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### Summary

Atmospheric aerosols affect climate and human health, but their source and chemical evolution is poorly understood. Organic aerosols (OA) constitute a major fraction of submicron atmospheric aerosols and are either emitted directly from the source (primary organic aerosol, POA) or formed from oxidation of precursor organic gases in the atmosphere (called secondary organic aerosols, SOA),<sup>1</sup> which is often a major fraction of OA.<sup>2</sup> Despite this progress, SOA formation mechanisms and dominance in atmospheric aerosols are poorly simulated by models. In collaboration with Aerodyne Research Inc., the UW group has been developing a new instrument to measure a suite of multifunctional organic compounds in both the gas and particle phases simultaneously. Here I present recent laboratory and field results obtained using this new instrument. We investigated the SOA yields, formation, reaction mechanisms and products, and gas-particle partitioning of the products from photo-oxidation of several alkanes in the Caltech environmental chamber. Previous smog-chamber studies<sup>3,4</sup> have shown that the oxidation of alkanes has a high SOA yield and can produce highly oxygenated SOA. We also took part in a recent field study to examine wintertime urban outflow in the UK, known as "Aeroflo/Clearflo" which occurred January – February, 2012. Preliminary data from that campaign are shown here and discussed in a breakout session.

# **New Instrumentation**

# MOVI-HRToF-CIMS Technique

- Pre-concentration and gas-particle separation using inertial impaction
- Temperature programmed desorption for volatility information
- Chemical ionization high-resolution time of flight mass spectrometry

## Micro-Orifice Volatilization Impactor<sup>5</sup>

100 nozzles, 150  $\mu$ m diameter; Cutpoint ( $d_{50}$ ): 0.13 µm; Flow: 10 L/min,  $\Delta P = 15\%$ 

Thermally isolated impaction surface; surface and body temperature controlled to within +/- 1K

## Chemical Ionization

Proton abstraction – selectively reacts with acids<sup>6</sup>  $CH_3C(O)O^- + RC(O)OH \rightarrow CH_3C(O)OH + RC(O)O^-$ 

Proton transfer – general ionization scheme  $H_3O(H_2O)_n^+ + RC(O)R' \rightarrow (n+1)H_2O + RC(O)R'H^+$ 

# High Resolution Time-of-Flight Mass Spectrometer



particle-phase carboxylic acids.

power as a function of m/Q (b), and an example ion resolution and composition assignment (c).







# SOA and OH maintained at ambient levels Mass Spectra – Positive Mode Gas-phase and the desidence of the later Like ike the ite is a second Particle-phase and the state of t

• Experiments used 28 m<sup>3</sup> Caltech chambers

and HONO + hv as OH source, respectively

Neutralized ammonium sulfate seed aerosol

• Dry conditions, RH < 7% and T = ambient

• Oxidation time scales of 18 and 36 hr



Fig. 3 (above): SOA mass concentration and yield during the 18hr oxidation of fractional Dodecane under Low-NOx conditions. SOA mass concentration reached 25 µg m<sup>-3</sup> by the end of the 18 hr oxidation. The maximum fractional yield is approximately 0.12. OH concentrations were 2 x  $10^6$  molec cm<sup>-3</sup>.

Fig. 4: Evolution of mass spectra in gas and particle-phases as a function of experiment time are presented here. Note the split in the mass axis and the signal on the right axis are multiplied by a factor of 500 and 50 relative to the left axis

Oxygenated products up to 450 Th are detected. There are differences between gas and particle-phase products, but the major products, 185 Th, 199 Th, and 203 Th are observed in both phases. The distribution and molecular composition of products evolves over 18hr, suggesting multiple generations arise from functionalization and fragmentation

# **Product Formation and Evolution**



mix of different generations of products.



1 22 23 24 25 26 Hours Since Lights On

Fig. 6: High resolution analysis shows multi-generational products at the same opposing time series (panels at right). A mix of "functionalization" (O – addition to backbone) and "fragmentation" (C loss from backbone) is apparent with an increasing O/C ratio as aging progresses.



# Towards a Molecular-Level Understanding of Organic Aerosol Evolution Using a Field-**Deployable High Resolution Time-of-Flight Chemical Ionization Mass Spectrometer**

**Dodecane Low-NO<sub>x</sub> 18hr Oxidation** • Low and high-NO<sub>x</sub> conditions, using  $H_2O_2$  + hv xperiment time (hr Experiment Time (hr)



Fig. 7 The Aeroflo field site was located ~45 km ESE of London in Kent (Detling). The area is rural and the meteorology during the campaign provided a range of source regions, air mass ages, temperature, and precipitation. The MOVI-HRToF-CIMS operated nearly continuously from January 22 – February 15.

# Preliminary Results: MOVI-HR-ToF-CIMS Time Series



products in gas and particle-phases. The MOVI-HRToF-CIMS is able to resolve differences in the time evolution of multigenerational products that contribute to SOA. Future analyses will address functionalization vs. fragmentation and volatility information obtained with MOVI (not shown). In addition, the MOVI-HRToF-CIMS successfully measured a large suite of compounds in the gas and particle phases during the Aeroflo deployment that will likely give unprecedented information on organic aerosol sources and evolution.

# **Acknowledgements / References**

07-Earth07F-0069

### **Field Measurements: Aeroflo/Clearflo U.K.**

DOE-ASR funded Aeroflo – an aerosol focused augmentation of the U.K. Clearflo field study of urban air quality and outflow in the U.K. A multi-institution team deployed a state-of-theart suite of in-situ and remote sensing instrumentation to the sources study and evolution of aerosol downwind of London and the European Continent from January 10 -February 15, 2012.



Figure 8. The three panels at left show a very small subset of ion signal time series recorded by the MOVI-HR-ToF-CIMS including gas and particle components detected using the acetate reagent ion (top and middle), and proton-transfer by  $H_3O^+(H_2O)_n$  clusters.



In the top panel, two ions (m/Q 138 and 170) detected analyzing the particle phase showed significant variations over the campaign that were fairly well correlated with AMS total organic aerosol mass. These ions also correlated well with that at m/Q 42 detected almost entirely in the gas-phase corresponding to isocyanic acid (HNCO), a biofuel combustion tracer. The ability to simultaneously monitor gas and particle phase composition with one instrument enhances the ability for robust source identification.

The middle and bottom panels illustrate that there are a number of ions detected in the particle phase that correlate well with AMS total organic aerosol mass loading (many not shown). Statistical techniques are to decompose these beina used multidimensional data into classes for a detailed source apportionment of organic aerosol in this study.

## High-Resolution Analysis of m/Q 138 During Sampling and Desorption

Fig. 9 As a final demonstration of the capabilities of the instrument, at left is plotted a high-resolution time series of ion signal with unit m/Q of 138 (see Fig. 8 above). Two different ions compose 37.8 137.85 137.9 137.95 138 138.05 138.1 138.15 138.2 the signal at this m/Q. During sampling, Fitted Data the ion at 137.96 Th dominates the During temperaturesignal. programmed desorption (6:25AM), the ion signal becomes dominated by that corresponding to a nitrophenol. This finding is consistent with the particle signal of m/Q 138 being correlated with a biofuel combustion tracer.

DOE ASR Program, Aerodyne Research Incorporated and AEROFLO COLLABORATORS! Also DOE SBIR, NSF CAREER ATM-0846183, NASA ES

1. Kanakidou, M., et al., (2005), Organic aerosol and global climate modeling: a review. Atmos. Chem. Phys., 5, 1053-1123. 2. Jimenez, J.L., et al., (2009), Evolution of organic aerosols in the atmosphere. Science, 326, 1525-1529. 3. Lim and Ziemann, (2009), Chemistry of secondary organic aerosol formation from OH radical-initiated reactions of linear, branched, and cyclic alkanes in the presence of NOx. Aerosol Sci. Tech., 43, 604-619. 4. Presto, A., et al., (2010), Secondary organic aerosol formation from high-NOx photo-oxidation of low-volatility precursors: n-Alkanes. ES&T 44, 2029 2034 5. Yatavelli, R. L. N., et al., (2010), Particulate organic matter detection using a micro-orifice volatilization impactor coupled to a chemical ionization mass spectrometer (MOVI-CIMS). Aerosol Sci. Tech., 44, 61-74. 6. Veres, P., et al., (2008), Development of Negative-Ion Proton-Transfer Chemical-ionization Mass Spectrometry (NI-PT-CIMS) for the Measurement of Gas-Phase Organic Acids in the Atmosphere. Int. J. Mass. Spec., 274, 48-55. 7. DeCarlo, P.F., et al., (2006), Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer. Anal. Chem., 78, 8281-8289. 8. Kroll, J., et al., (2011), Carbon oxidation state as a metric for describing the chemistry of atmospheric organic aerosol. Nature Chem., 3, 133-139