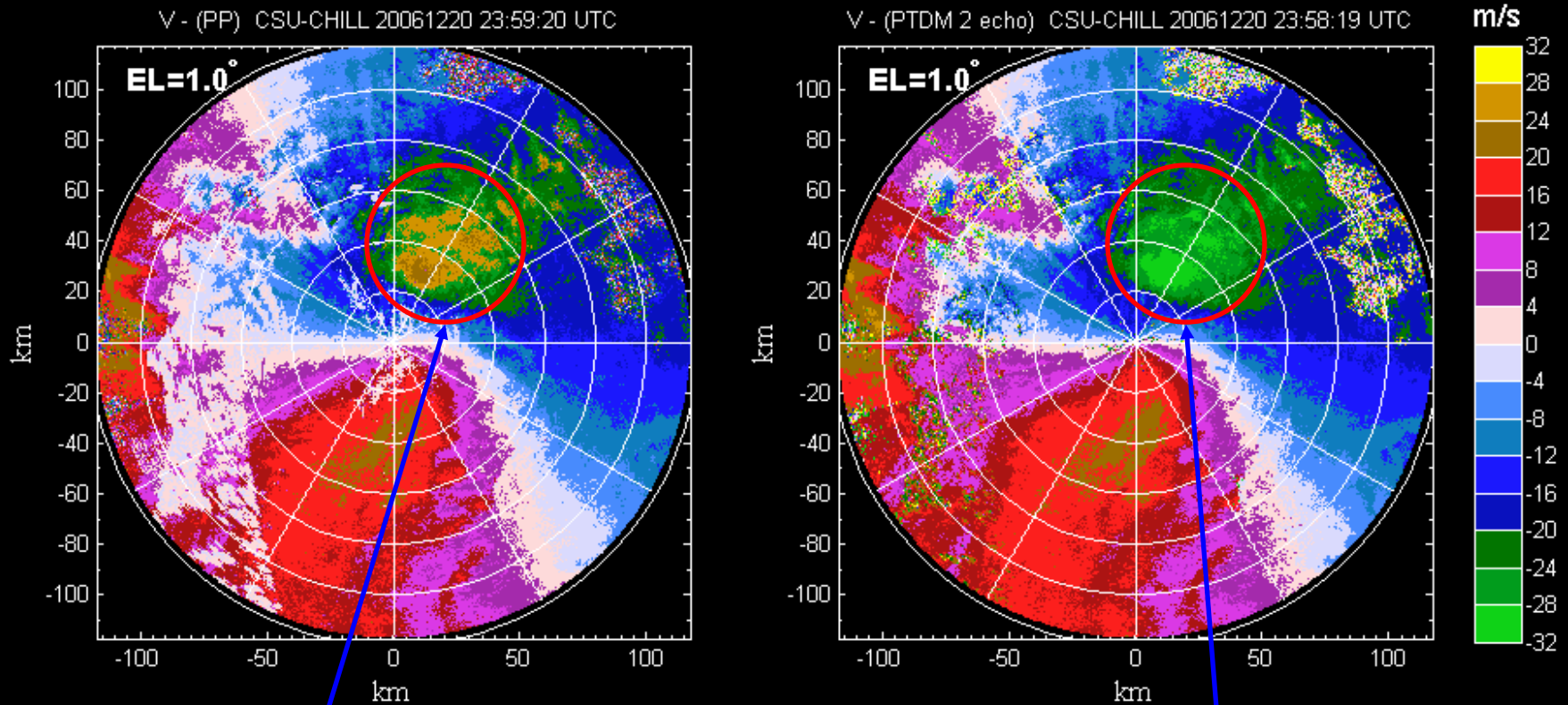
# Range-velocity Ambiguity Mitigation Methods



- If  $v_{max}$  is increased then  $r_{max}$  decreases correspondingly ( Range-velocity ambiguity)
- Fundamental limitation of pulsed Doppler radar transmitting uniformly spaced pulses

# Radial Velocity Folding in Severe Weather

Velocity measurements with uniform PRT and staggered PRT with CSU-CHILL 2006-Dec-20



Velocity folding from -27 m/s to 27 m/s

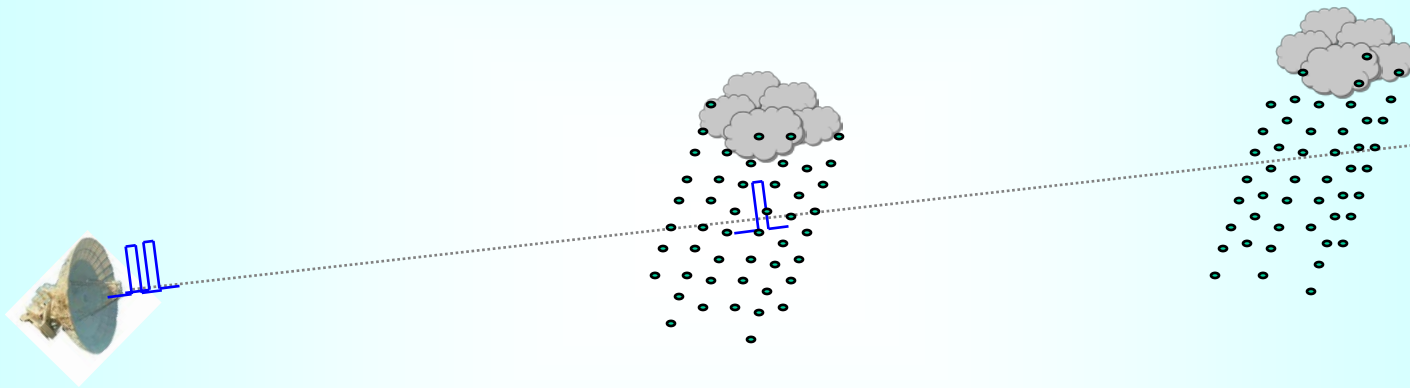
Unfolded velocity folding using staggered PRT waveform

# Range-velocity Ambiguity Mitigation Methods

- Phase coding to mitigate range ambiguity
  - Random phase coding
  - Systematic phase coding
- Staggered pulsing to mitigate
  - Staggered PRT
  - Staggered PRF
- Polarization diversity to mitigate range ambiguity

# Range-overlay due to shorter PRT

- In order to obtain reasonable unambiguous velocities the Doppler radar's PRT is significantly shorter



- The shorter PRTs results in range-overlaid echoes which gives erroneous measurements of the Doppler spectral moments

● : Resolution volume  
 $r_m$  : Maximum unambiguous range  
 $T$  : Pulse repetition time (PRT)

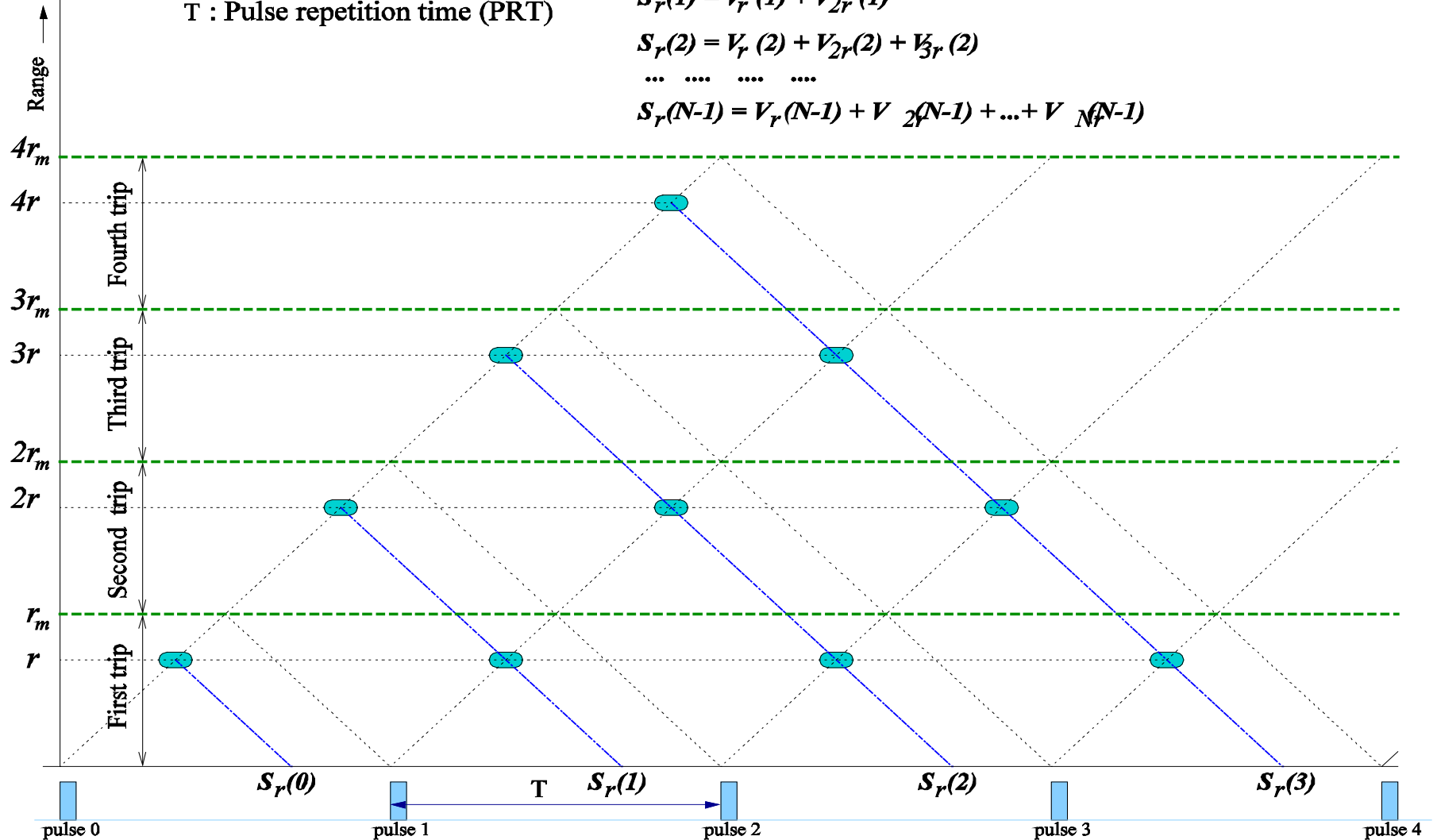
$$S_r(0) = V_r(0)$$

$$S_r(1) = V_r(1) + V_{2r}(1)$$

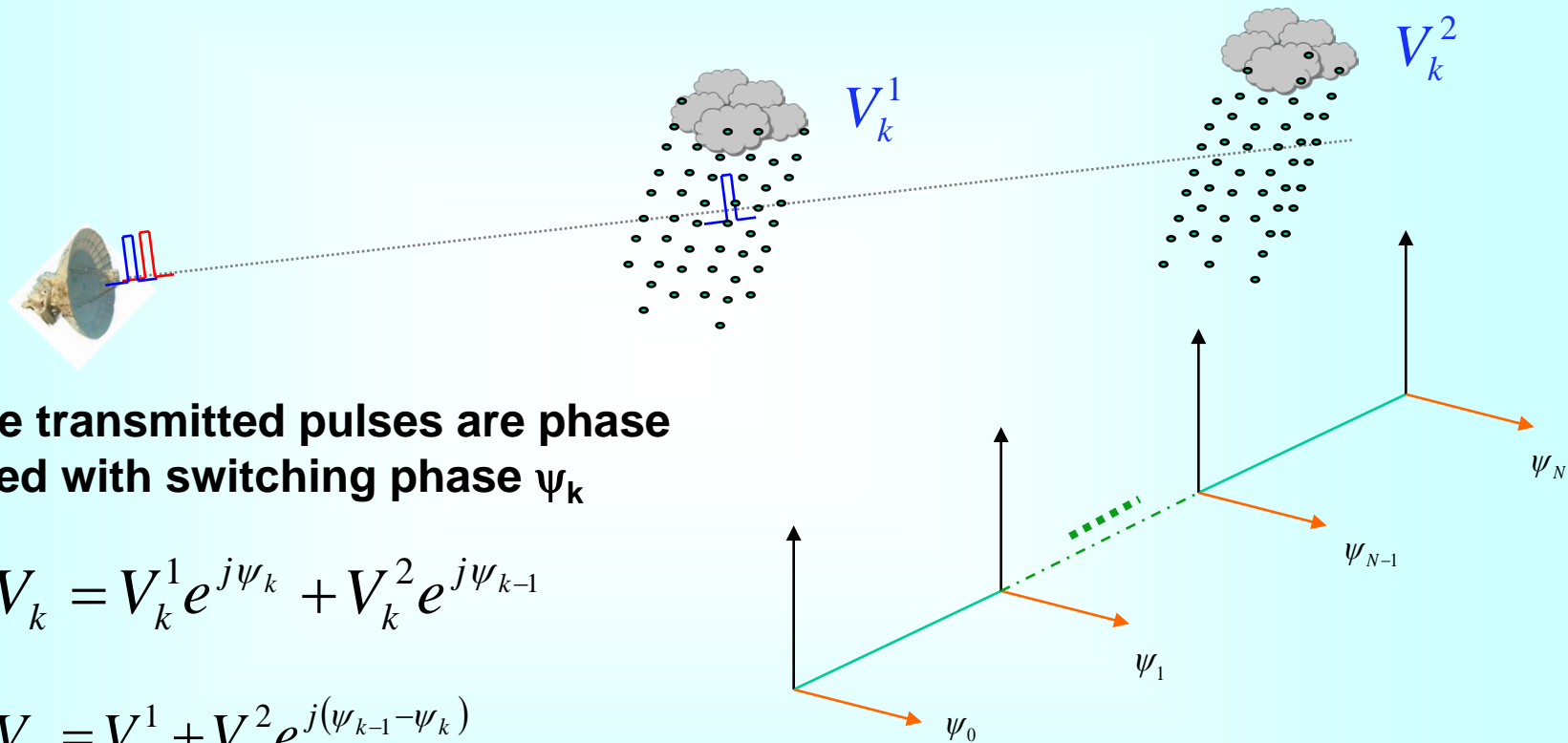
$$S_r(2) = V_r(2) + V_{2r}(2) + V_{3r}(2)$$

... ..

$$S_r(N-1) = V_r(N-1) + V_{2r}(N-1) + \dots + V_{Nr}(N-1)$$



The transmitted pulses are phase coded and the received signal is cohered for the first and second trip




- The transmitted pulses are phase coded with switching phase  $\psi_k$

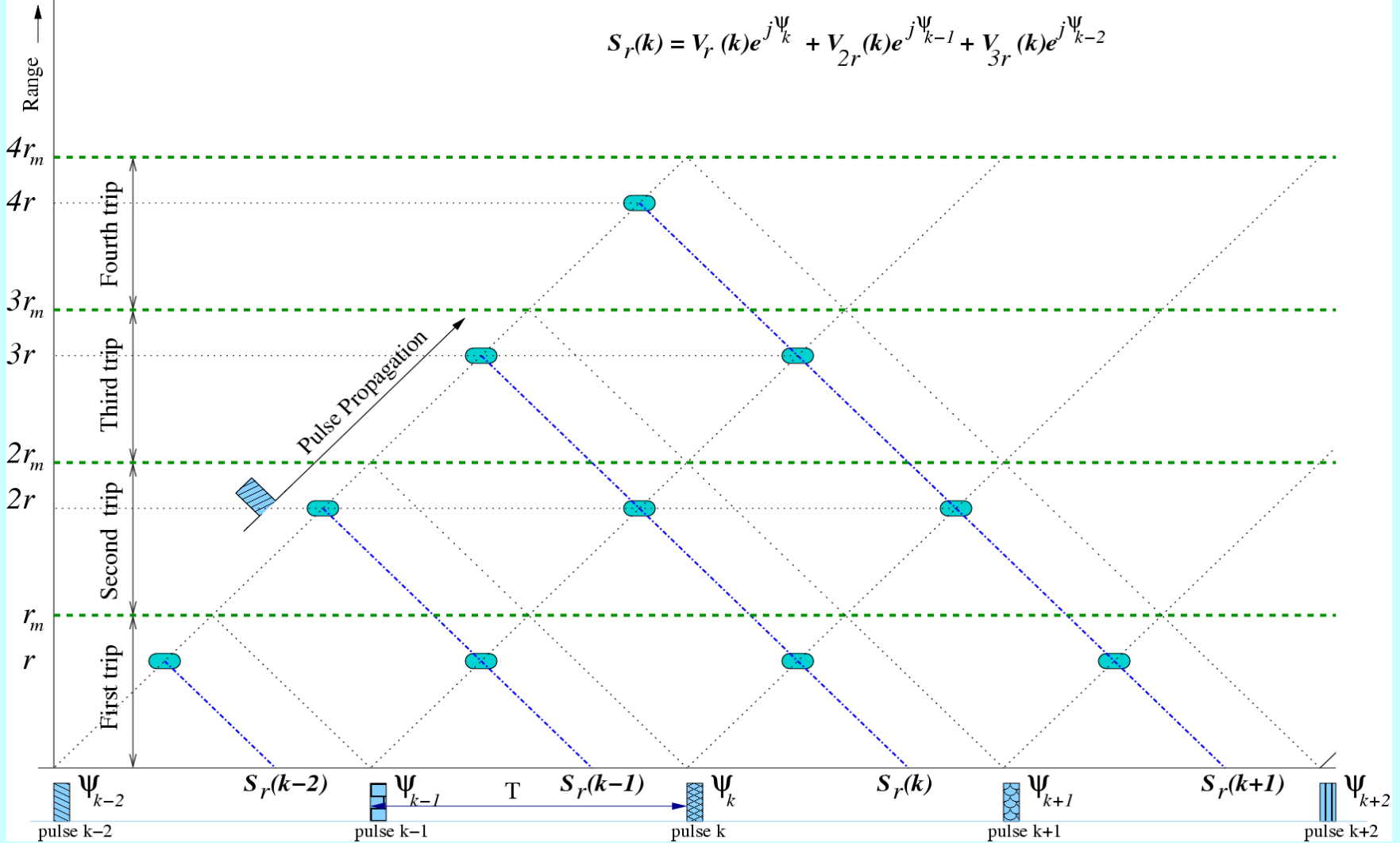
$$V_k = V_k^1 e^{j\psi_k} + V_k^2 e^{j\psi_{k-1}}$$

$$V_k = V_k^1 + V_k^2 e^{j(\psi_{k-1} - \psi_k)}$$

Where  $e^{j(\psi_{k-1} - \psi_k)} = e^{j\phi_k}$  is the modulation code

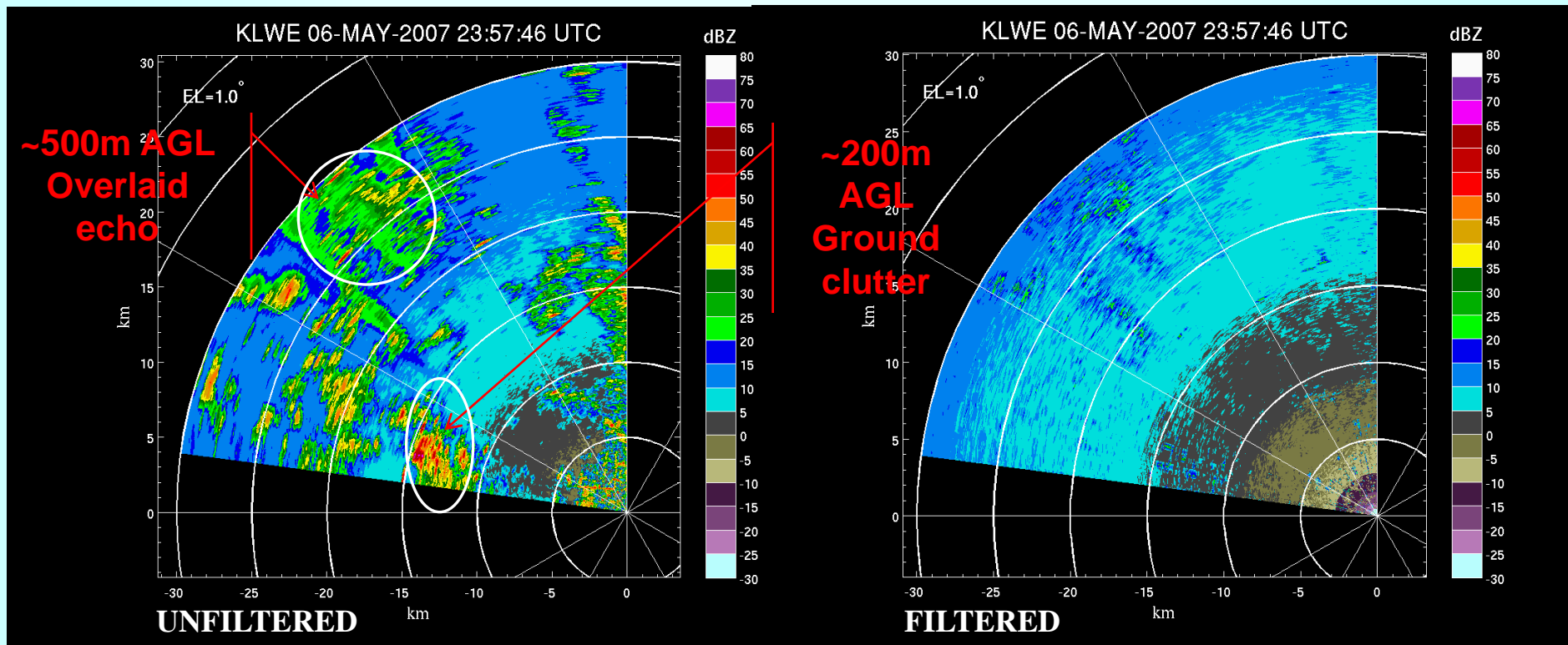
-  : Resolution volume
- $r_m$  : Maximum unambiguous range
- T : Pulse repetition time (PRT)

$$S_r(k) = V_r(k)e^{j\psi_k} + V_{2r}(k)e^{j\psi_{k-1}} + V_{3r}(k)e^{j\psi_{k-2}}$$





# Random Phase Coding For Range-Overlay Suppression



Before overlaid echo suppression and clutter filtering. Data collected on May 06, 2007 at Lawton (EL=1 deg).

After overlaid echo suppression and clutter filtering. Data collected on May 06, 2007 at Lawton (EL=1 deg).

# Staggered PRT for Mitigating Velocity Ambiguity

- A periodic block pulsing scheme can be used for range –velocity ambiguity mitigation
- In general we can have  $T_1, T_2, T_3, \dots, T_n$  such that we satisfy

$$\sum_{j=1}^n T_j = T$$

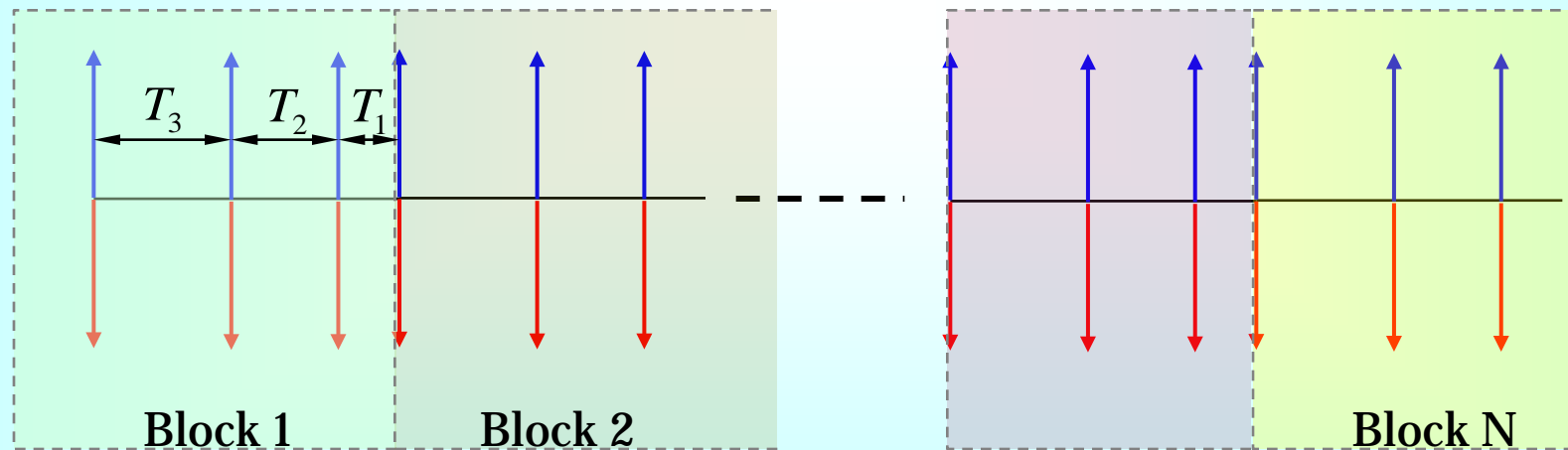
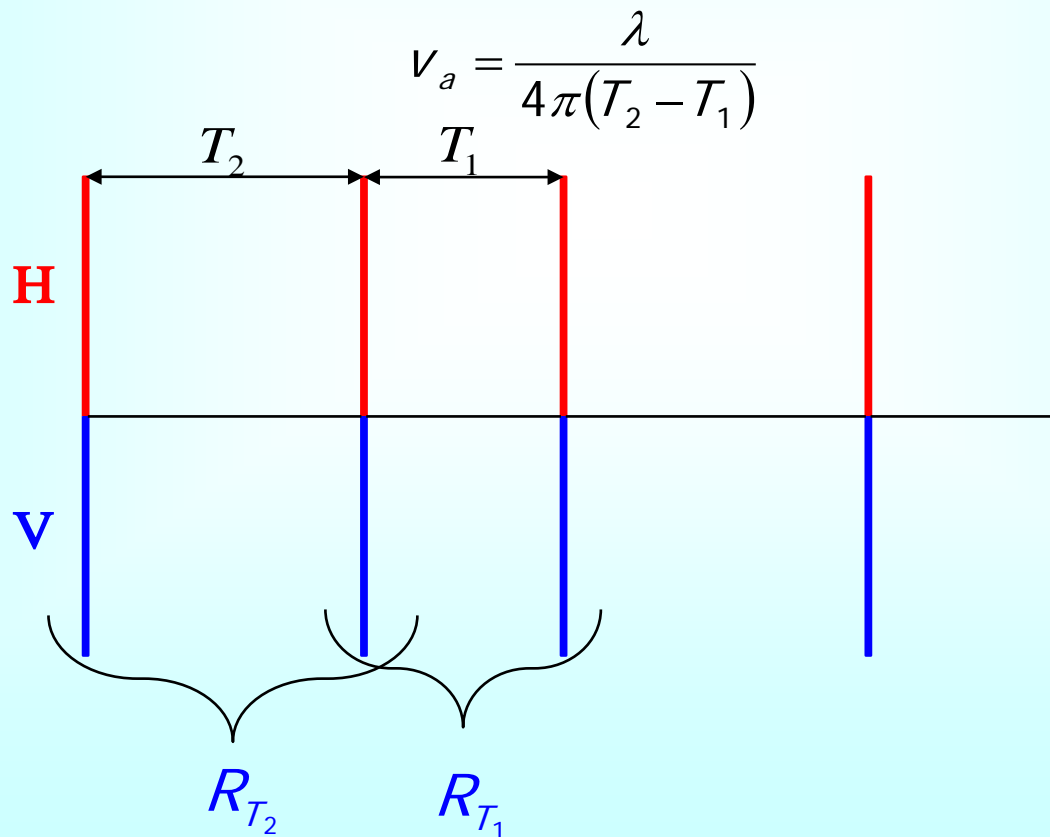


Illustration a periodic block pulsing scheme in hybrid mode of operation for a dual polarized radar

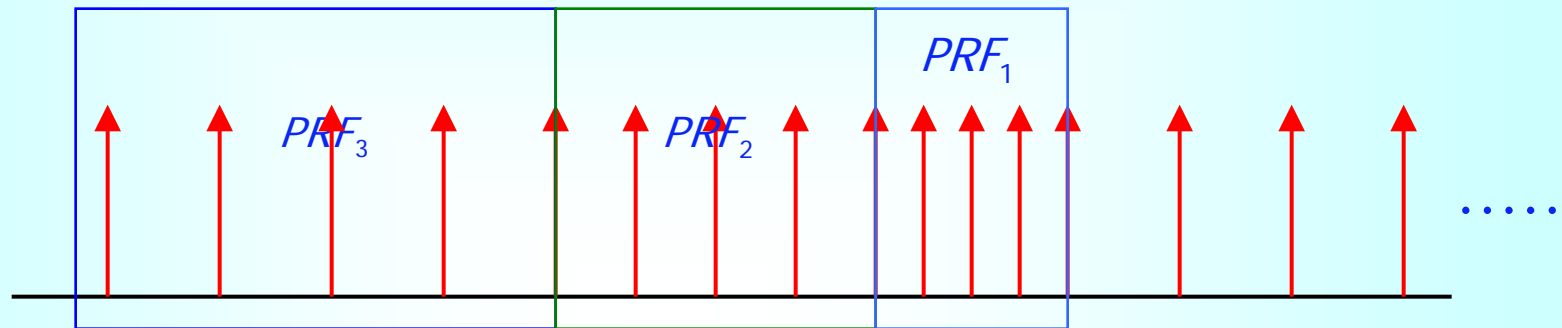
# Staggered PRT for Mitigating Velocity Ambiguity

- Staggered PRT scheme with two PRTs  $T_1$  and  $T_2$  will increase the maximum unambiguous velocity to



# Staggered PRF

- In the staggered PRF technique pulses are transmitted with different PRFs

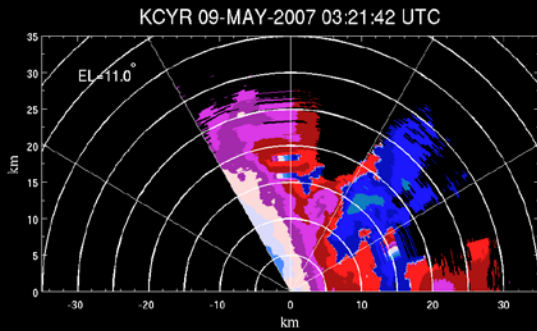


- Dual-PRF waveform are more common in weather radars to increase the unambiguous velocity
- The maximum unambiguous velocity obtained from Dual-PRF waveforms is given by

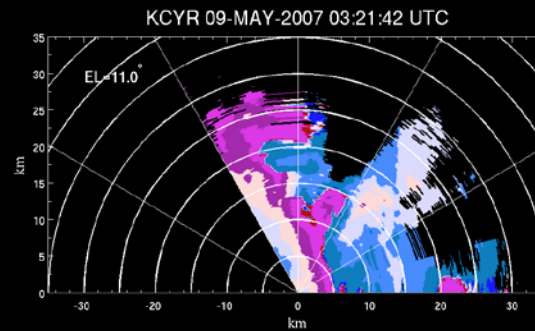
$$V_a = \frac{V_{a1}V_{a2}}{V_{a1} - V_{a2}}$$

- The unfolding procedure of velocity  $V_1$  is similar to the unfolding procedure of staggered PRT

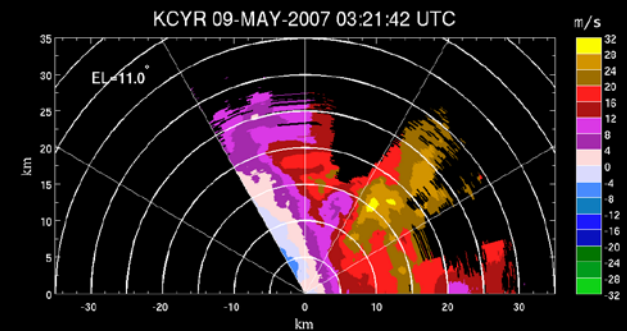
# Velocity Unfolding



**PRF=2.4 kHz**



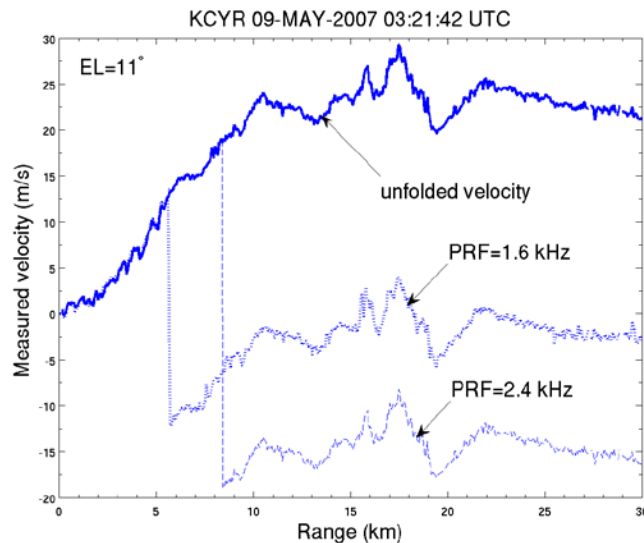
**PRF=1.6 kHz**



**Unfolded**

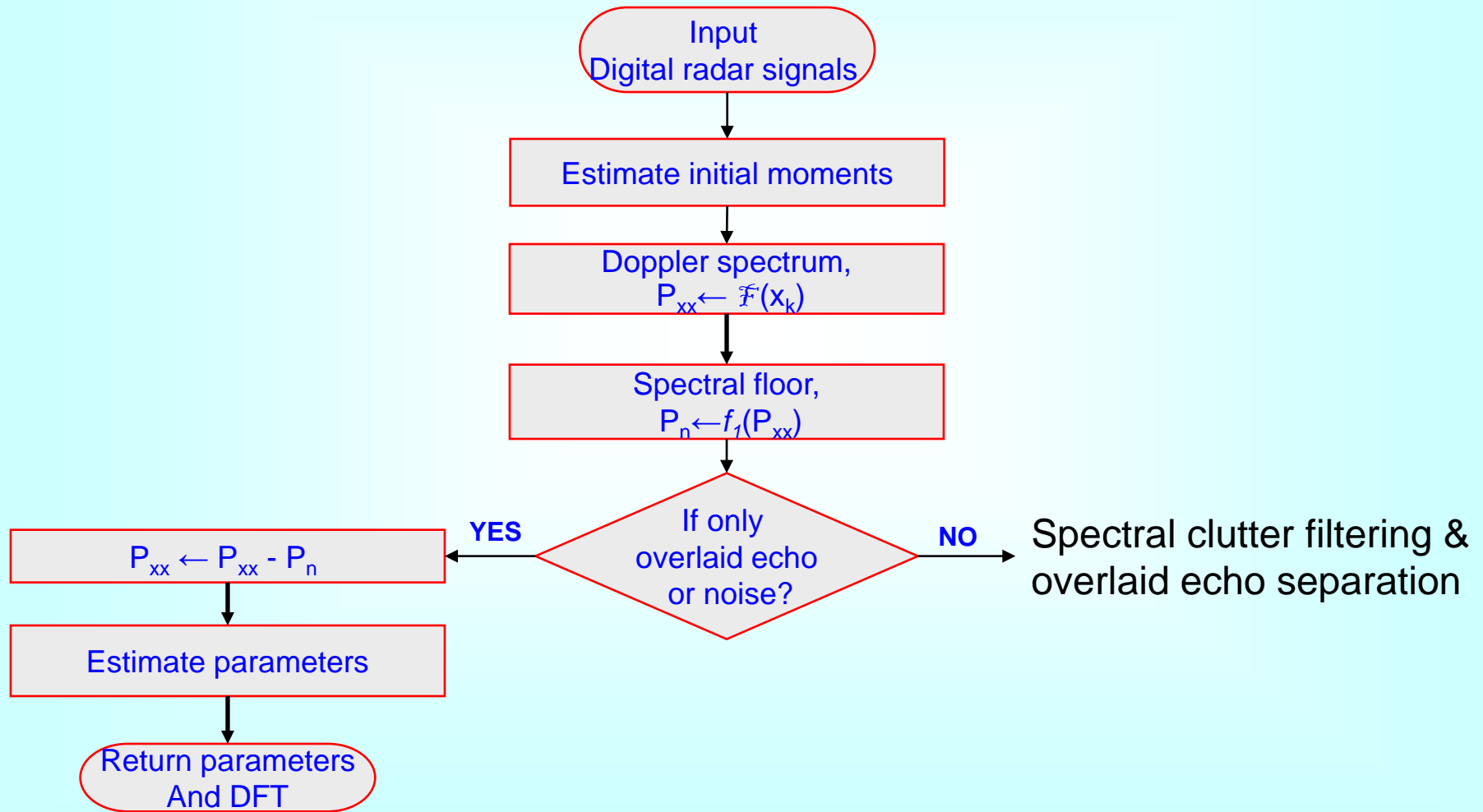
Velocity measurements on a radial-by-radial basis with dual-PRF

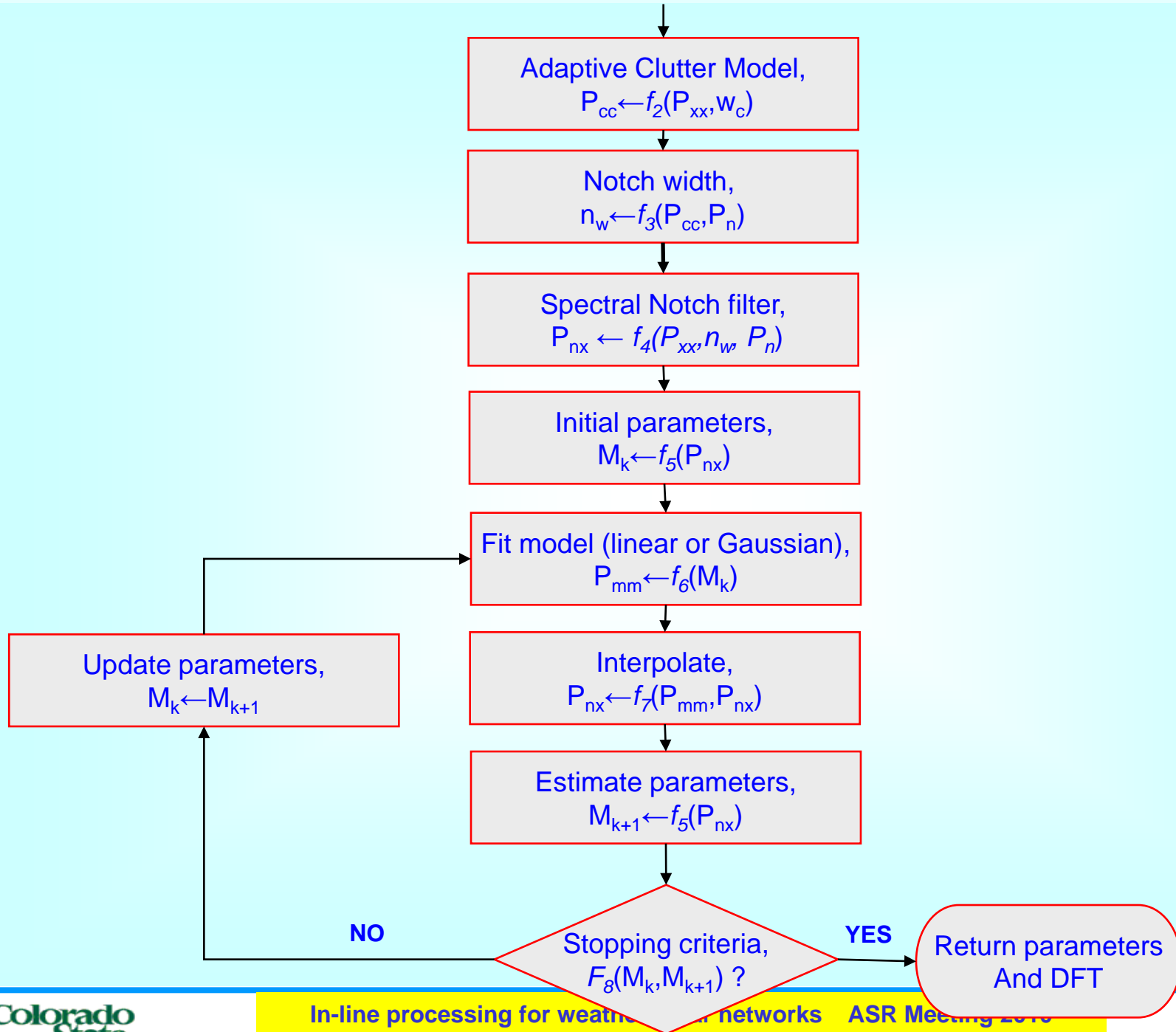
Unfolded velocity with dual-PRF



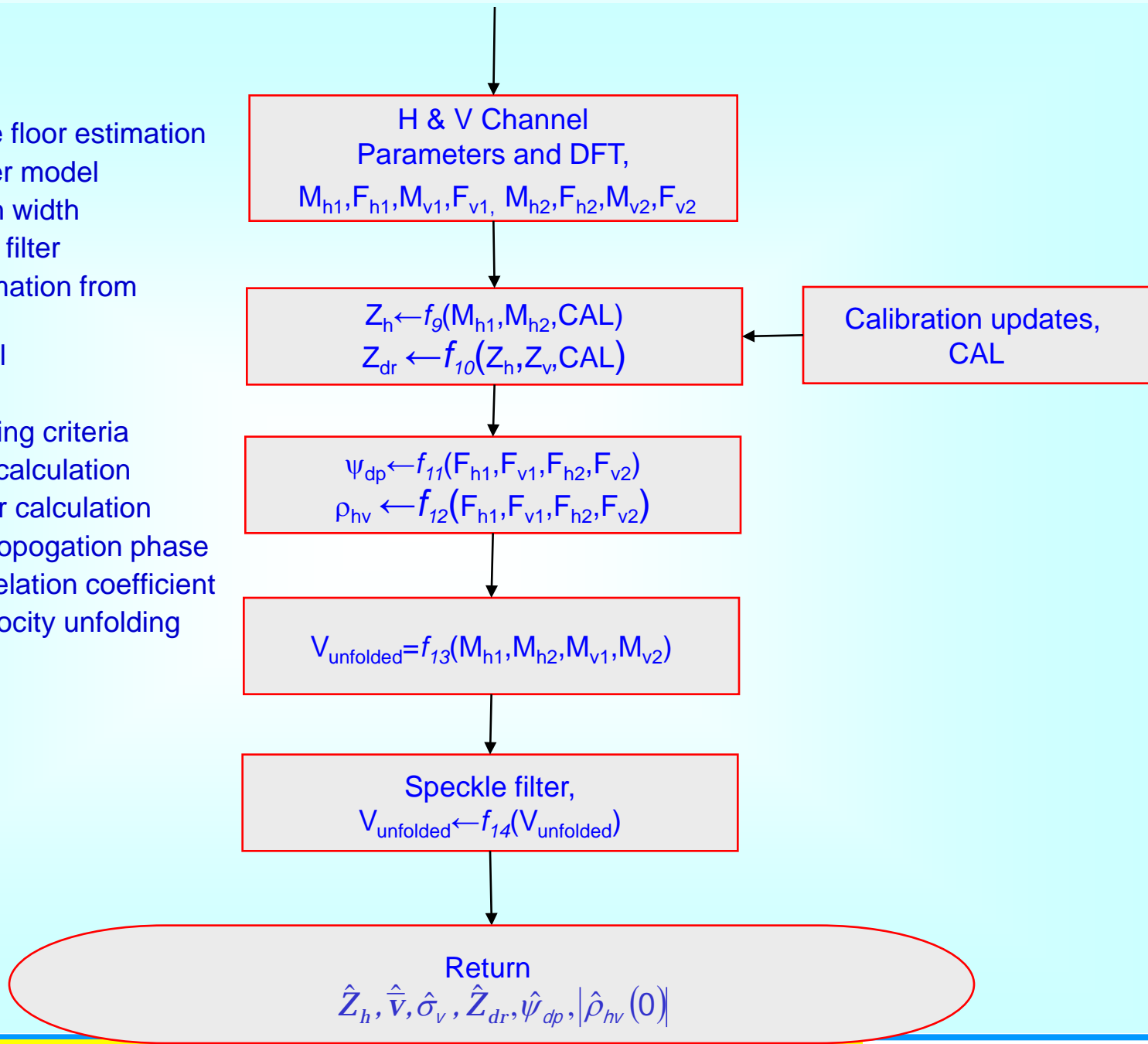
Doppler velocity from Dual-PRF. Data collected on May 09, 2007 at Cyril (EL=11 deg).

# Radar Product Generation





- $f_1$ = Dynamic noise floor estimation
- $f_2$ = Adaptive clutter model
- $f_3$ = Adaptive notch width
- $f_4$ = Spectral notch filter
- $f_5$ = Moments estimation from spectrum
- $f_6$ = Spectral model
- $f_7$ = Interpolation
- $f_8$ = Iteration stopping criteria
- $f_9$ = Calibrated  $Z_h$  calculation
- $f_{10}$ = Calibrated  $Z_{dr}$  calculation
- $f_{11}$ = Differential propagation phase
- $f_{12}$ = Co-polar correlation coefficient
- $f_{13}$ = Dual-PRF velocity unfolding
- $f_{14}$ = Speckle filter

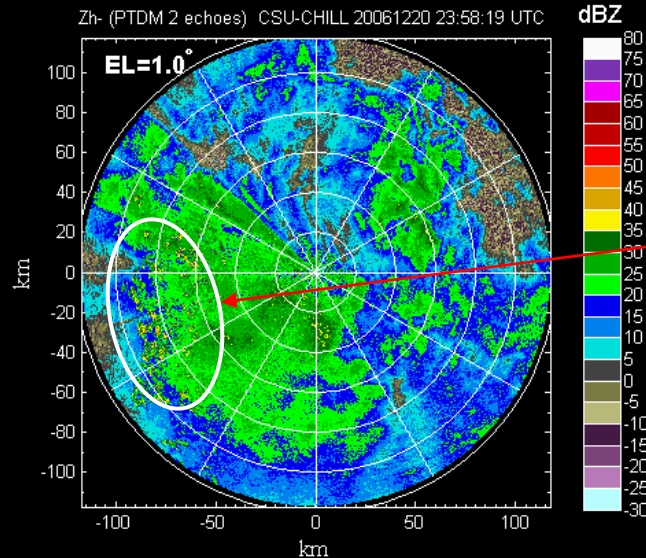
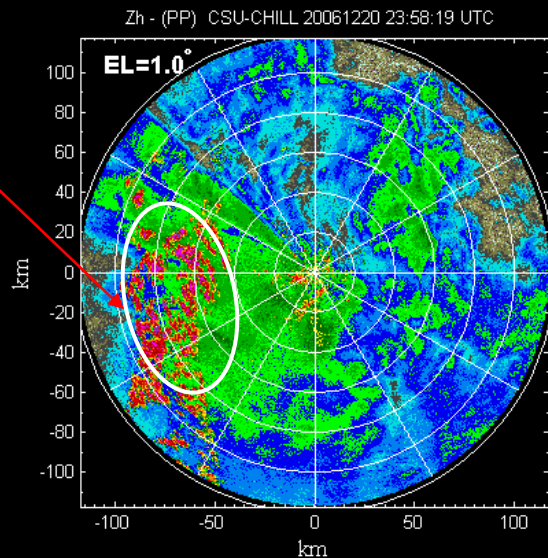




# Data Products with CSU-CHILL

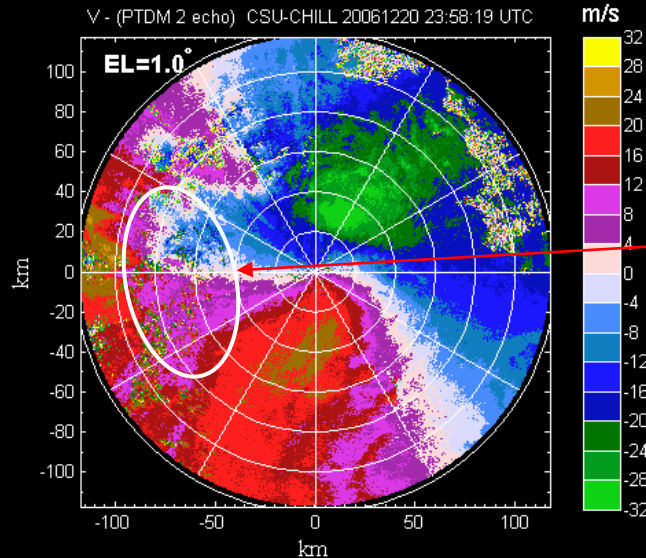
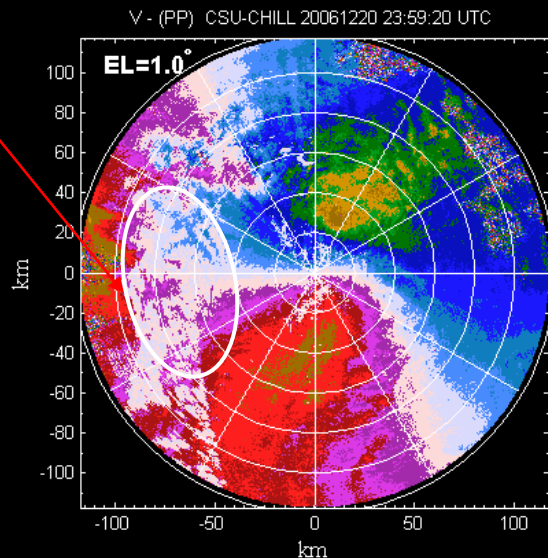
Data collected with PRT = 1 ms; N=64 on Dec 20, 2006 at 23:58:19 UTC

Severe clutter from Rocky mountains



Reflectivity after spectral clutter filtering

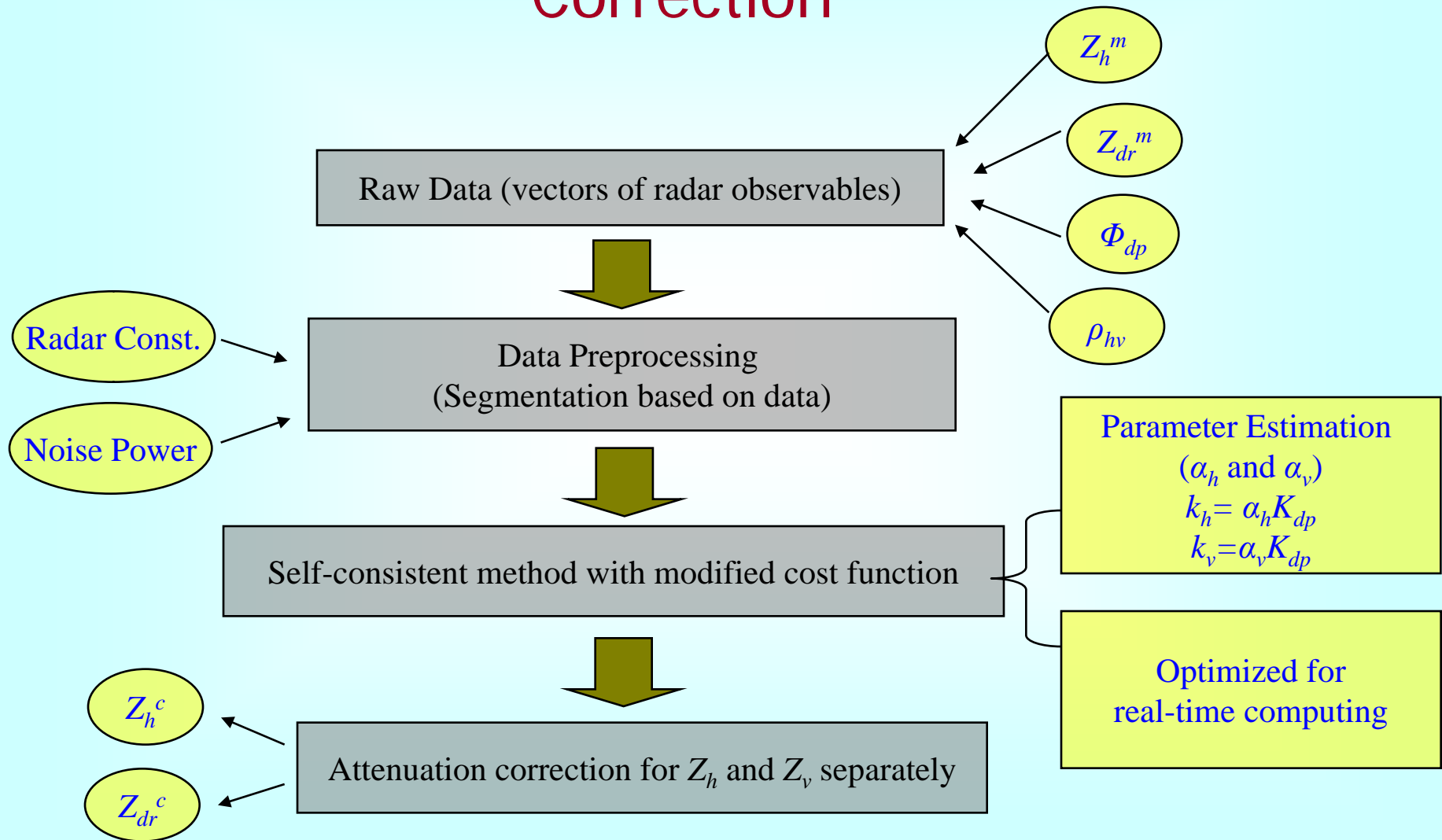
Velocities biased due to clutter



Velocities after clutter filtering

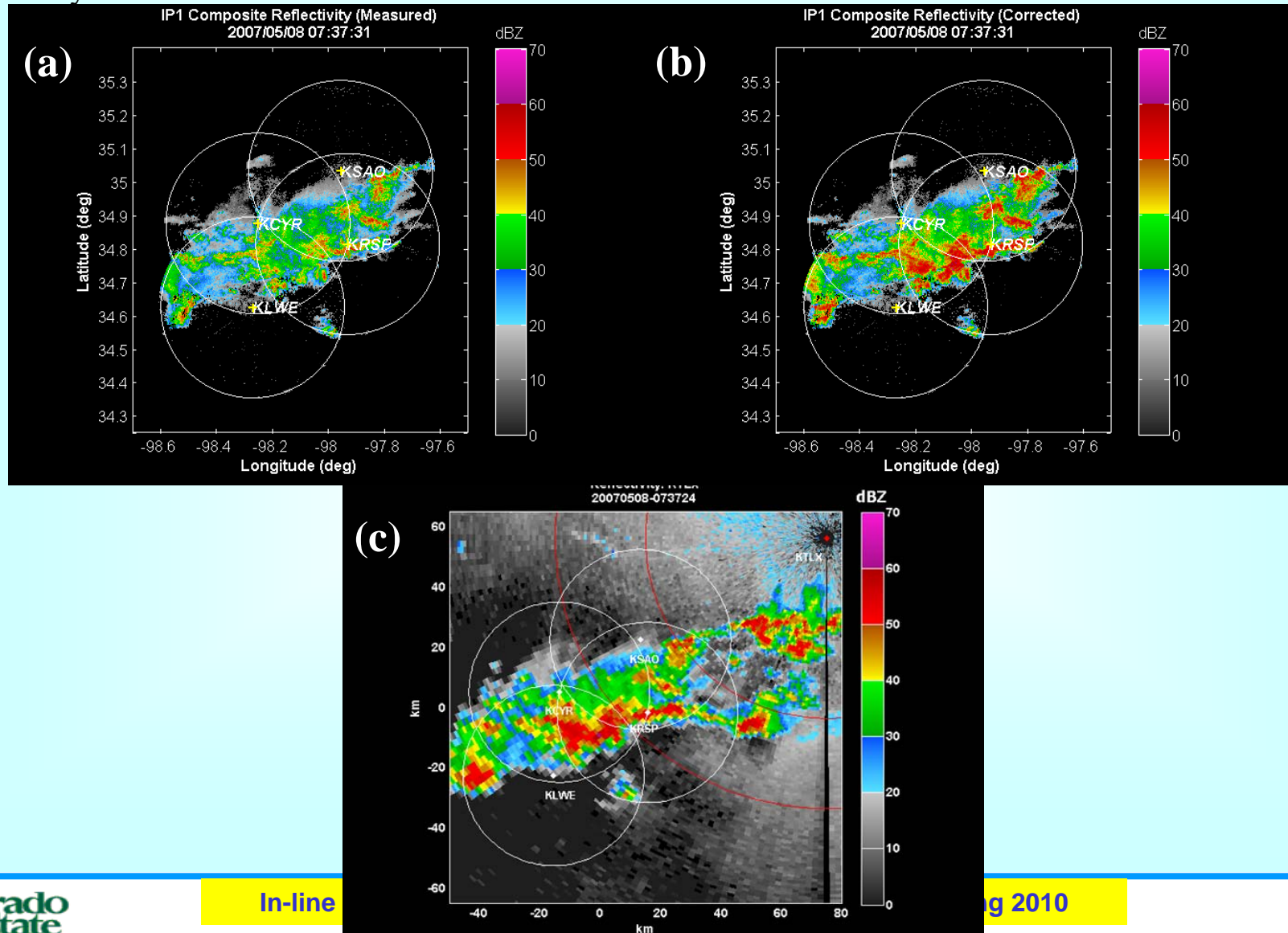
# Attenuation correction

# IP1 Real-time Dual-Pol Based Attenuation Correction



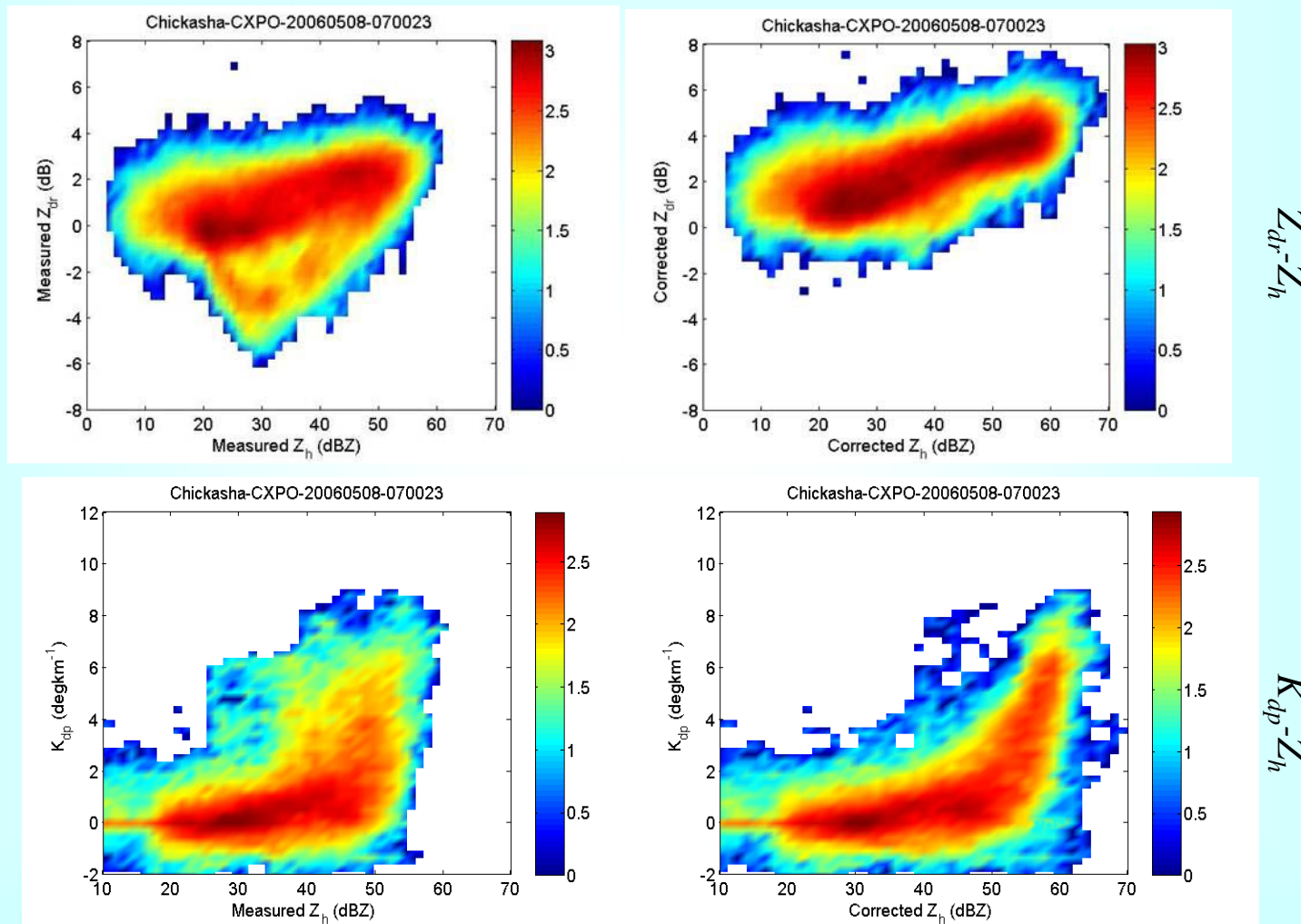
# Dual-Pol Based Attenuation Correction

Reflectivity maps at 07:37:31 May 8 2007 and Nexrad reflectivity map at 07:37:24 May 8 2007. (a) IP1 reflectivity before attenuation correction (b) IP1 reflectivity after attenuation correction (c) Nexrad reflectivity.



# Dual-Pol Based Attenuation Correction

$Z_{dr}-Z_h$  and  $K_{dp}-Z_h$  2D histogram plots for a squall line case on May 8, 2006 before and after attenuation correction.



Before attenuation correction

After attenuation correction

# QPE

# Rainfall Algorithms

$$R(Z, Z_{dr}) = CZ_h^a \mathcal{Z}_{dr}^b \quad \text{or} \quad CZ_h^a 10^{0.1bZ_{dr}(dB)}$$

$$R(K_{dp}) = 129 \left( \frac{K_{dp}}{f} \right)^{b_2}$$

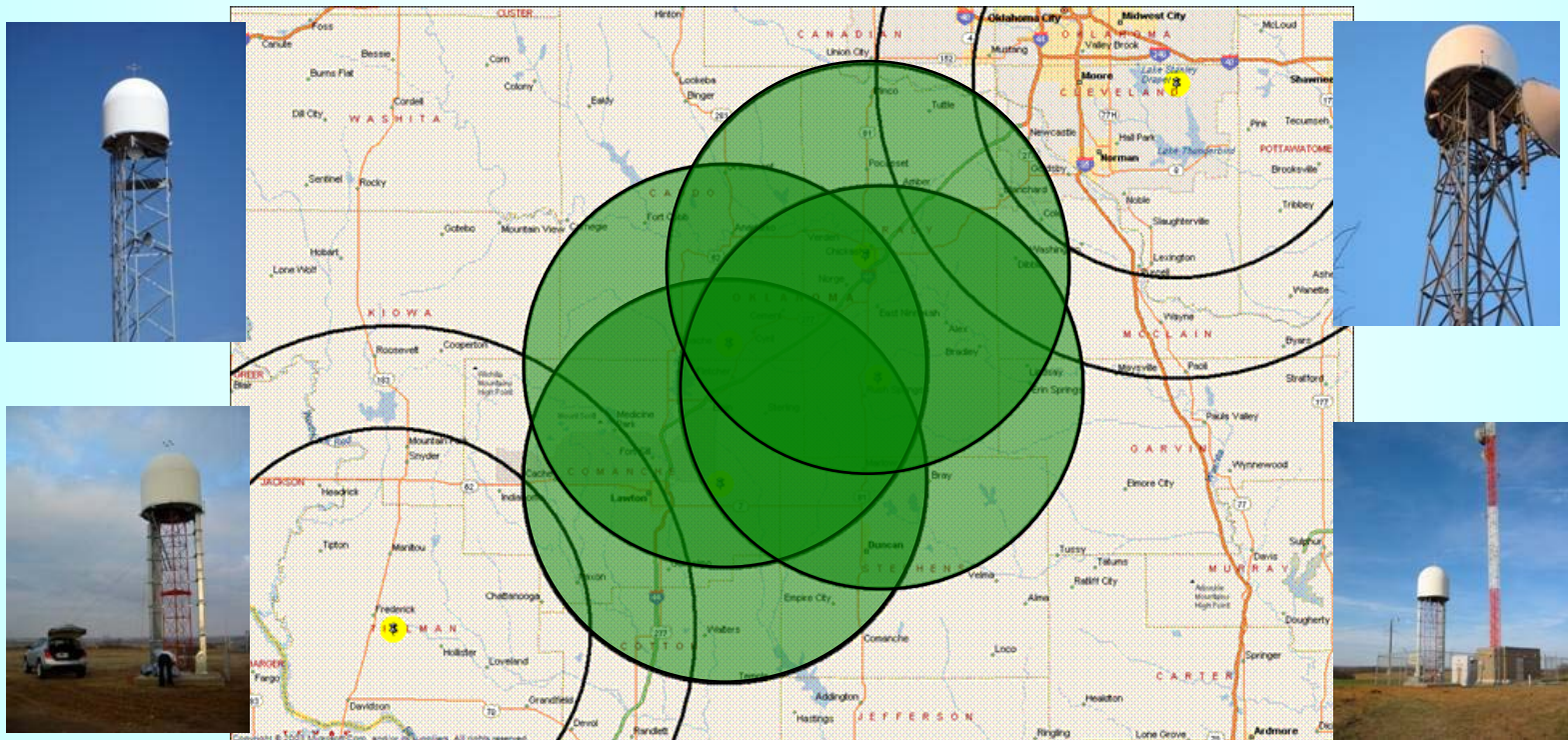
$$R(K_{dp}, Z_{dr}) = C_3 K_{dp}^{a_3} \mathcal{Z}_{dr}^{b_3} \quad \text{mm h}^{-1}$$

The use of  $K_{dp}$  to estimate rainfall has a number of advantages over power measurements:

- independent of receiver and transmitter calibrations,
- unaffected by attenuation,
- relatively immune to beam blockage,
- unbiased by presence of hail or other 'spherical' ice particles in the resolution volume

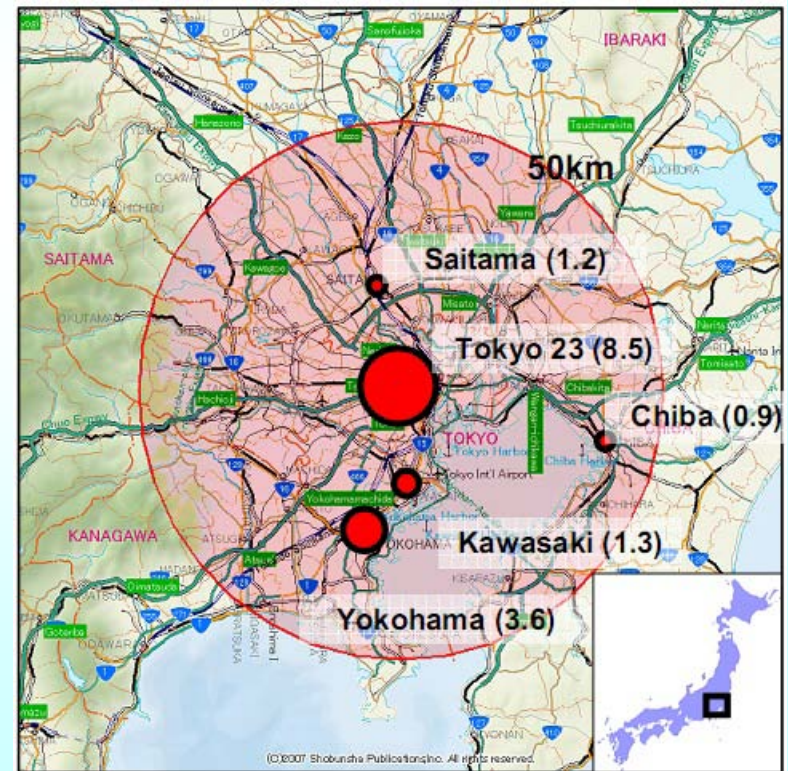
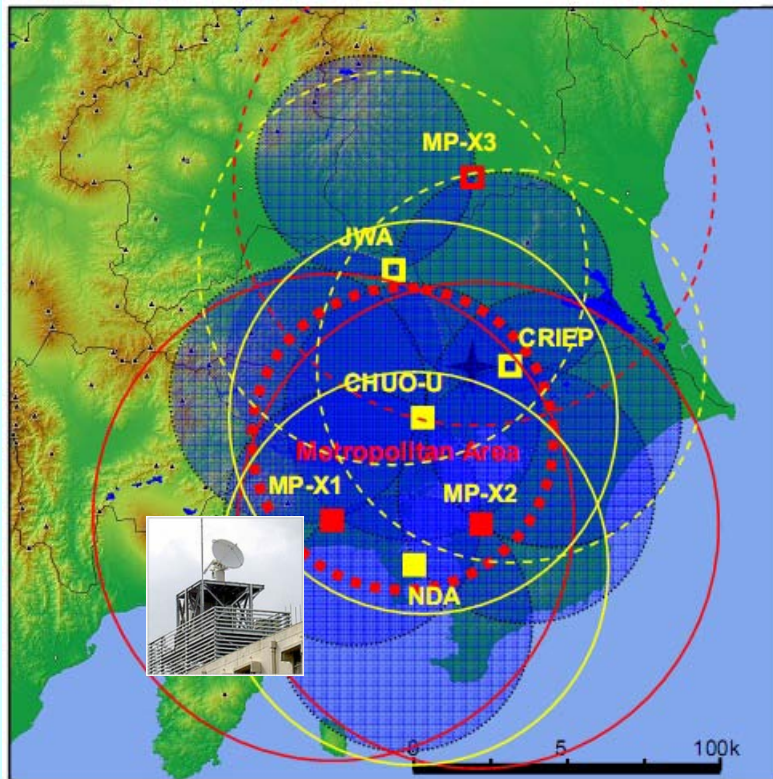


- The radar network
  - Southwestern Oklahoma, ~7000 sq km
  - Four X-band, dual-polarization radar



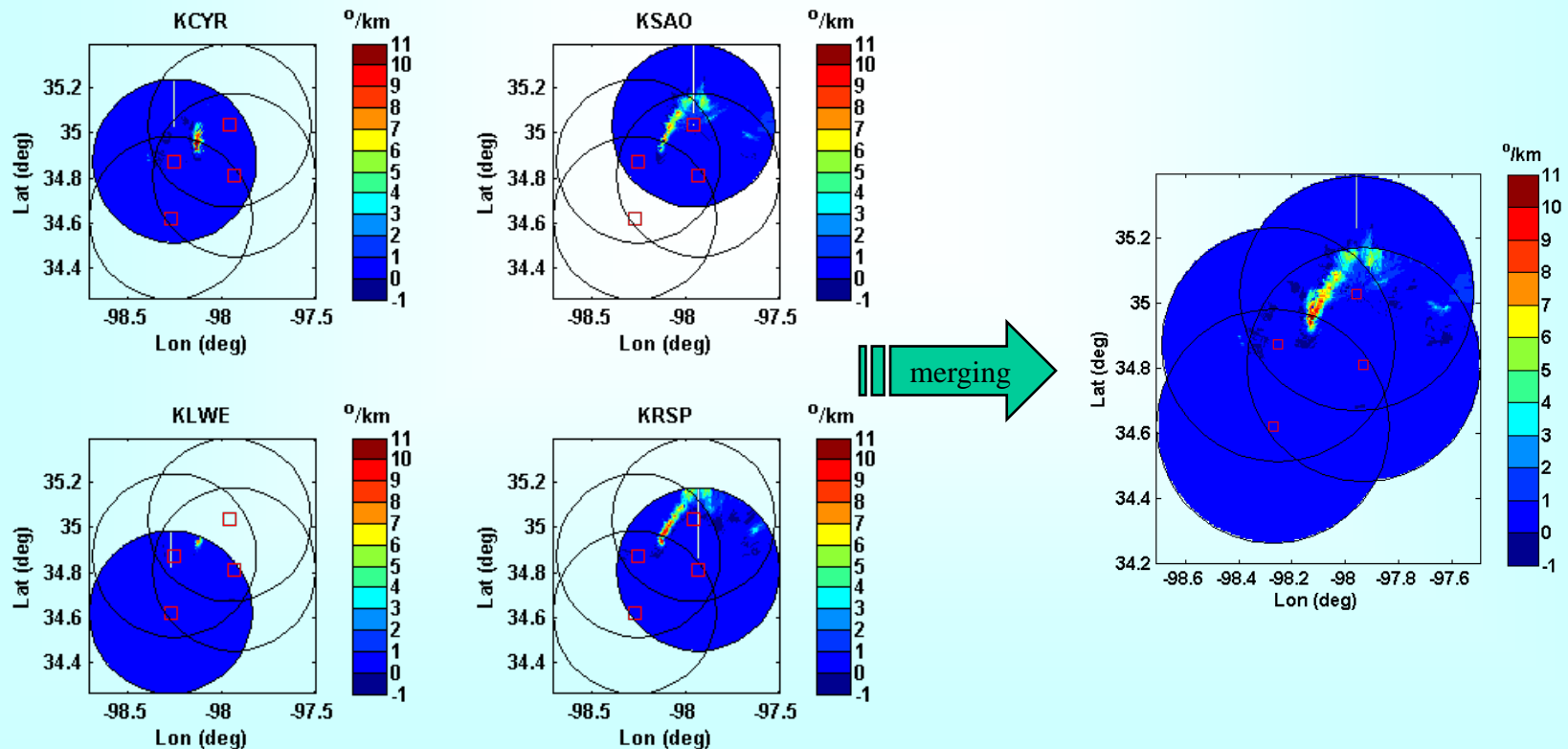
# NIED's X-Net Test Bed

- Led by NIED, National Research Institute for Earth Science and Disaster Prevention, Japan
  - Over one of the most populous and densest metropolitan regions in the world



# $K_{dp}$ and Composition

- $K_{dp}$  based rainfall conversion is attractive at X-band
  - Responds well to low rainfall rate
  - Avoids the uncertainty in attenuation correction
  - Immune to calibration factors across the network





# Rainfall Conversion

- $R$ - $K_{dp}$  based rainfall estimation was implemented in CASA's IP1 test bed.

$$R = 0.6\pi \times 10^{-3} \int v(D) D^3 N(D) dD$$

$$K_{dp} = \frac{\pi^2}{6\lambda} C \int (1-r) D^3 N(D) dD$$

A scaled version of KOUN's rainfall estimation is tested (based on local measured DSDs). \*

$$R = 47.3 K_{dp}^{0.791} \text{ mm/hr}$$

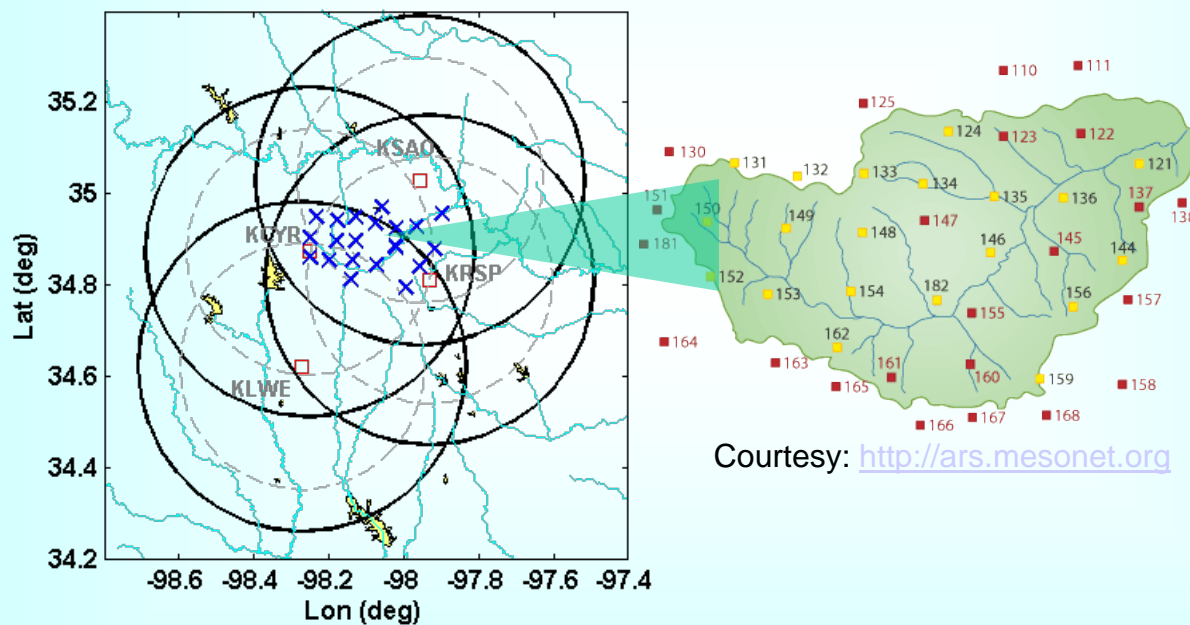


$$R = 18.15 K_{dp}^{0.791} \text{ mm/hr}$$

\* Ryzhkov, A.V., S.E. Giangrande, and T.J. Schuur, 2005: Rainfall Estimation with a Polarimetric Prototype of WSR-88D. J. Appl. Meteor., **44**, 502–515.

# Gauge Comparison

- USDA ARS Micronet – A rain gauge network located at the center of the IP1 test bed



## Little Washita

Size: 611 km<sup>2</sup>

Mean annual precipitation: 760 mm

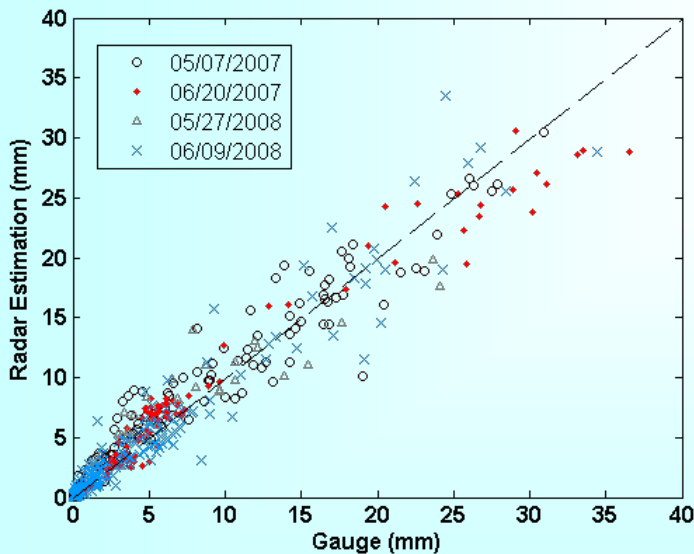
Gauge network: 20 tip-bucket stations

# QPE Evaluation

- Metrics:
  - Normalized bias:
  - Normalized error:

$$\langle e \rangle_N = \frac{\langle R_R - R_G \rangle}{\langle R_G \rangle}$$

$$NSE = \frac{\langle |R_R - R_G| \rangle}{\langle R_G \rangle}$$



Performance of hourly rainfall estimates in comparable weather radar systems using  $K_{dp}$  based QPE algorithms.

Radar System or Network	Total Events Analyzed	NSE (%)
IP1( total )	29	22.76
Instantaneous		42
Hourly		15
MP-X	3	14.8

# QPE Evaluation

- A composite  $R-K_{dp}$  based estimation was implemented in NIED's X-Net test bed.
- X-Net Ground Validation
  - Three events: 2 stratiform, 1 typhoon

$$R = 19.63 K_{dp}^{0.823} \text{ mm/hr}$$

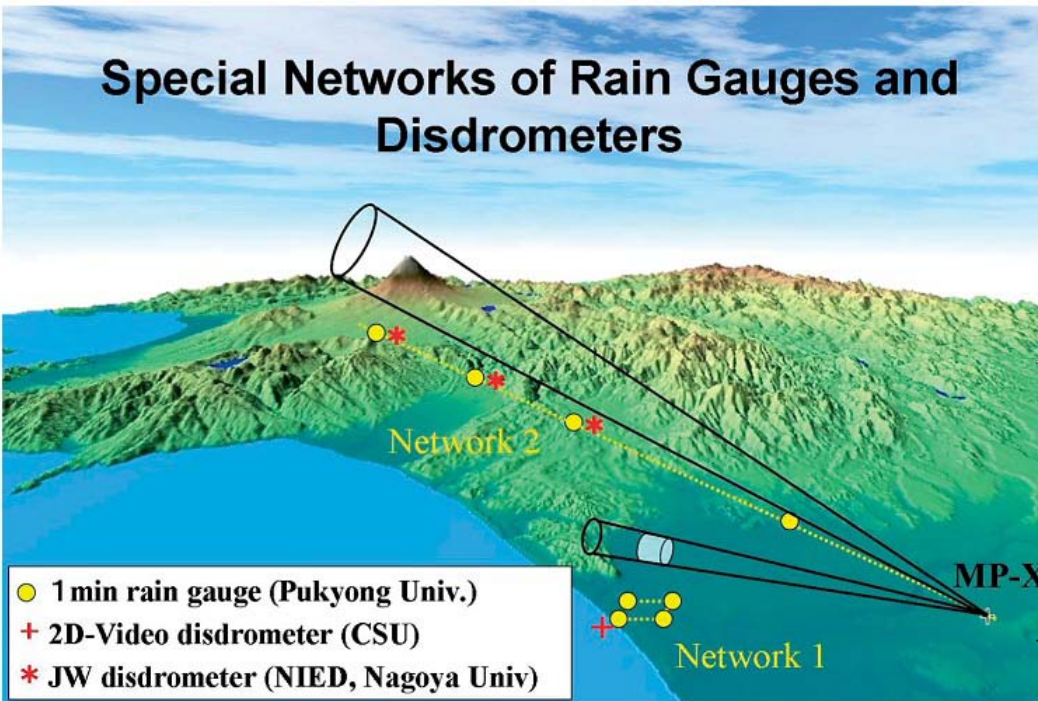
$$K_{dp} > 0.3^\circ / \text{km} \text{ and } Z_h > 35 \text{ dBZ}$$

$$R = 7.07 \times 10^{-3} Z_h^{0.819} \text{ mm/hr}$$

*Also based on locally measured DSDs*

15 min	NE (%)	21.1
	NB (%)	-2.9
1 h	NE (%)	14.8
	NB (%)	-1.1
3 h	NE (%)	11.4
	NB (%)	-1.0

## Special Networks of Rain Gauges and Disdrometers

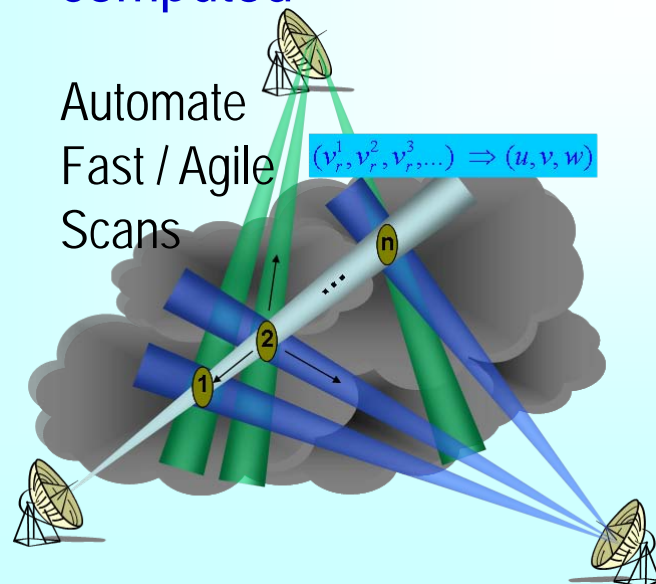


# Dual Doppler



# Dual-Doppler Scan and Retrieval

- Real-time Dual-Doppler wind velocity retrieval system has been developed and installed in IP1, based on proven algorithms and computation tools.
  - 3-D observations from the IP1 radars are gridded and merged, fused into a common analysis grid
  - Both horizontal wind field and vertical wind component are computed



$$\begin{bmatrix} u \\ v \end{bmatrix} = \begin{bmatrix} \sin \phi_1 \cos \theta_1 & \cos \phi_1 \cos \theta_1 \\ \sin \phi_2 \cos \theta_2 & \cos \phi_2 \cos \theta_2 \end{bmatrix}^{-1} \begin{bmatrix} v_r^1 \\ v_r^2 \end{bmatrix} + \text{low level approximation}$$

$$\begin{bmatrix} \sin \phi_1 \cos \theta_1 & \cos \phi_1 \cos \theta_1 \\ \sin \phi_2 \cos \theta_2 & \cos \phi_2 \cos \theta_2 \end{bmatrix}^{-1} \begin{bmatrix} \sin \theta_1 \\ \sin \theta_2 \end{bmatrix} (w + w_t)$$

# Dual-Doppler Scan and Retrieval

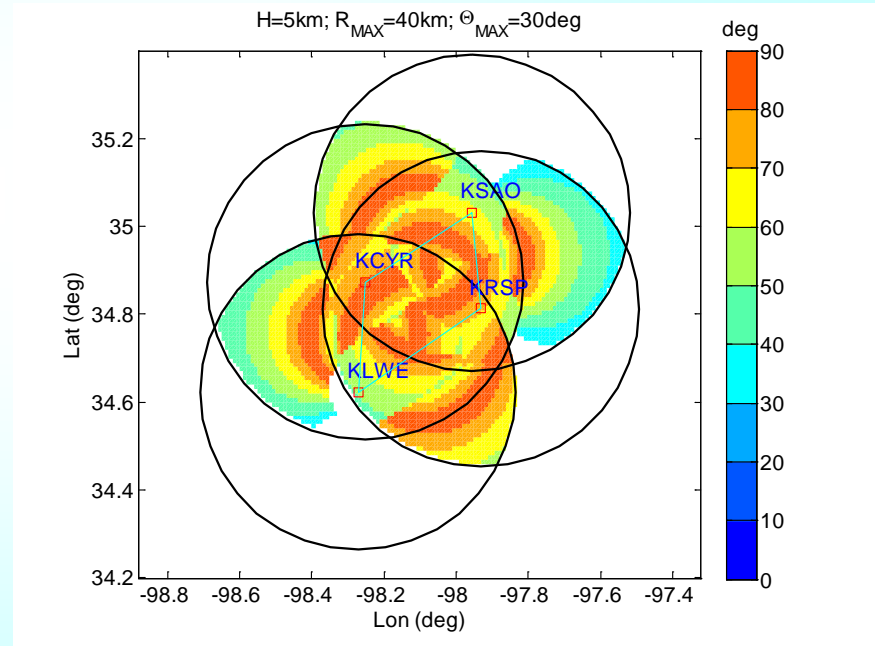
- The upgrades over the last two storm seasons improved the capability of the IP1 radar network for dual-Doppler wind observations

Error assessment at lower altitudes

$$\frac{\sigma_{u'}^2 + \sigma_{v'}^2}{2\sigma_{v_r}^2} = \frac{1}{\sin^2(\phi_1 - \phi_2)} \frac{\cos^2 \theta_1 + \cos^2 \theta_2}{2\cos^2 \theta_1 \cos^2 \theta_2}$$

Unique Capabilities

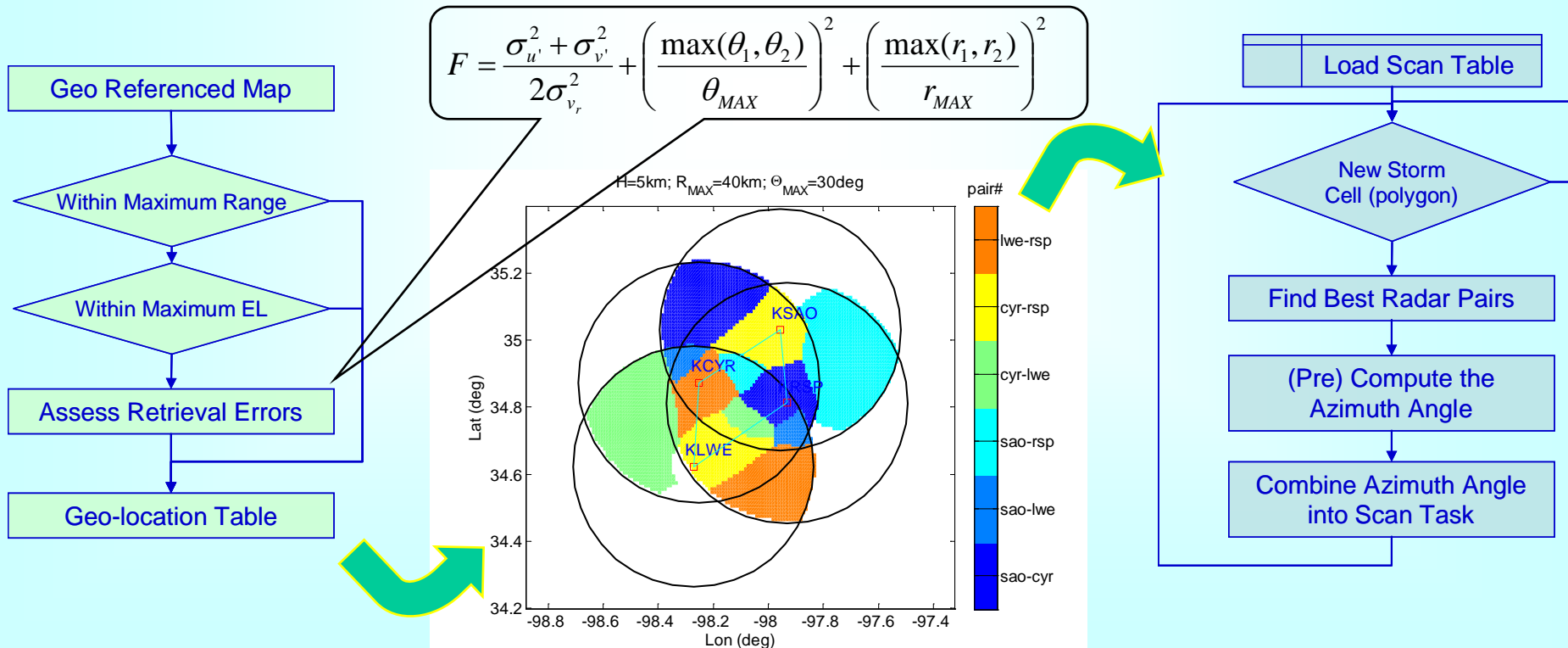
- ❖ Real-time processing and control

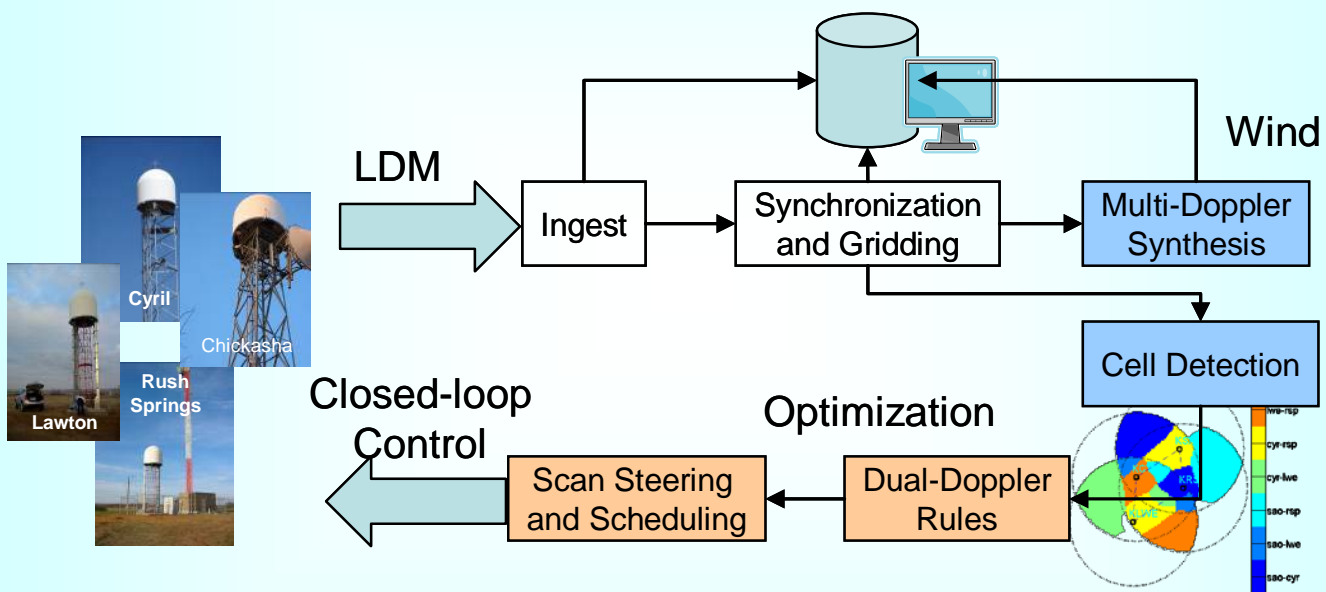


Best Dual-Doppler Angles in the IP1 Coverage at 5 km AGL

# Dual-Doppler Scan and Retrieval

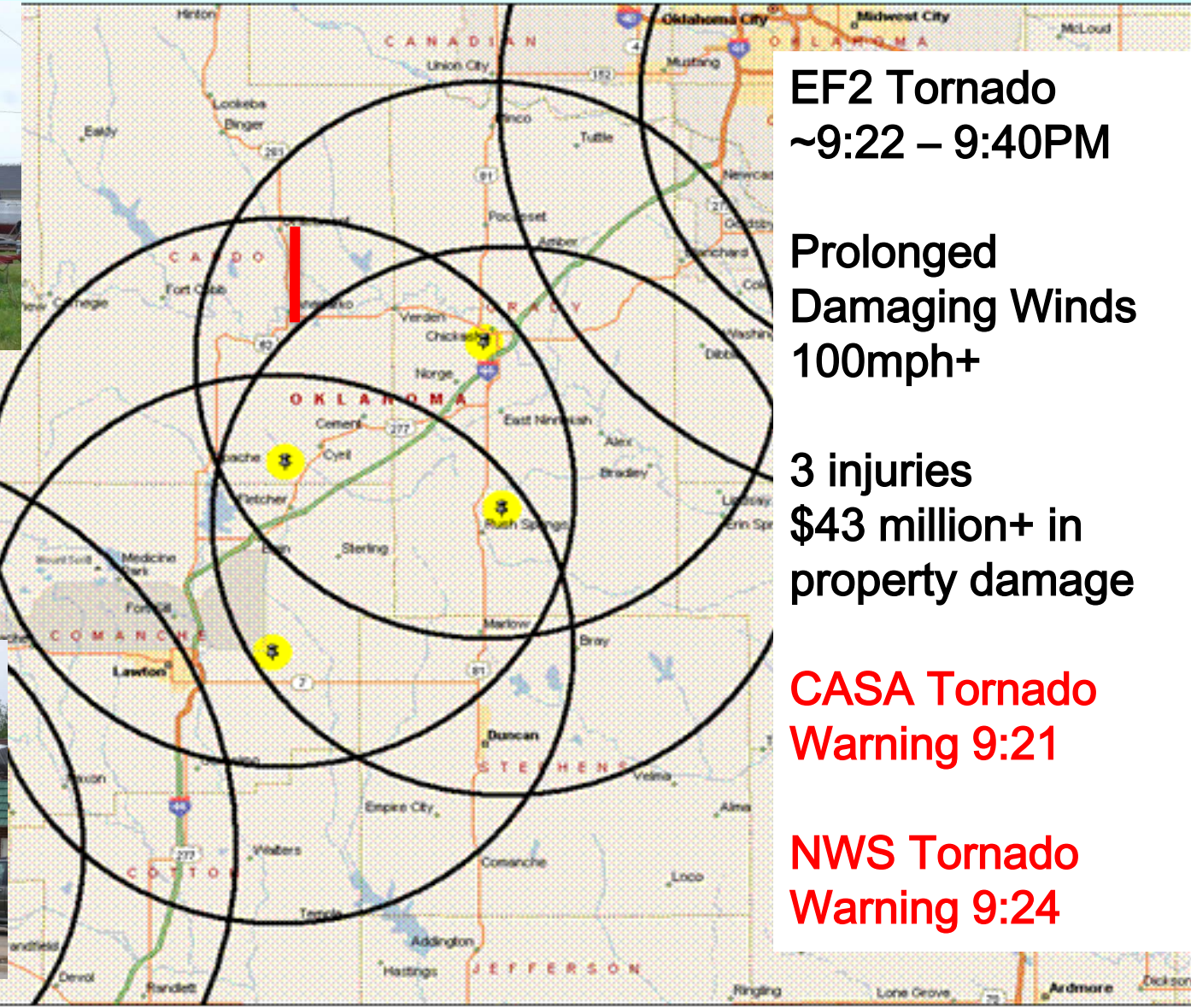
- Dual-Doppler scan strategy
  - Best beam crossing angle for Dual-Doppler retrieval;
  - Lowest elevation angle to reduce scan interval (tilts);
  - Closest range to efficiently use the power budget.







# Anadarko Tornado and Damaging Winds Event – May 13, 2009



**EF2 Tornado**  
**~9:22 – 9:40PM**

**Prolonged**  
**Damaging Winds**  
**100mph+**

**3 injuries**  
**\$43 million+ in**  
**property damage**

**CASA Tornado**  
**Warning 9:21**

**NWS Tornado**  
**Warning 9:24**

# IP1 Radar Network Dual-Doppler Tornado Observation: 2009-May-14

02:28:07 UTC

02:29:07 UTC

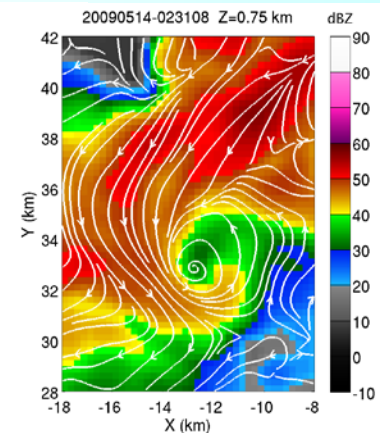
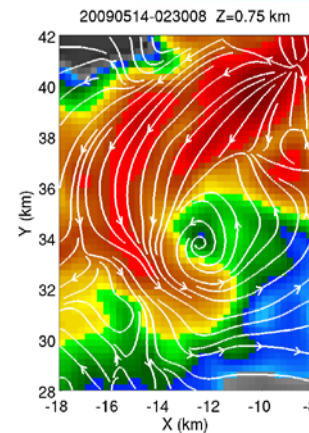
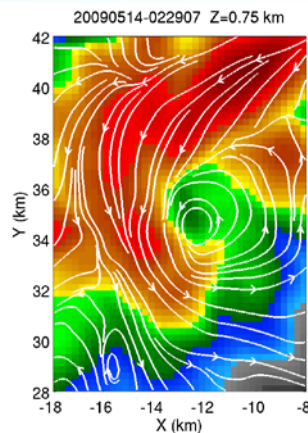
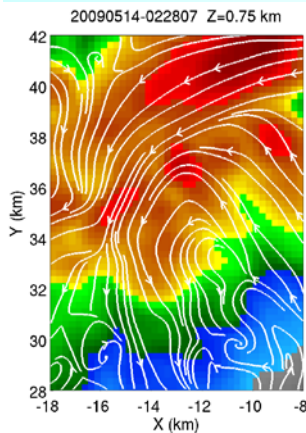
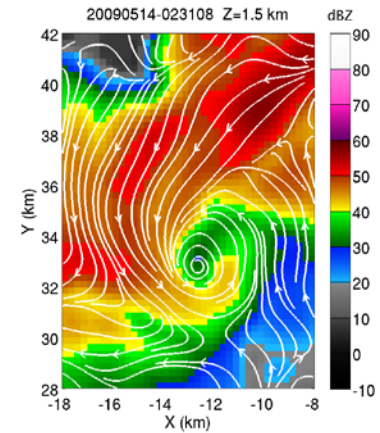
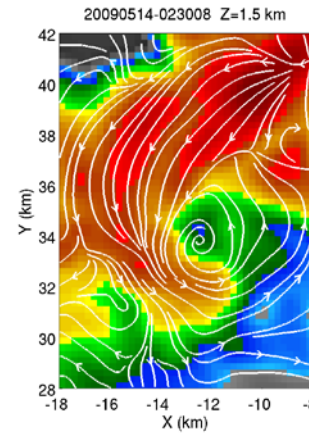
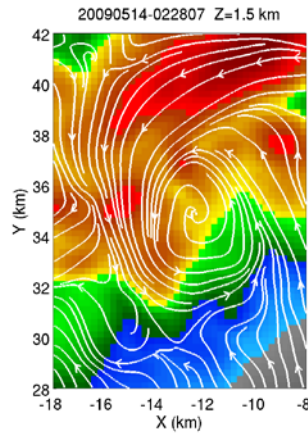
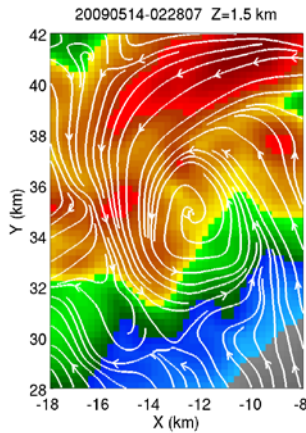
02:30:08 UTC

02:31:08 UTC

Altitude (AGL, km)

1.1

0.35



Time



# IP1 Radar Network Dual-Doppler Tornado Observation: 2009-May-14

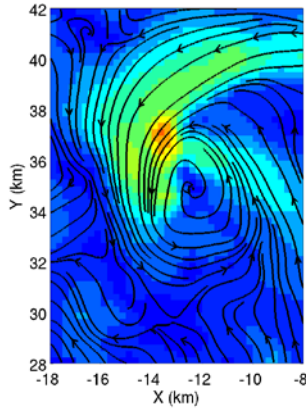
02:28:07 UTC

02:29:07 UTC

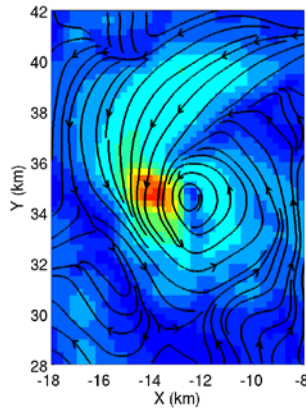
02:30:08 UTC

02:31:08 UTC

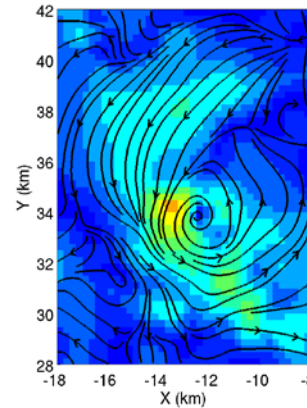
20090514-022807 Z=1.5 km



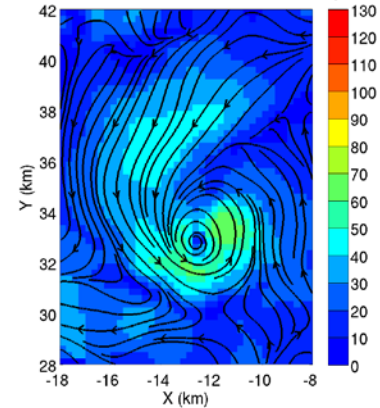
20090514-022907 Z=1.5 km



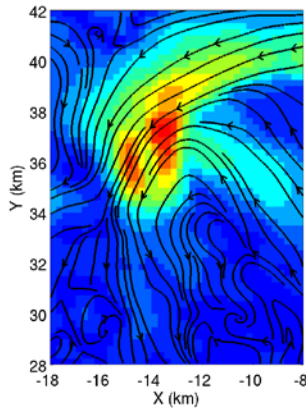
20090514-023008 Z=1.5 km



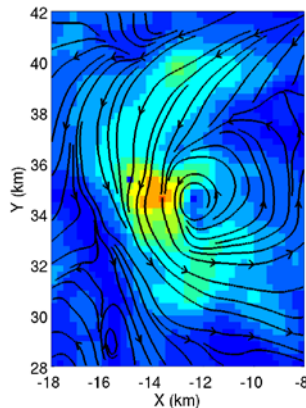
20090514-023108 Z=1.5 km



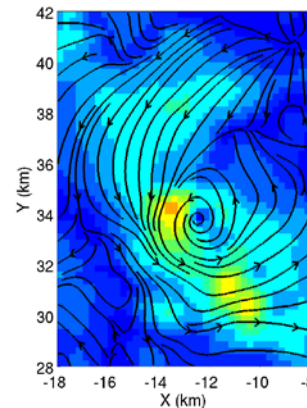
20090514-022807 Z=0.75 km



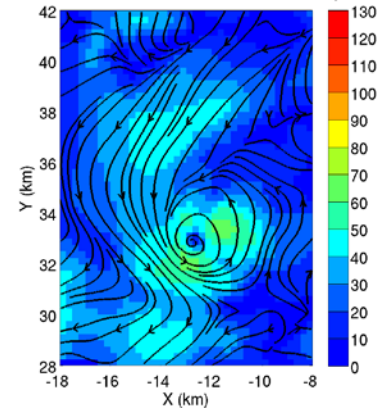
20090514-022907 Z=0.75 km



20090514-023008 Z=0.75 km



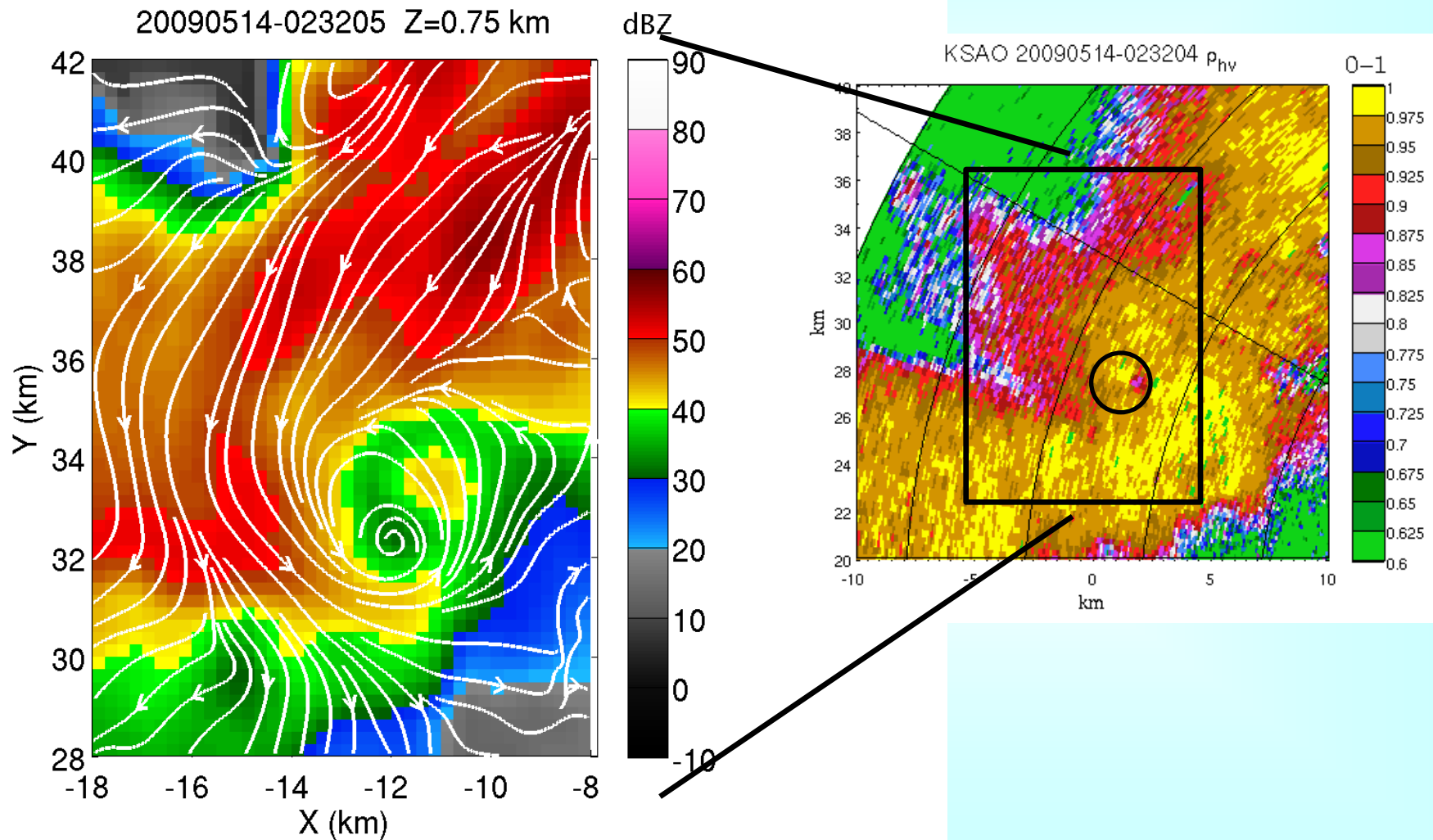
20090514-023108 Z=0.75 km



Time

# IP1 Radar Network Dual-Doppler Tornado Observation: 2009-May-14

## Drop in co-polar correlation due to debris





# Retrievals Products that are fairly stable and can be considered operationally viable

- ❑ Quantitative Precipitation Estimation ( 2D )
- ❑ Drop Size Distribution ( Quasi 3D )
- ❑ Water content ( 2 D )
- ❑ Hydrometeor Classification ( 3 D )
- ❑ Dual –Doppler based products such as Vorticity ( 2 D )

*Quality of all these products depends on the quality of in-line processing*

# Thank you