

Towards Understanding and Parameterizing Cold Pools over tropical Oceans

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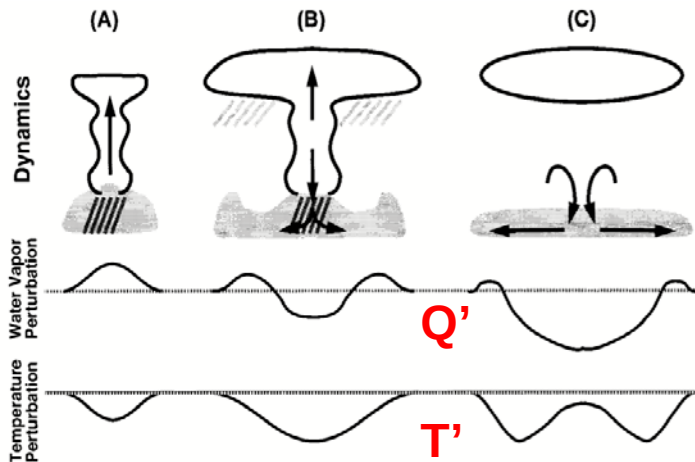
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2014 ASR Science Team Meeting

Why are cold pools important?

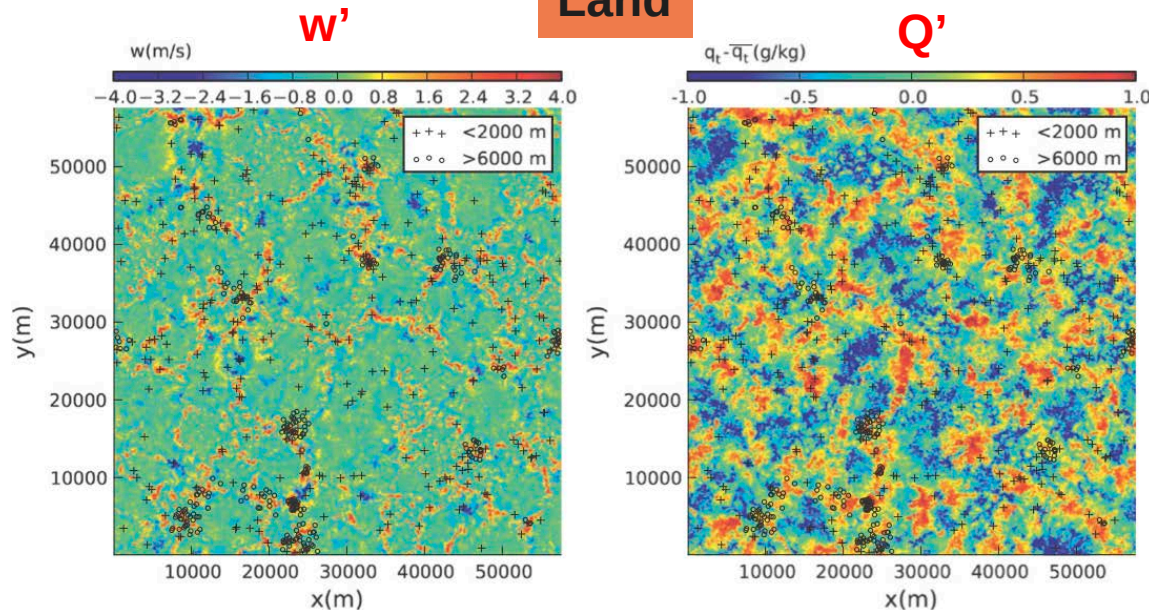
- ▶ An important mechanism for organized convection
 - Accelerates transition from shallow-to-deep convection [Khairoutdinov and Randall 2006]
 - Deep convection preferentially develops at the edge of cold pool (moist & unstable; forced ascend at gust front) [Tompkins 2001; Lima and Wilson 2008; Boing et al. 2012]
- ▶ Their effects are not treated in most GCMs

Oceanic



Tompkins (2001)

Land



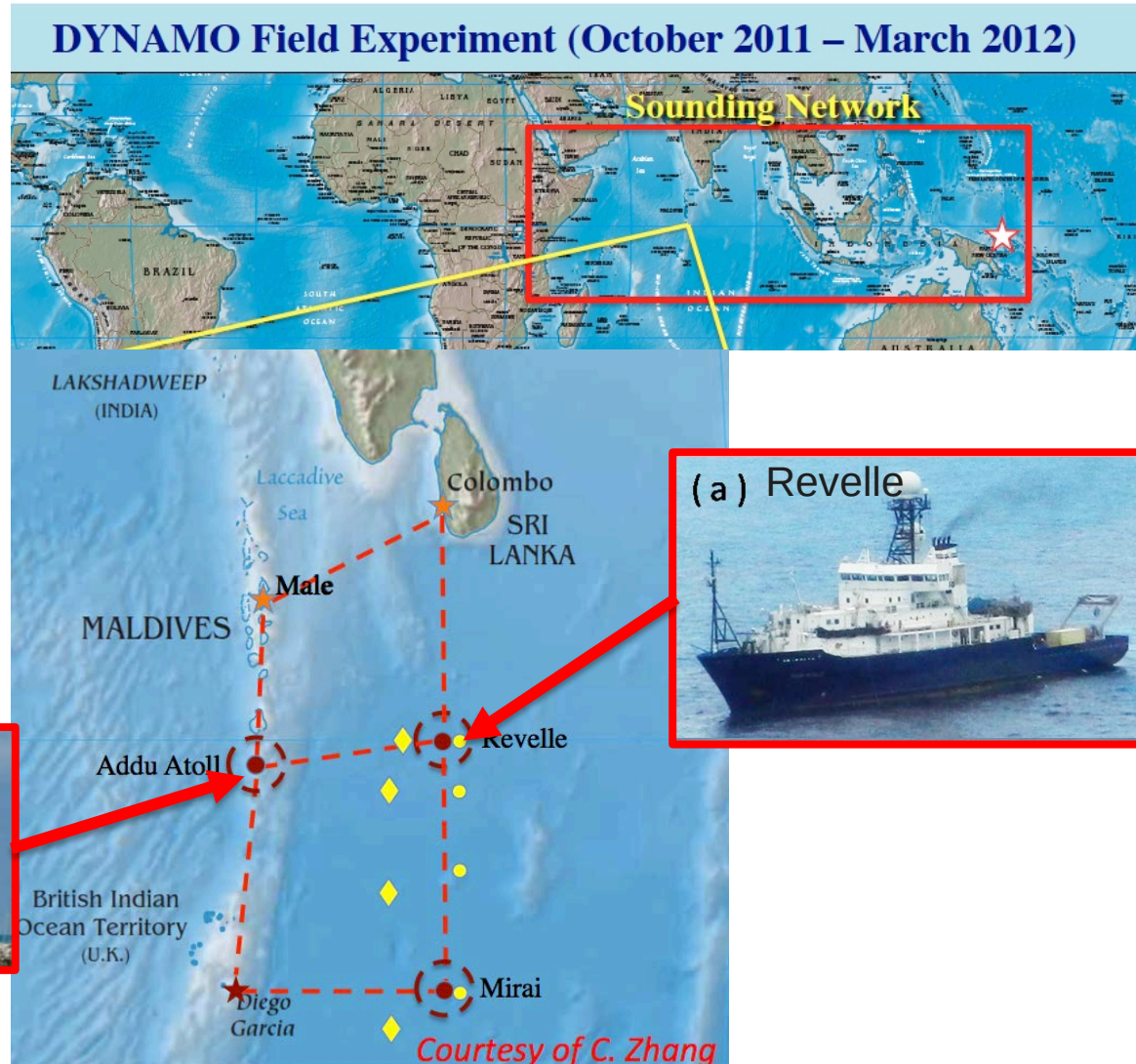
Boing et al. (2012)

Objectives

- ▶ **Use AMIE/DYNAMO observations to validate high-resolution WRF simulation in correctly representing cold pool statistics**
- ▶ **Parameterize effects of cold pool by using data from high-resolution simulation**

AMIE/DYNAMO Experiment

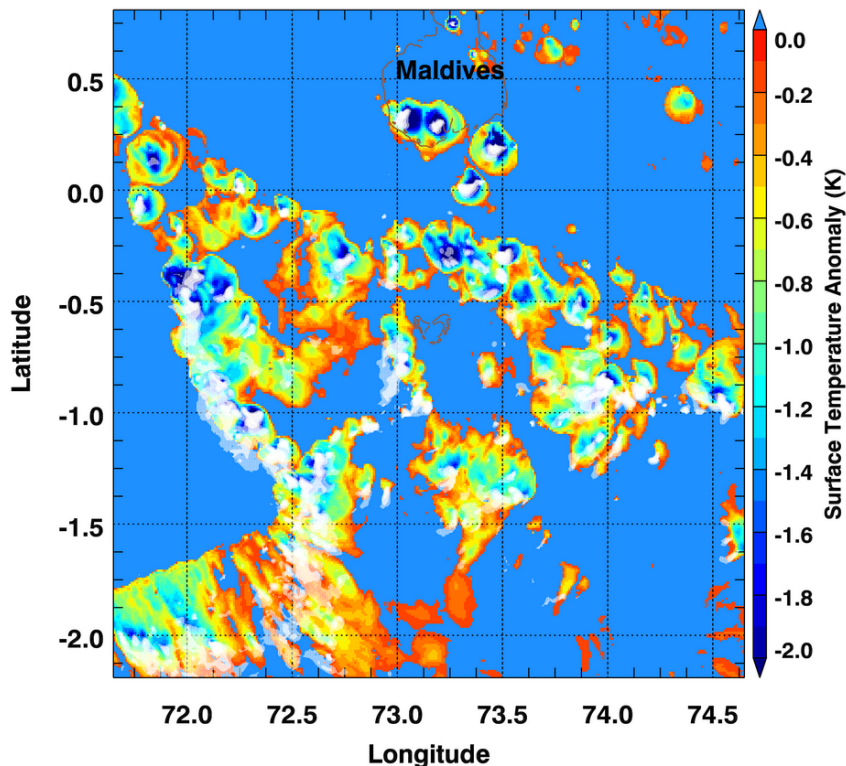
- ▶ Observations used:
- ▶ S-Pol radar convective cloud statistics and precipitation
- ▶ Revelle surface meteorology observations



High-res WRF Simulation

- Real-world simulation in large domain to obtain cold pool statistics

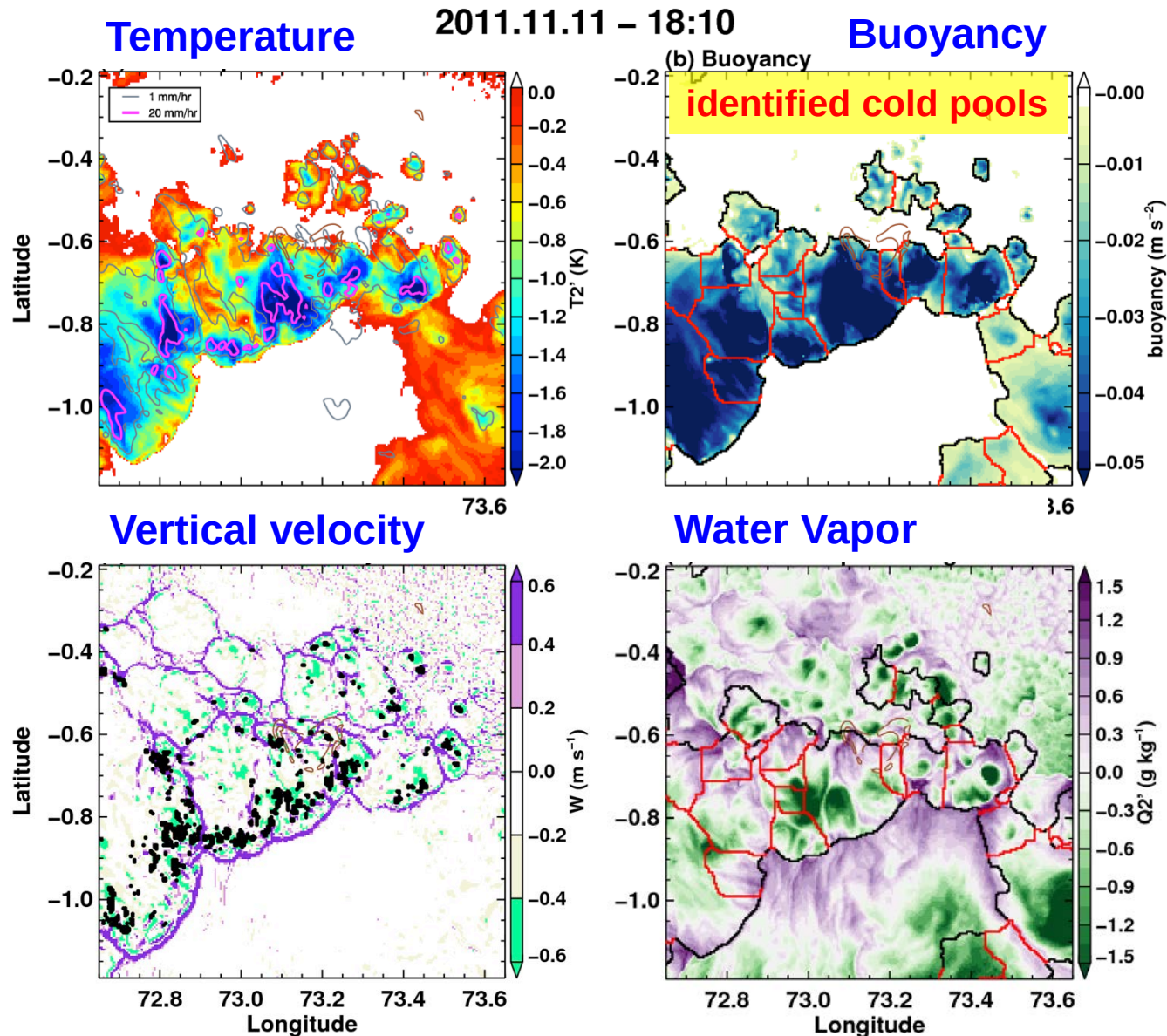
Simulated Surface Temperature and Cloud
2011.11.23 - 06:20



Model Setup	
Period	Nov 1 – 20, 2011
Domain	1000 x 500 km
Horizontal res.	500 m
Vertical res.	50 m (lowest 1 km) 40 levels
Microphysics	Thompson
Forcing	ERA-I 6-hourly analysis
PBL	UW
Surface	MO

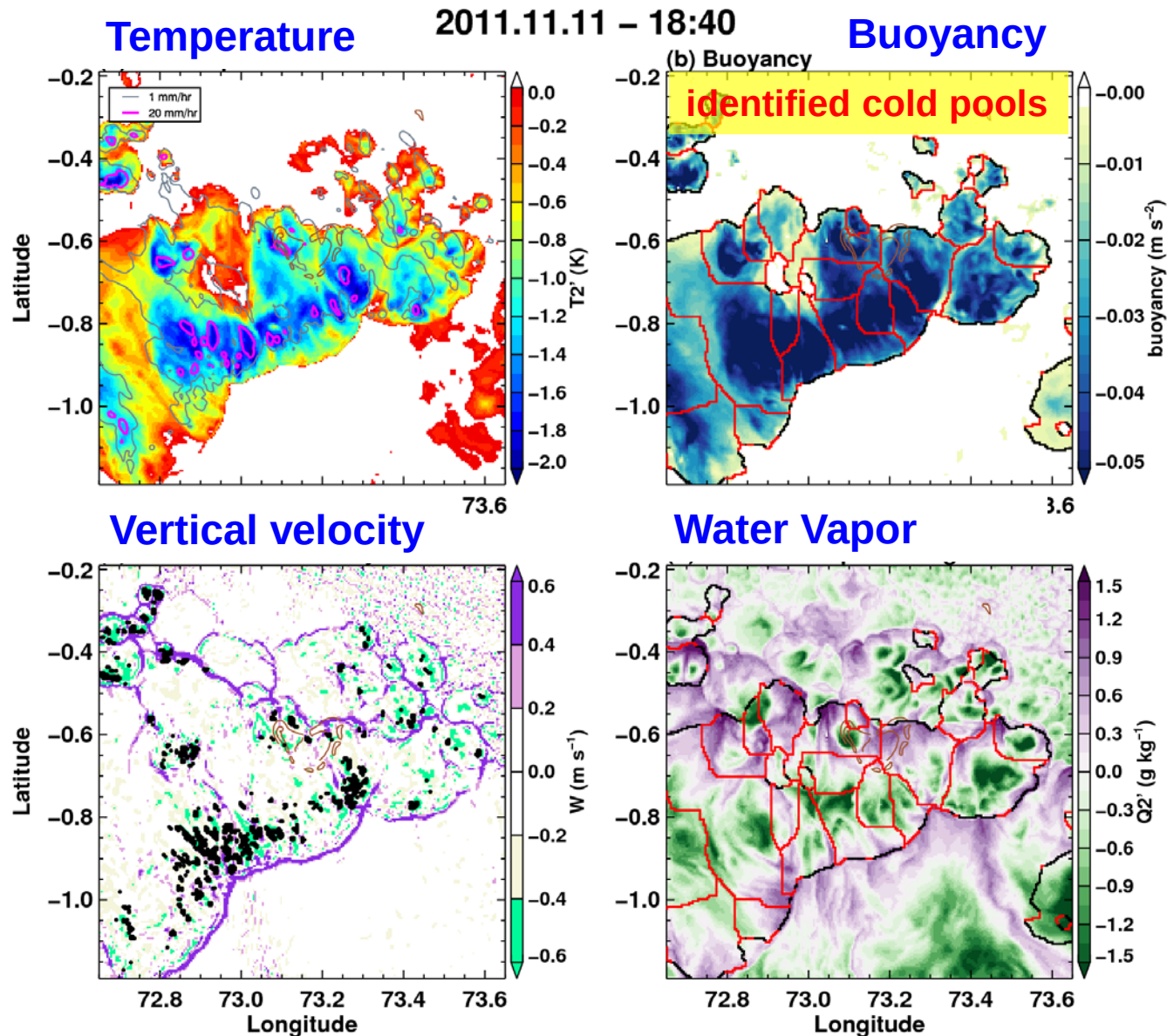
Examples of Simulated Cold Pool

- Identify and track individual cold pools using buoyancy
- 2m water vapor: moist anomaly at the edge, dry anomaly in the center
- Enhanced boundary layer vertical velocity at gust front

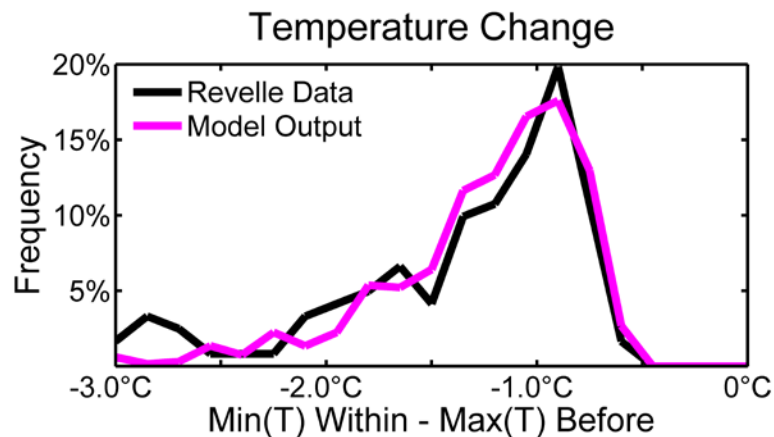


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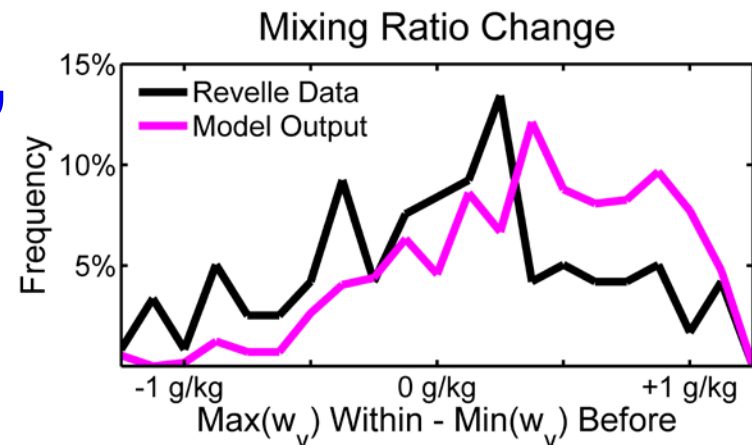
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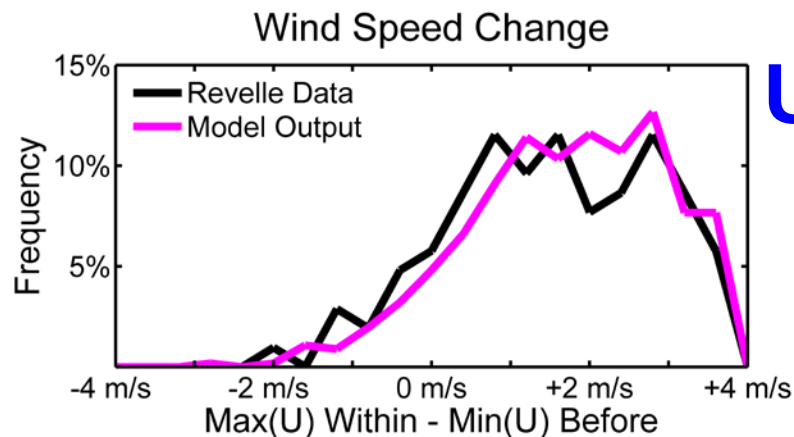
Comparison with Revelle Surface Meteorology



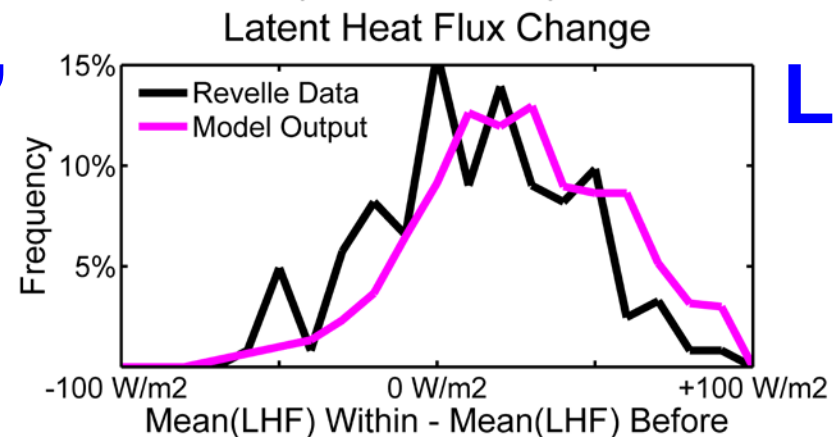
T'



Q_v'



U'



LH

- Identify cold pools using the same method
- Model produces comparable changes in temperature, wind speed and LH flux statistics with observations, but have a moist bias

Parameterize Cold Pool Effects on Wind Gustiness

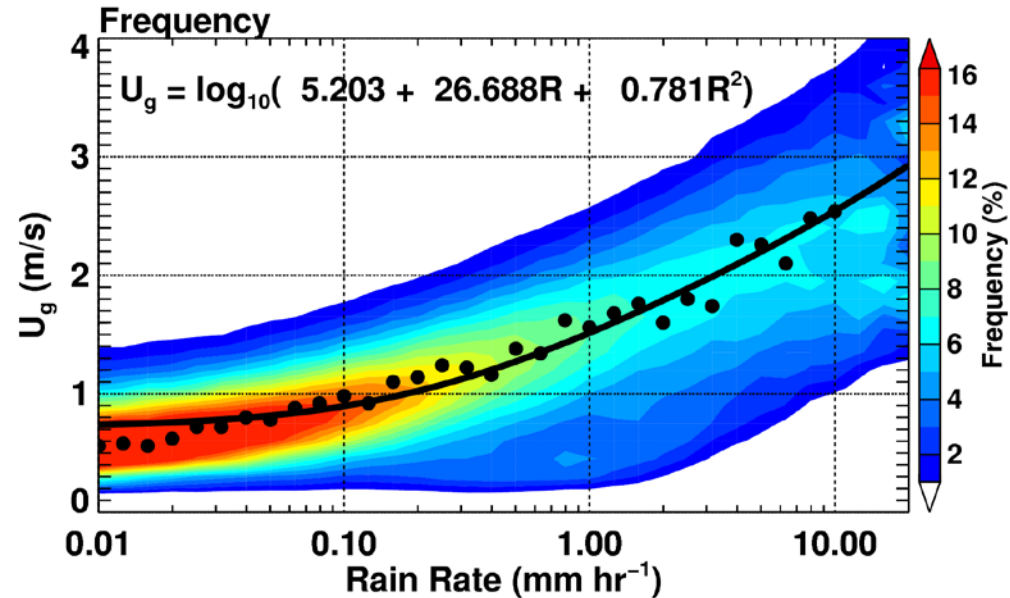
- ▶ Surface wind gustiness (Zeng et al. 2002) is a result of precipitation driven cold pool, it enhances surface fluxes
- ▶ Wind gustiness (U_g):

$$U_g = \left(U^2 - U_v^2 \right)^{1/2}$$

$$\left\{ \begin{array}{l} \text{vector wind : } U_v = \left(\bar{u}^2 + \bar{v}^2 \right)^{1/2} \\ \text{scalar wind : } U = \left(u^2 + v^2 \right)^{1/2} \end{array} \right.$$

- ▶ Parameterize U_g based on CRM simulation

Wind Gustiness vs. Precipitation (CRM)

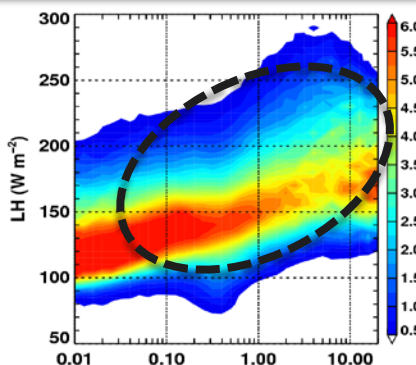
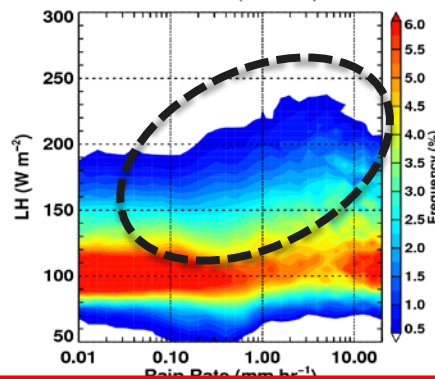
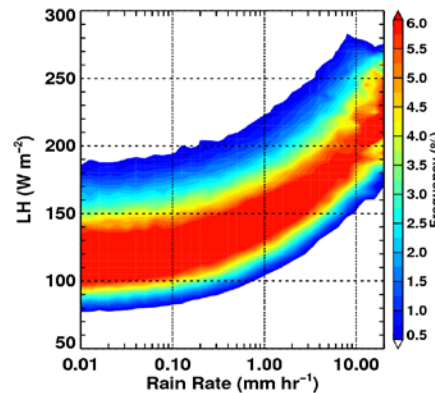


$$U_g = f(Rain)$$

Impacts of Parameterization on Surface Fluxes

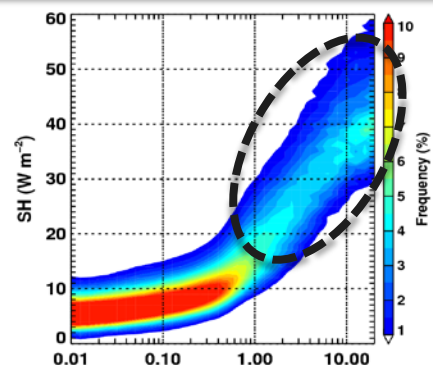
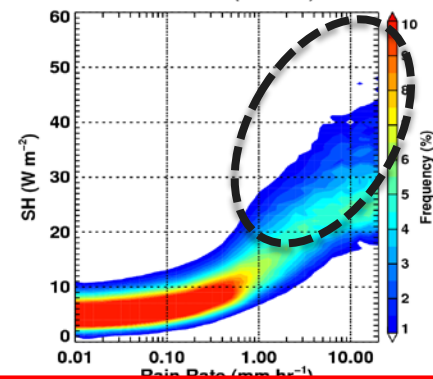
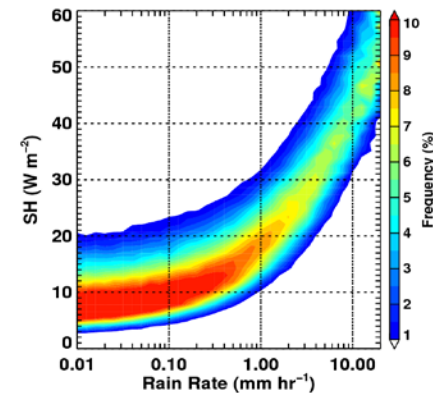
- ▶ CRM results show clear relationship of LH & SH flux with precipitation
- ▶ Control simulation: (20-km resolution)
 - LH has no sensitivity to precipitation
 - SH sensitivity is lower
- ▶ Simulation with parameterized U_g improves SH & LH flux and compares better with CRM

Latent Heat Flux



Rain Rate (mm/hr)

Sensible Heat Flux



Rain Rate (mm/hr)

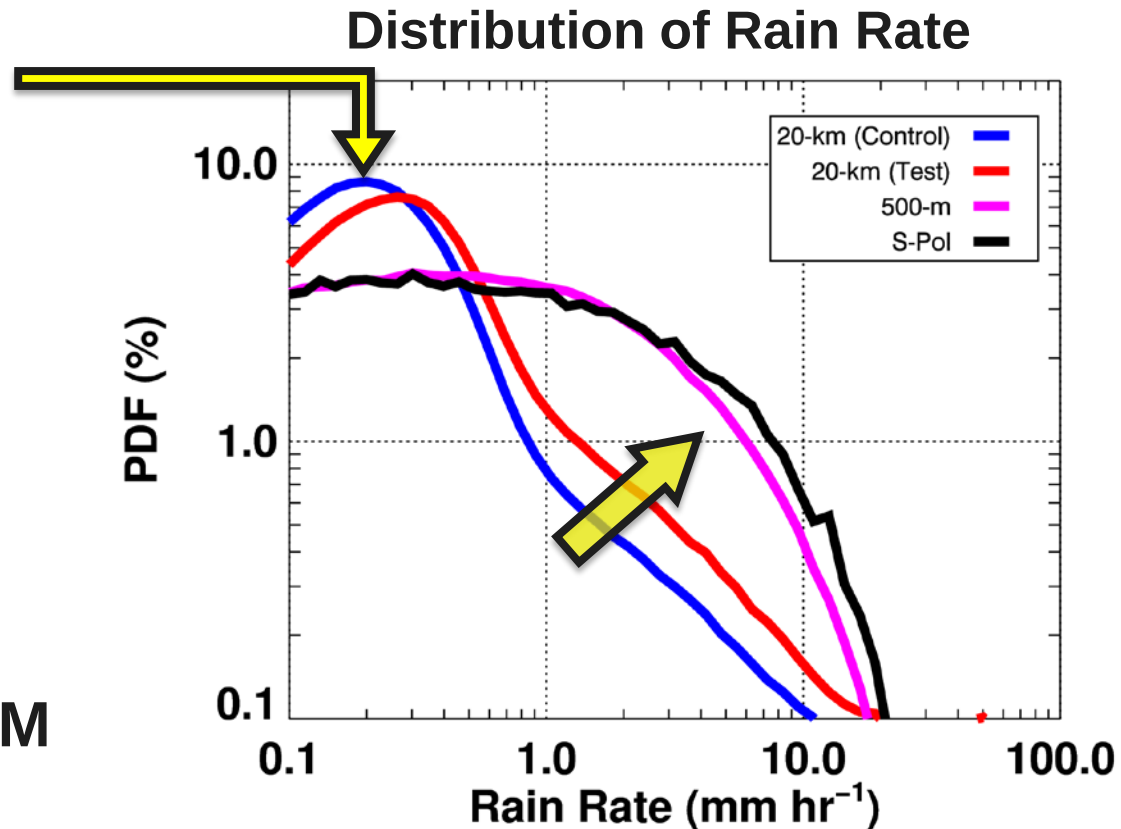
CRM

20-km
Control

20-km
with U_g

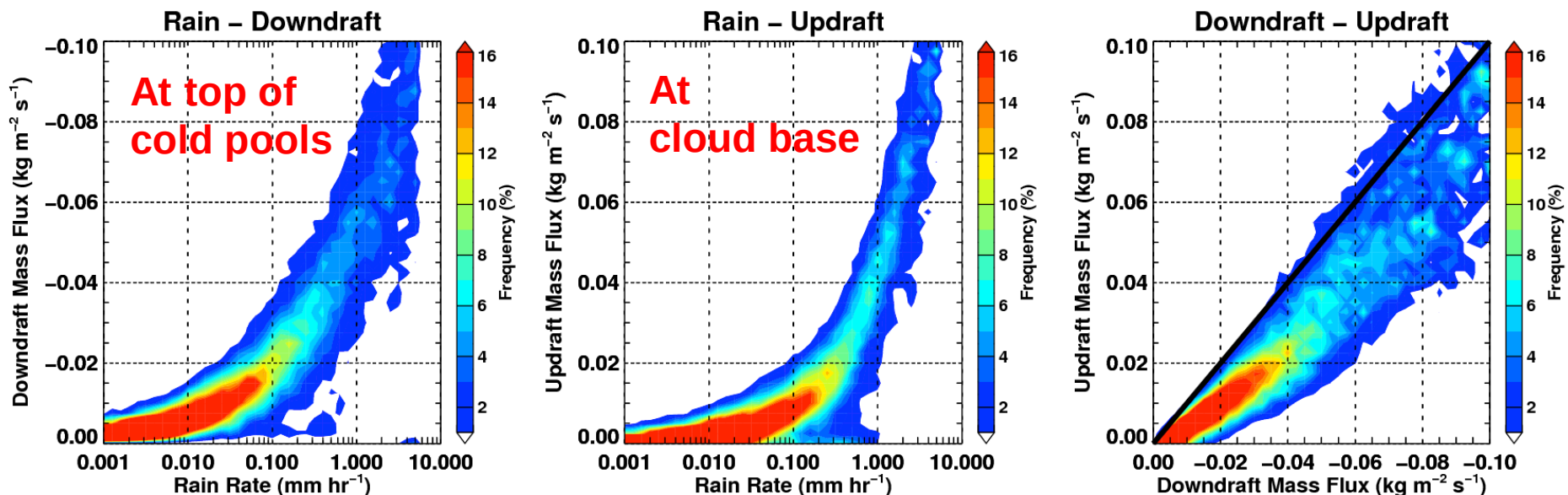
Feedback on Precipitation

- ▶ Too much drizzle is a common problem in GCMs (Dai 2006)
- ▶ Enhances surface fluxes increases light to moderate rain rate
- ▶ But rain rates are still underestimated compared to both CRM and S-Pol
- ▶ **Next-step: cold pool dynamic effect**



S-Pol: derived rain rate
Model: direct rain rate

Cold Pool Effects on Secondary Updraft



- ▶ Identify downdraft and **secondary updraft** associated with cold pools, average in 100 km box (GCM resolution)
- ▶ Strong relationship between precipitation, downdraft mass flux at cold pool top, secondary updraft mass flux
- ▶ Will work on parameterizing this dynamic effect

Summary

- ▶ High-resolution WRF simulation are performed over tropical Indian Ocean during AMIE/DYNAMO, model produced comparable observed precipitation and cold pool statistics
- ▶ Cold pool effects on enhancing surface fluxes through wind gustiness are parameterized
- ▶ Future work will also parameterize updraft enhancement
- ▶ Test parameterizations on climate models

