## Surface Energy Fluxes over the Eastern North

# Pacific Measured during MAGIC

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#### Acknowledgements

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Atmospheric System Research (ASR) Program



Atmospheric Radiation
Measurement (ARM) Program



#### Importance of Surface Fluxes

Surface fluxes provide important (and often dominant) contributions to energy and moisture in the MBL.

Surface fluxes are controlling factors for the dynamics of the MBL and for cloud formation.

Surface fluxes are boundary conditions for models.

## Surface Fluxes during MAGIC

1-minute time series of surface energy fluxes during MAGIC are available on the ARM archive

Two data files are available: "flux.mat" and "magic\_flux.txt" To access these data,

- 1) go to <a href="www.arm.gov/campaigns/amf2012magic">www.arm.gov/campaigns/amf2012magic</a>
- 2) go to "Bulk Aerodynamic Fluxes" under "Campaign Data Sets"
- 3) click "Order Data"

Be sure to read the file "magic\_flux\_readme\_arm.txt" which is also in the archive, and the document "OnDataProcessing" at www.rmrco.com/cruise/magic/data/OnDataProcessing/

#### Sea-Air Fluxes (watch sign conventions!)

$$H_{\text{net}}(\uparrow = +) = (H_{\text{sens}} + H_{\text{lat}}) - (R_{\text{sw}} + R_{\text{sw}}) - (R_{\text{lw}} + R_{\text{lw}}) + H_{\text{rain}}$$

H<sub>net</sub>: net heat flux

H<sub>sens</sub>: sensible heat flux

H<sub>lat</sub>: latent heat flux

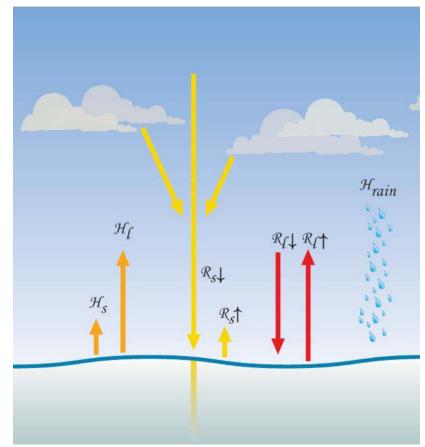
R<sub>sw</sub>↓: SW downwelling irradiance

R<sub>sw</sub>: SW upwelling irradiance

R<sub>lw</sub>: LW downwelling irradiance

R<sub>Iw</sub>: LW upwelling irradiance

H<sub>rain</sub>: heat from rain at T ≠ T<sub>surf</sub>



Adapted from Fairall and Bradley, 2006

## Bulk Aerodynamic Method

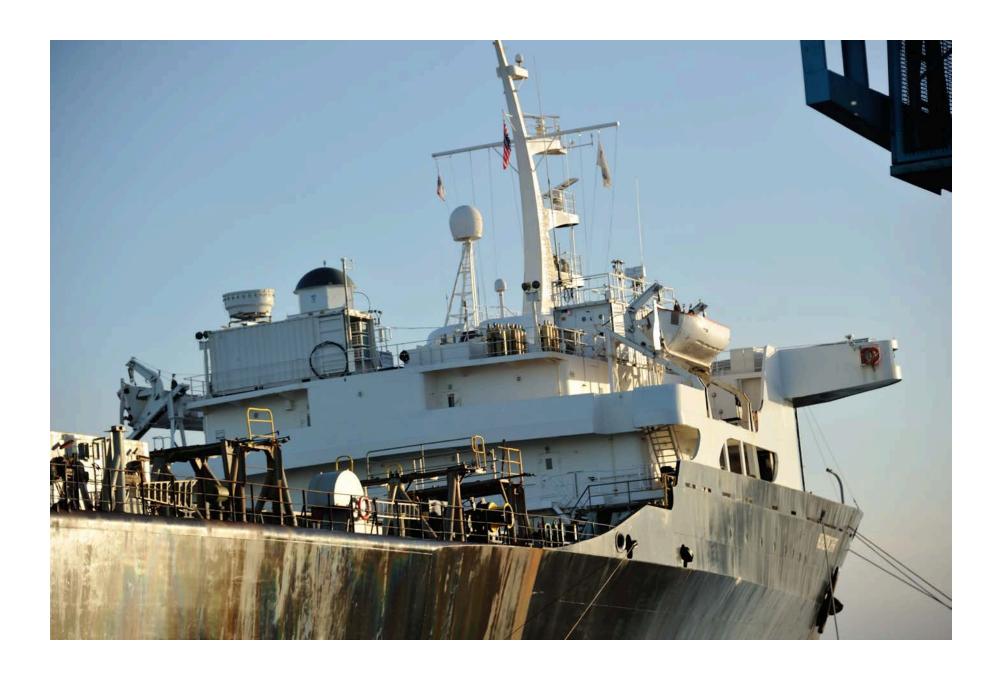
 $H_{sens}$  and  $H_{lat}$  are calculated using the TOGA-COARE algorithm, with the following measured quantities as inputs:

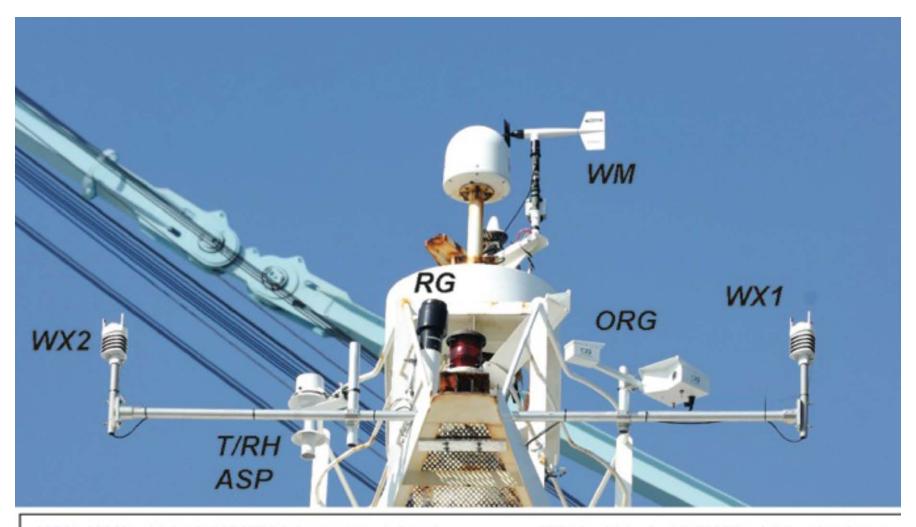
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U_{10} (wind speed)
     T<sub>air</sub> (air temperature)
     SSST (sea surface skin temperature)
     RH (relative humidity)
H_{\text{sens}} = \rho_{\text{air}} C_{\text{p,air}} C_{\text{H}} (U_{10} - U_{\text{sea}}) * (\Theta_{\text{surf}} - \Theta)
     \Theta \approx T + 0.01z (for z in meters)
H_{lat} = \rho_{air} L_w C_E (U_{10} - U_{sea}) * (q_{surf} - q)
     q_{surf} = 0.98 q_{sat}(T_{surf})
     q = (RH/100) q_{sat}(T)
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 $C_{\rm H} \approx C_{\rm E} \approx 0.001$ , with little  $U_{10}$  dependence (for neutral stability)

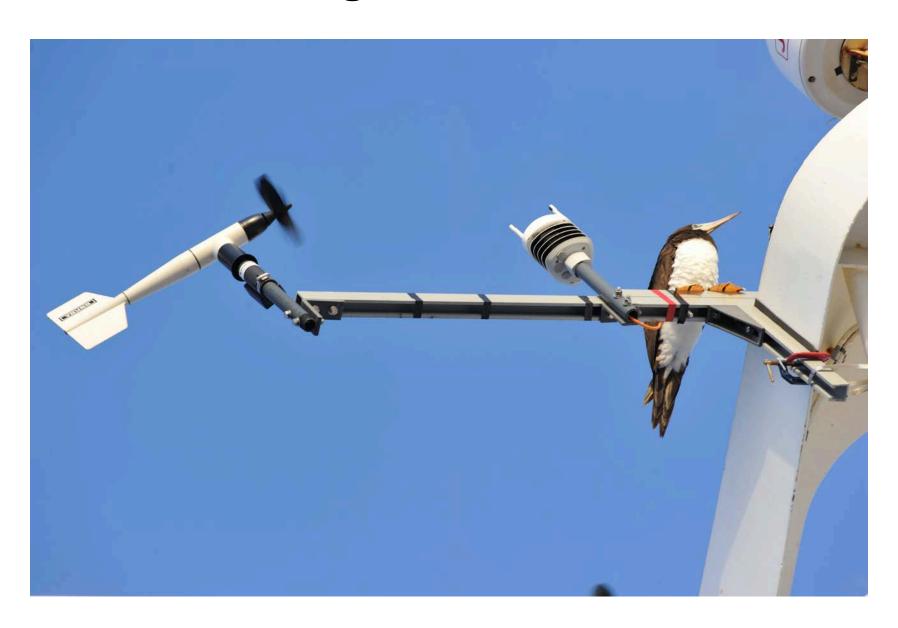


The meteorological mast is ~27 m above sea level.



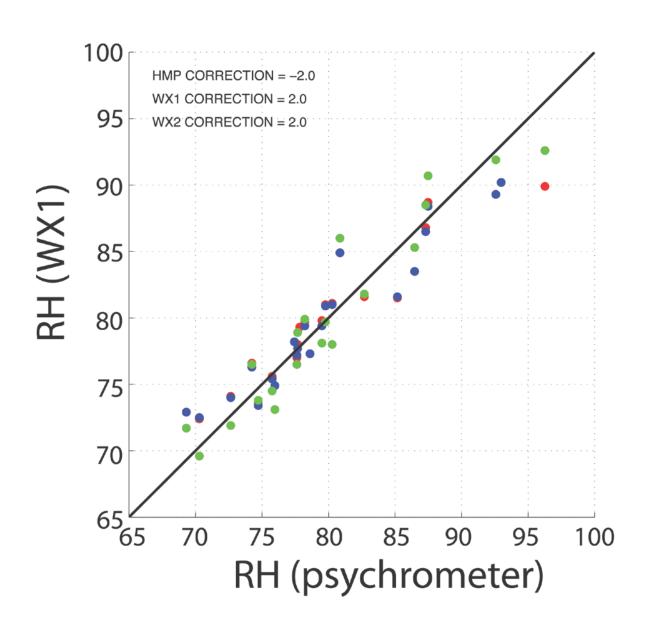


WX1, WX2 = Vaisala WXT520 (ws,wd,ta,rh,bp,ri) RG = Young 50203 Siphon Rain Gauge WM = Young 05106 Wind Monitor T/RH = Vaisala HMP155, ASP = Aspirator ORG = OSI Optical Rain Gauge





#### RH Calibration with Psychrometer



#### Sea Surface Temperature



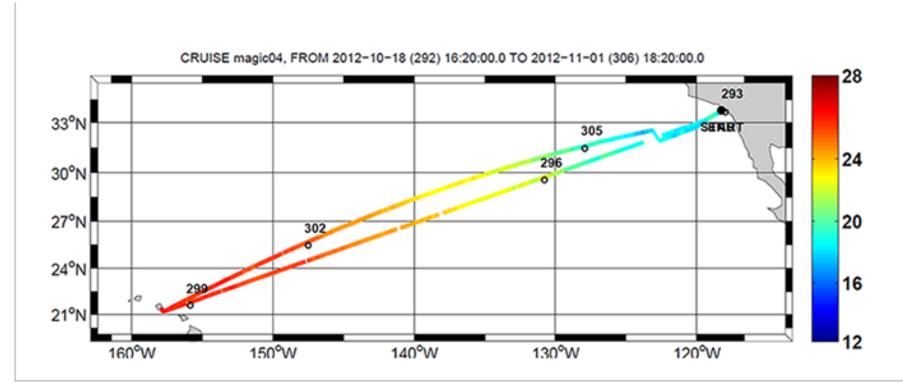


IRT ISAR

IRT (Wintronics): Heitronics KT19.85 Infrared Radiation Pyrometer ISAR - Infrared Sea Surface Temperature Autonomous Radiometer

Both detect in the 9.6-11.5  $\mu$ m window.

#### Sea Surface Temperature



Leg04 (2012-10-18 to 2012-11-01)

#### Radiometric Fluxes

 $R_{sw}\downarrow$ : measured with Precision Spectral Pyranometer (PSP)

 $R_{sw}\uparrow$ : calculated from  $R_{sw}\downarrow$ , sea surface albedo, solar angle

R<sub>Iw</sub>: measured with Precision Infrared Radiometer (PIR)

RIWT: calculated from measured SSST, sea surface emissivity

#### Radiation Measurements

PSP SPN PIR

Portable Radiation Package (PRP)



PRP with FRSR

Precision Spectral Pyranometer (PSP) – downwelling broadband, 0.285–2.8  $\mu$ m Precision Infrared Radiometer (PIR) – downwelling broadband, 4–50  $\mu$ m Sunshine Pyranometer (SPN1) – total and diffuse irradiance Fast Rotating Shadowband Radiometer (FRSR)

- 10 nm wide channels at 415, 500, 680, 870, and 940 nm
- direct normal irradiance, diffuse irradiance, total irradiance; also AOT

#### Hrain calculated from measured rainfall rates

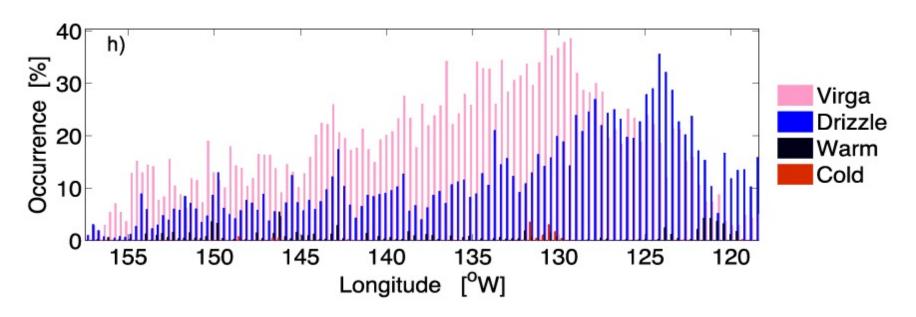
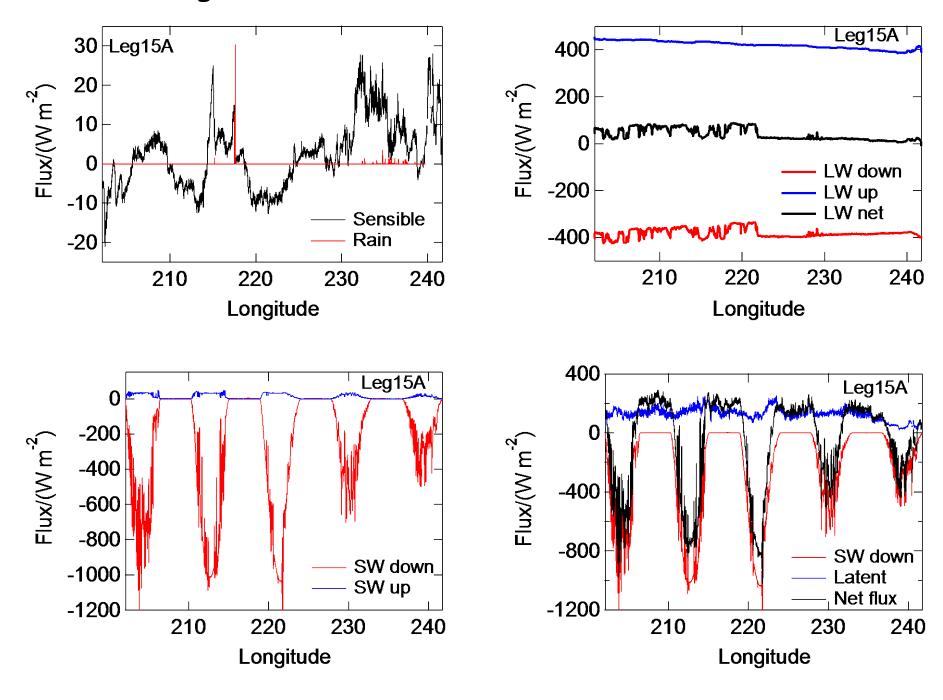
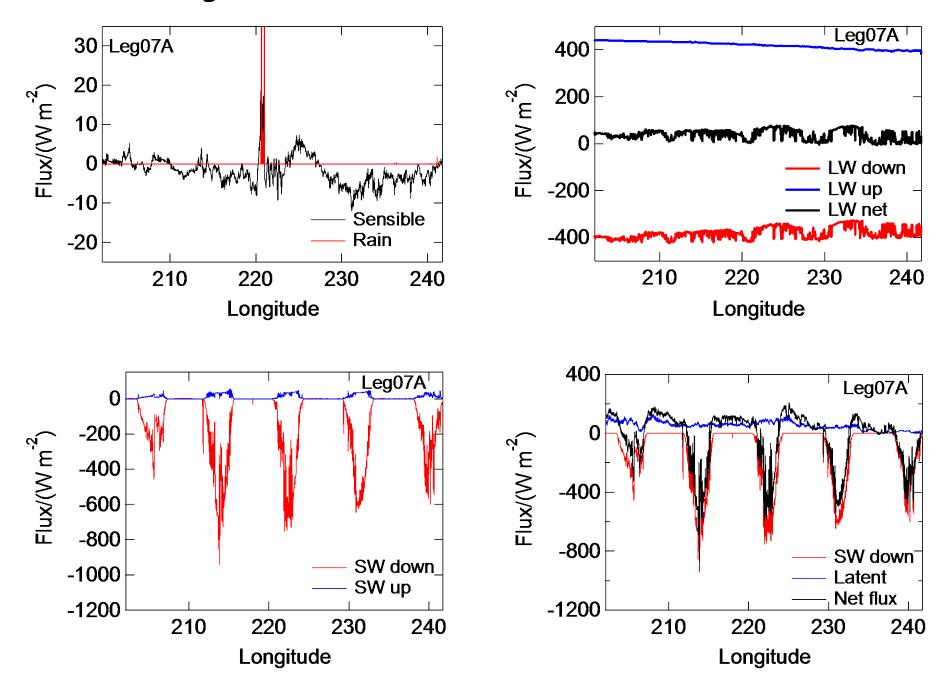


Figure from Xiaoli Zhou, McGill University

#### Leg15A (2013-07-20 to 2013-07-25) Fluxes



#### Leg07A (2012-12-01 to 2012-12-06) Fluxes



# Sea-Air Fluxes (up = +)

	Overall	Leg15A	Leg07A
	$(n \approx 200,000)$	(n ≈ 6,800)	(n ≈ 6,900)
Rain	1.5 ± 14	$0.03 \pm 0.5$	0.04 ± 1
Sensible	7 ± 10	2 ± 8	-2 ± 3
Latent	114 ± 58	128 ± 42	52 ± 28
LW <sub>up</sub>	420 ± 18	421 ± 17	419 ± 15
LW <sub>down</sub>	-377 ± 26	-382 ± 18	-381 ± 23
LWnet	46 ± 25	40 ± 25	38 ± 22
SW <sub>up</sub>	11 ± 14	12 ± 14	9 ± 13
SWdown	-210 ± 302	-241 ± 317	-123 ± 189
SWnet	-200 ± 291	-229 ± 305	-114 ± 177
Net	-31 ± 292	-61 ± 305	-25 ± 178

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LWup	420 ± 18	421 ± 17	419 ± 15
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LWnet	46 ± 25	40 ± 25	38 ± 22
5W <sub>up</sub>	11 ± 14	12 ± 14	9 ± 13
SWdown	-210 ± 302	-241 ± 317	-123 ± 189
SWnet	-200 ± 291	-229 ± 305	-114 ± 177
Net	-31 ± 292	-61 ± 305	-25 ± 178

#### Issues

Possible sampling bias (day/night sampling might not be equal)

Moving target - days start and stop at different locations and length of day changes as ship moves

#### Conclusions

H<sub>net</sub> is generally downward, but with large variability

H<sub>sens</sub> provides only a very minor component

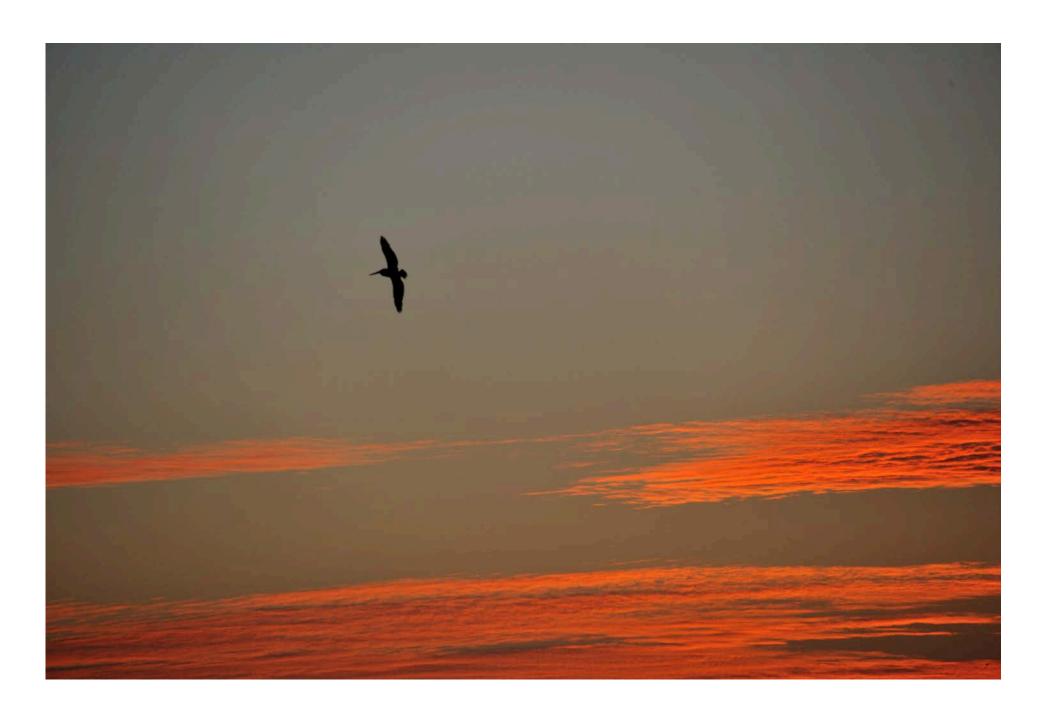
H<sub>lat</sub> is a major (upward) component, summer larger than winter

 $R_{sw}\downarrow$  is large and downward, but with large diurnal variability and distinct longitudinal patterns (due to clouds)

R<sub>sw</sub> provides only a minor contribution

 $R_{lw}$  is largest downward component, but nearly cancelled by  $R_{lw}$  such that  $R_{lw,net}$  is upward but small

Hrain is quite variable but very small



Thank you!