

Modeling Radiative Impact of Aerosols over S Asia Constrained by Observation of Vertical Distribution

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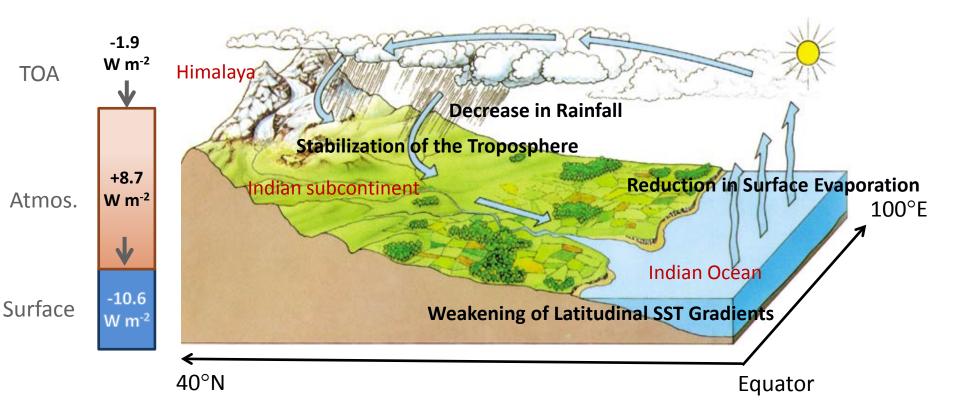
Argonne National Laboratory

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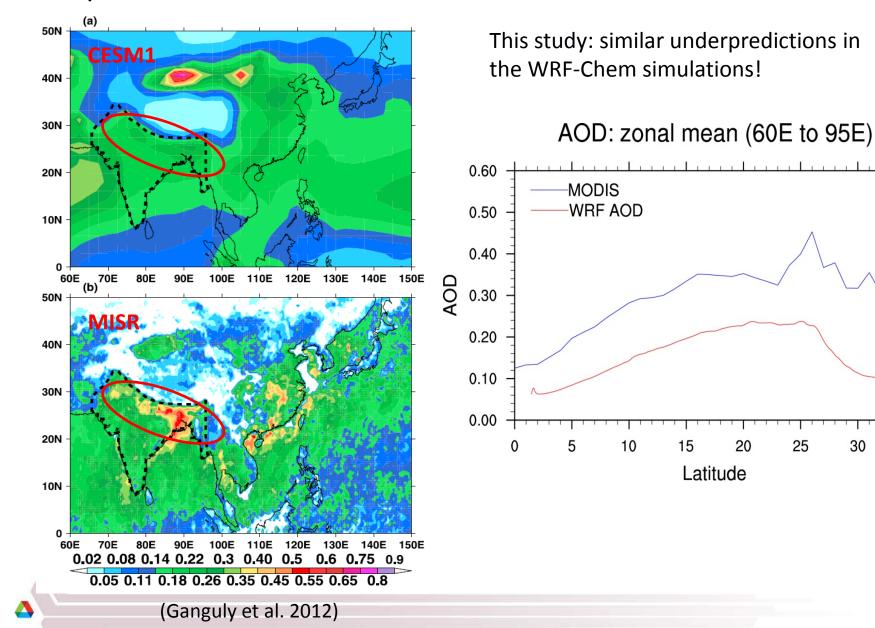


High aerosol loadings in S Asia influence the regional energy balance and hydrological cycle

Regional hydrological cycle



Column aerosol optical depth (AOD) underestimated by models in comparison with satellite data



Questions

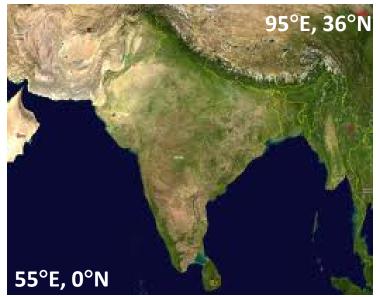
- How does the simulated aerosol extinction profile relate to the underestimated column AOD?
- How does the uncertainty in aerosol vertical distribution affect the simulated atmospheric heating and thermodynamics?

WRF-Chem Specification

- WRF/Chem 3.3 with modifications
- **Domain**: 55E ~ 95E and 0 ~ 36N
- Grid size:12km x 12km
- Vertical layers: 27
- Chemistry: MOZCART
 - MOZART gas-phase chemistry
 - GOCART aerosols:
 - Sulfate
 - BC and OC (hydrophilic and hydrophobic)
 - Dust (0.5, 1.4, 2.4, 4.5, and 8 μm in effective radius)
 - Sea salt (0.3, 1.0, 3.2, and 7.5 μm)
 - No SOA

Aerosols feedback to radiative transfer

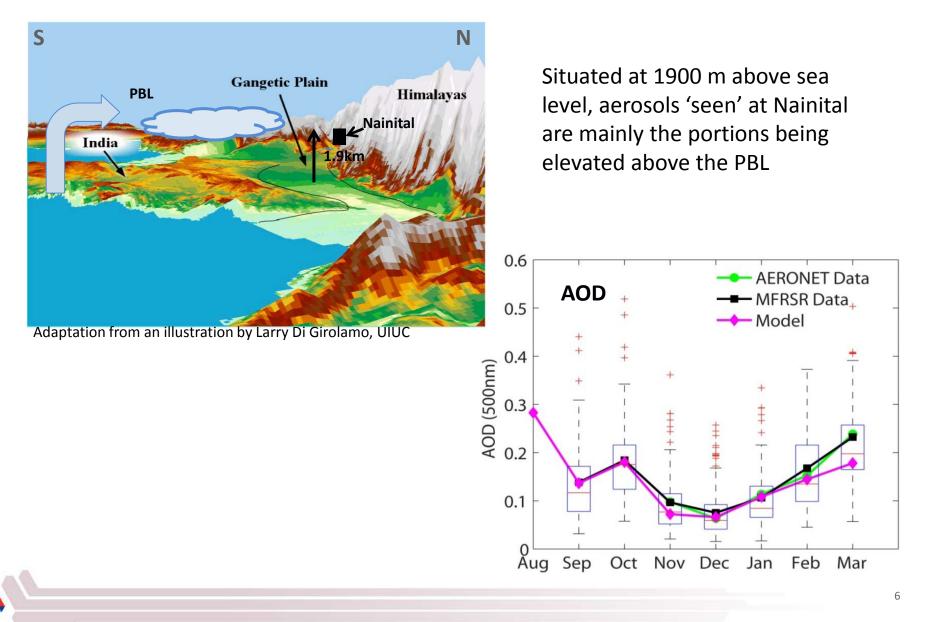




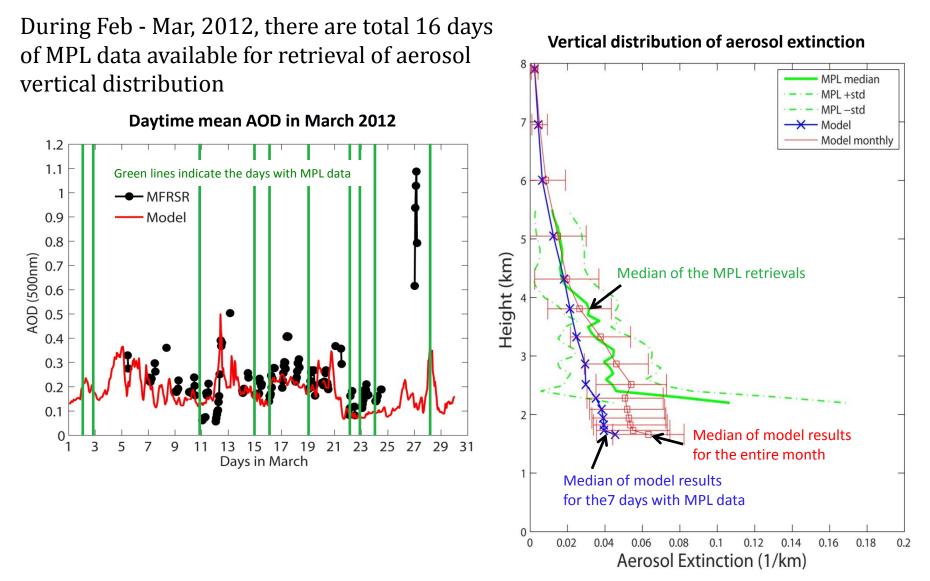
- Anthropogenic emissions: SO₂, BC, and OC for year 2010 (Lu et al., 2011); others from EDGAR;
- Physics/Dynamics:
 - RRTMG for shortwave and longwave radiation
 - MYJ TKE PBL scheme
 - Thompson cloud microphysics
 - CAM5 cumulus parameterization

Ganges Valley Experiment (GVAX)

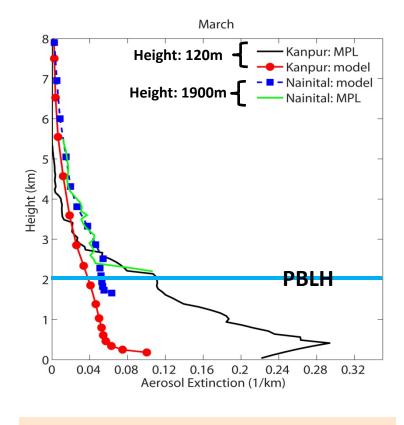
Surface measurements of aerosols were made from Aug, 2011 to Mar, 2012 at Nainital



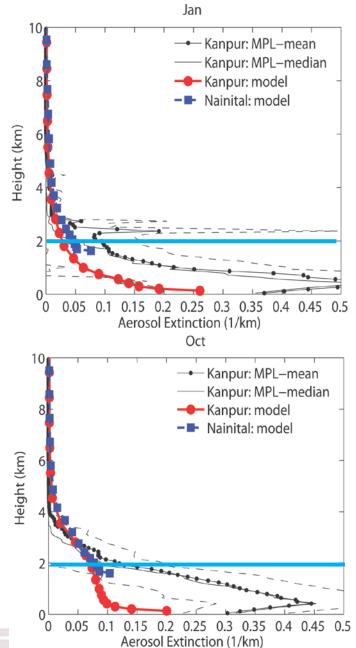
Nainital/GVAX: vertical profile of aerosol extinction



High-altitude vs. low- altitude locations

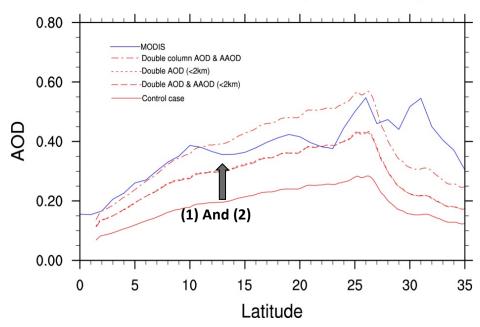


Underestimation of the column AOD is mainly due to the underestimation in the lower model levels, especially in the PBL



Sensitivity studies constrained by observations

- ✓ Control run
- ✓ Sensitivity experiments:
 - 1) Double AOD below 2km; same SSA (increasing scattering and absorbing aerosols);
 - 2) Double AOD below 2km; larger SSA (increasing scattering aerosols only);
 - 3) Double AOD and AAOD in the entire column (Included in the Poster)



AOD: zonal mean (60E to 95E)

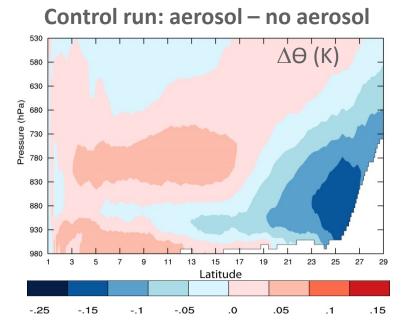
Aerosol direct radiative forcing

Regional mean (W m ⁻²)	Control run	Double AOD (<2km), same SSA	Double AOD (<2km), larger SSA
ΤΟΑ	-3.0	-4.9	-5.4
Atmosphere	+6.3	+9.3	+6.3
Surface	-9.3	-14.2	-11.7



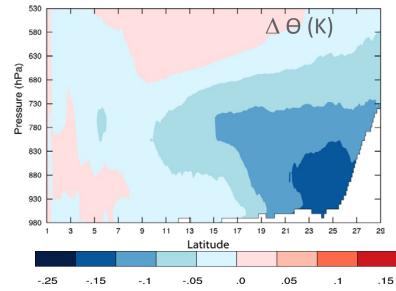
Changes in potential temperature

 $Tendency_{\Theta} = [radiation] + [microphysics] + [pbl] + [cumulus] + [mixing, diffusion]$

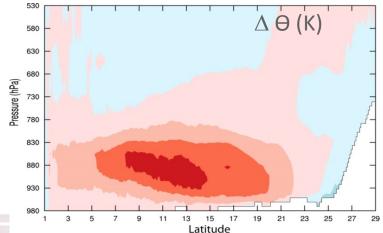


- Aerosol-radiation interactions lead to different responses in potential temperature over land and over the ocean
- With same increase in AOD, potential temperature is enhanced by increasing both scattering and absorption (same SSA), while reduced by increasing aerosol scattering only (larger SSA)

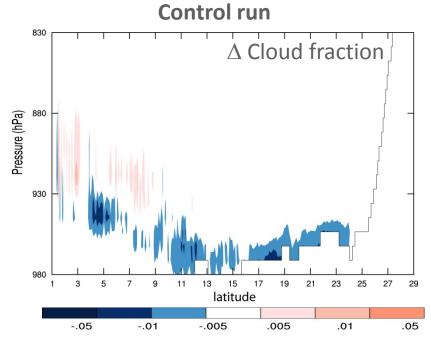
Double AOD (<2km) larger SSA - Control



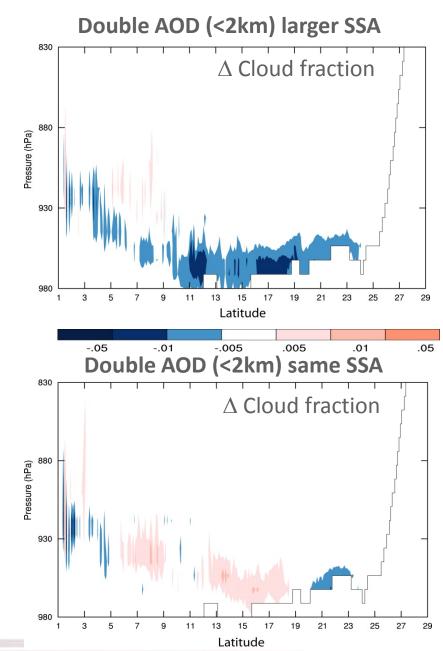
Double AOD (<2km) same SSA - Control



Impact on cloud cover



- Aerosol-ratiation-cloud interactions lead to reductions in cloud cover
- With same increase in AOD, cloud cover is increased by increasing both scattering and absorption (same SSA), while reduced by increasing aerosol scattering only (larger SSA)



Conclusion

- Comparison with in surface aerosol measurements and extinction profiles suggest that underestimation of AOD in S Asia is mainly due to aerosol underpredictions in PBL
- The resulting responses in the potential temperature and cloud cover by constraining the aerosol extinction profiles suggest that increase of absorbing aerosols in PBL enhances convection (heating/more cloud cover), as scattering aerosols stabilize the lower atmosphere (cooling/less cloud cover).
- To understand aerosol-radiation-cloud interactions, we need more quality data to evaluate and improve vertical distribution of aerosol extinction and absorption simulated by models.