

**Simulated aerosol radiative forcing  
by overlying and entrained smoke aerosol:  
Importance of background conditions and longwave impacts**

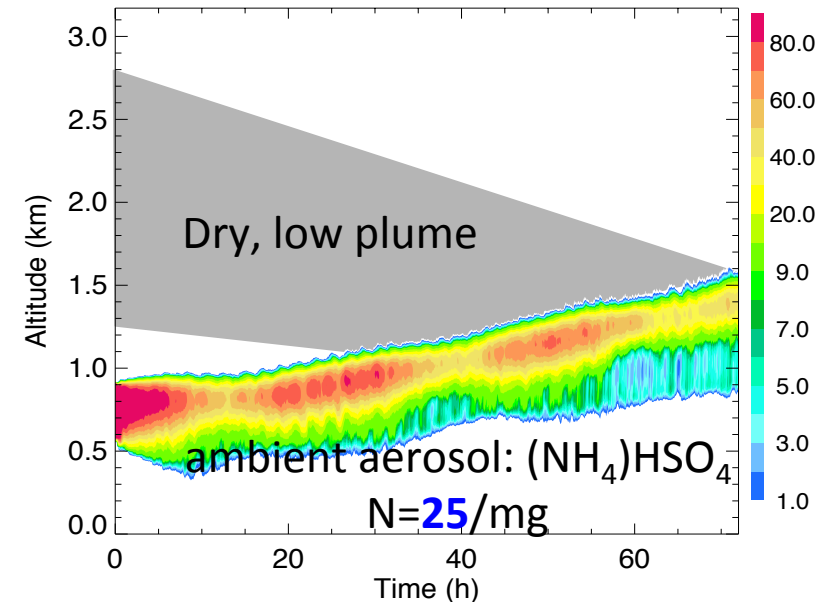
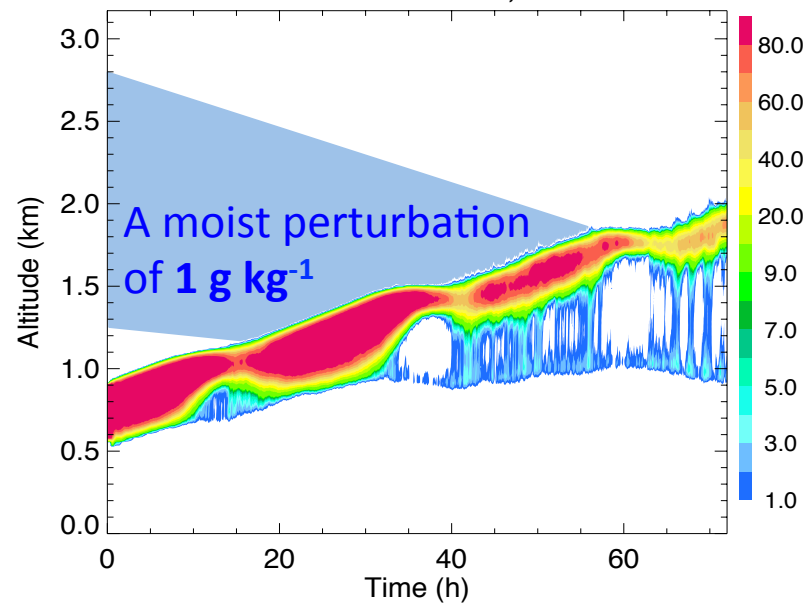
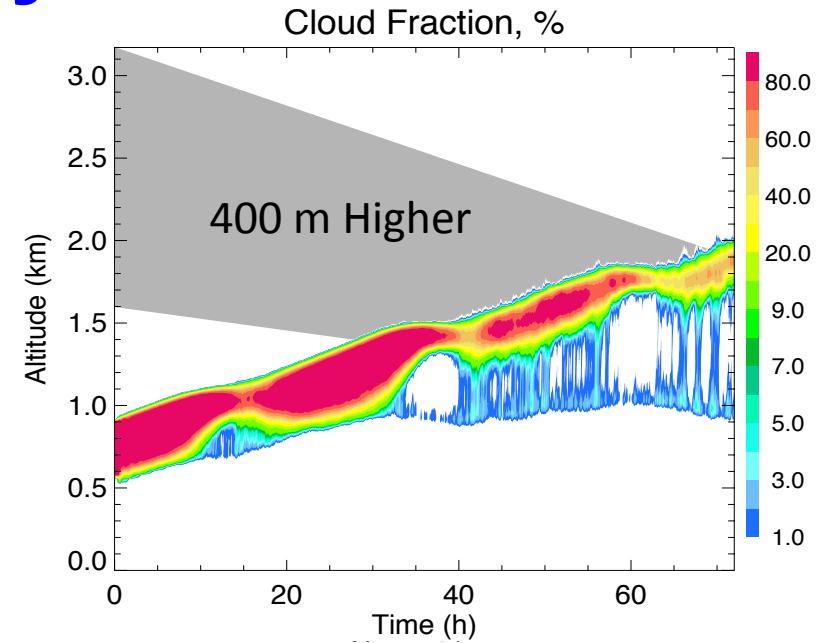
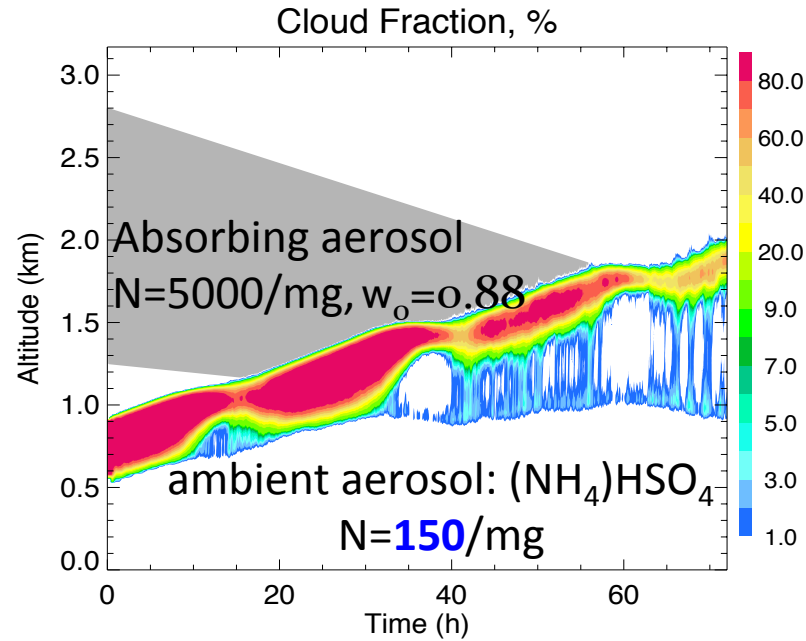
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# Background

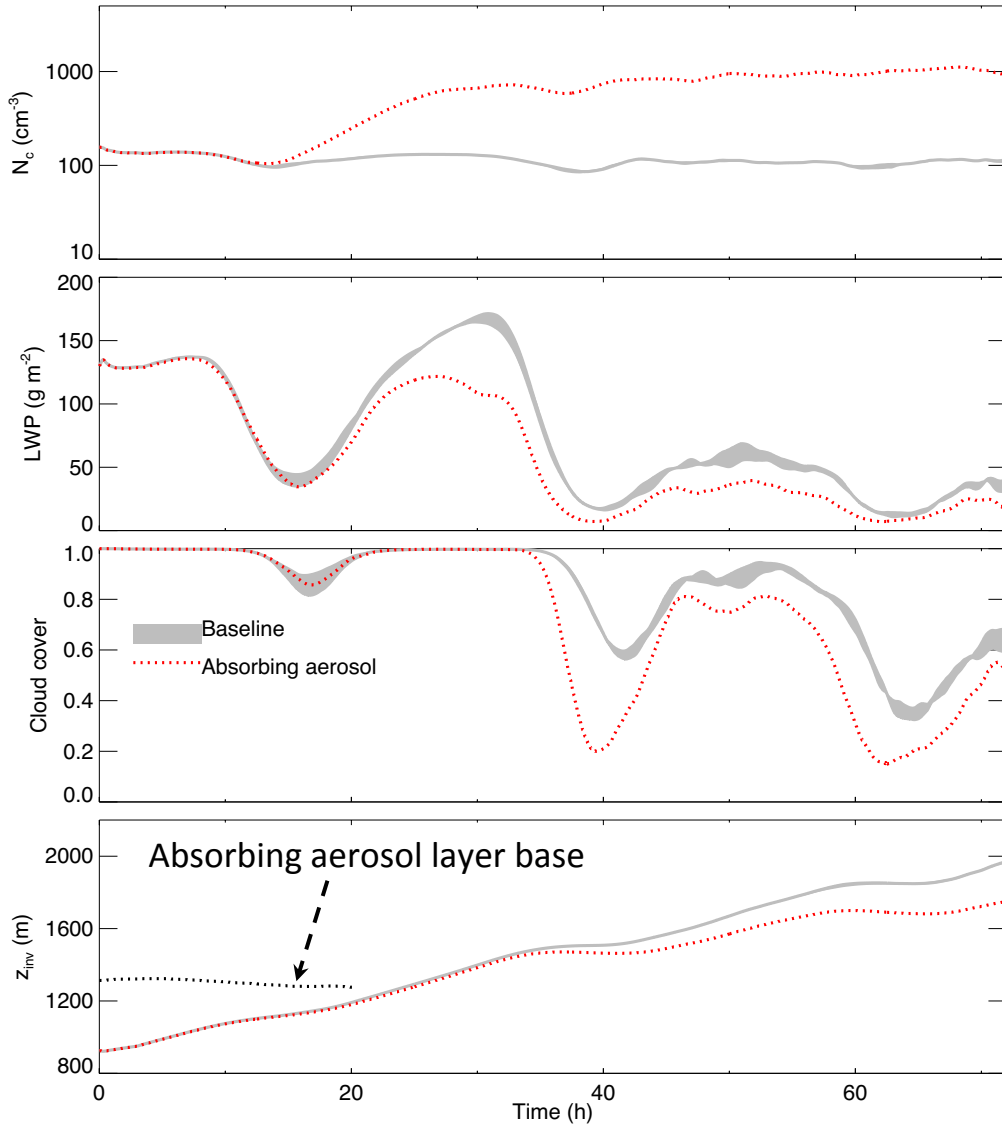
- Semi-direct effect (Hansen et al. 1997): aerosol solar absorption evaporates embedded clouds, **+ive shortwave (SW) radiative forcing at TOA (DF)**
- Johnson et al. (2003): heating *immediately above* Sc strengthens inversion, reduces entrainment drying, thickens clouds, **-ive SW DF**
- Here: consider effect of overlying absorbing aerosol (active as CCN) on transition of Sc → TrCu, **+ive SW RF? -ive SW DF?**
- Unperturbed: advection across warmer SSTs drives progressive entrainment, cloud-top LW cooling cannot maintain well-mixed PBL → buildup of heat and moisture in surface layer drive shallow Cu (Wyant et al. 1997)
- Yamaguchi et al. (2015): absorbing aerosol *delays* transition via slower entrainment, less drizzle, **-ive change in SW cloud radiative forcing,  $\downarrow CRF = \downarrow F - \downarrow F(\text{clear sky})$**
- **What about longwave (LW)  $\downarrow F$ ?**

# Simulation strategy

- 10.8 x 10.8 x 3.2 km grid
- 2-moment microphysics
- meteorology: base case used in de Roode et al. (2016)  
LES intercomparison  
(from NE Pacific trajectories)
- ambient aerosol:  $k=0.55$ ,  
 $r_g=0.05 \mu\text{m}$ ,  $\sigma_g=1.2$
- absorbing layer:  $k=0.2$ ,  
 $r_g=0.12 \mu\text{m}$ ,  $\sigma_g=1.3$



# Indirect plus semi-direct forcings



- Transition hastened and strengthened (Y15: transition delayed instead)
- Net **negative** radiative forcing  $DF$  (but positive  $\Delta\text{CRF}$ , opposite of Y15)
  - SW: Twomey overcomes reduction in CF and LWP
  - LW: shallower PBL, reduced CF both increase upwelling LW

	$DF$ ( $\text{W m}^{-2}$ )		
	SW	LW	TOTAL
Day 2	-0.5	-2.6	-3.1
Day 3	-1.2	-6.0	-7.2

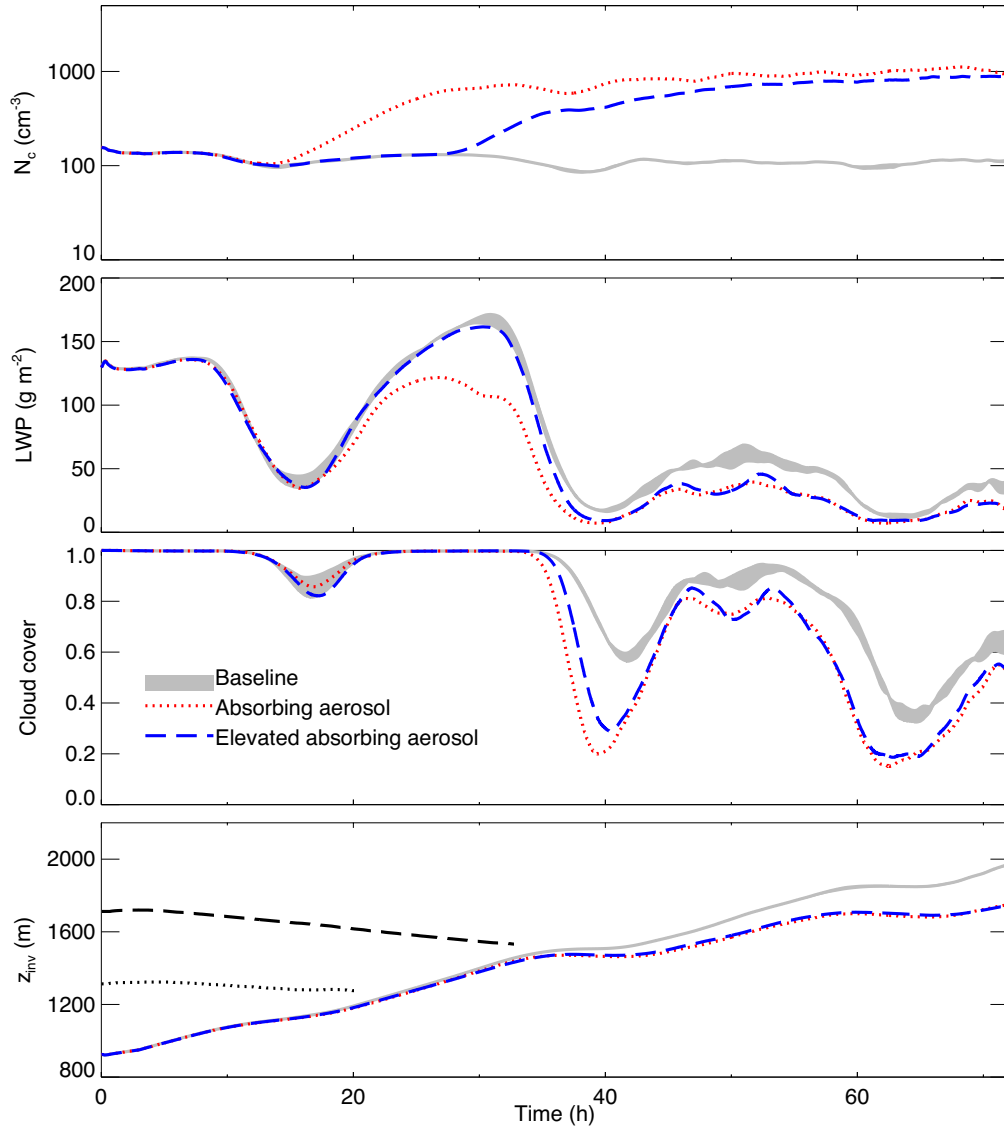
## Indirect forcings

		SW	LW
<b>Cloud layer</b>			
Twomey effect	$N_c \uparrow$	—	
Sedimentation $\downarrow$	LWP $\downarrow$	+	
	CF $\downarrow$	+	—

## Semi-direct forcings

		SW	LW
<b>FT heating</b>			
Inversion strength $\uparrow$	CF $\uparrow$	—	+
	$Z_i \downarrow$		—
<b>PBL heating</b>			
RH $\downarrow$ stabilization $\uparrow$	CF $\downarrow$	+	—
	$Z_i \downarrow$		—

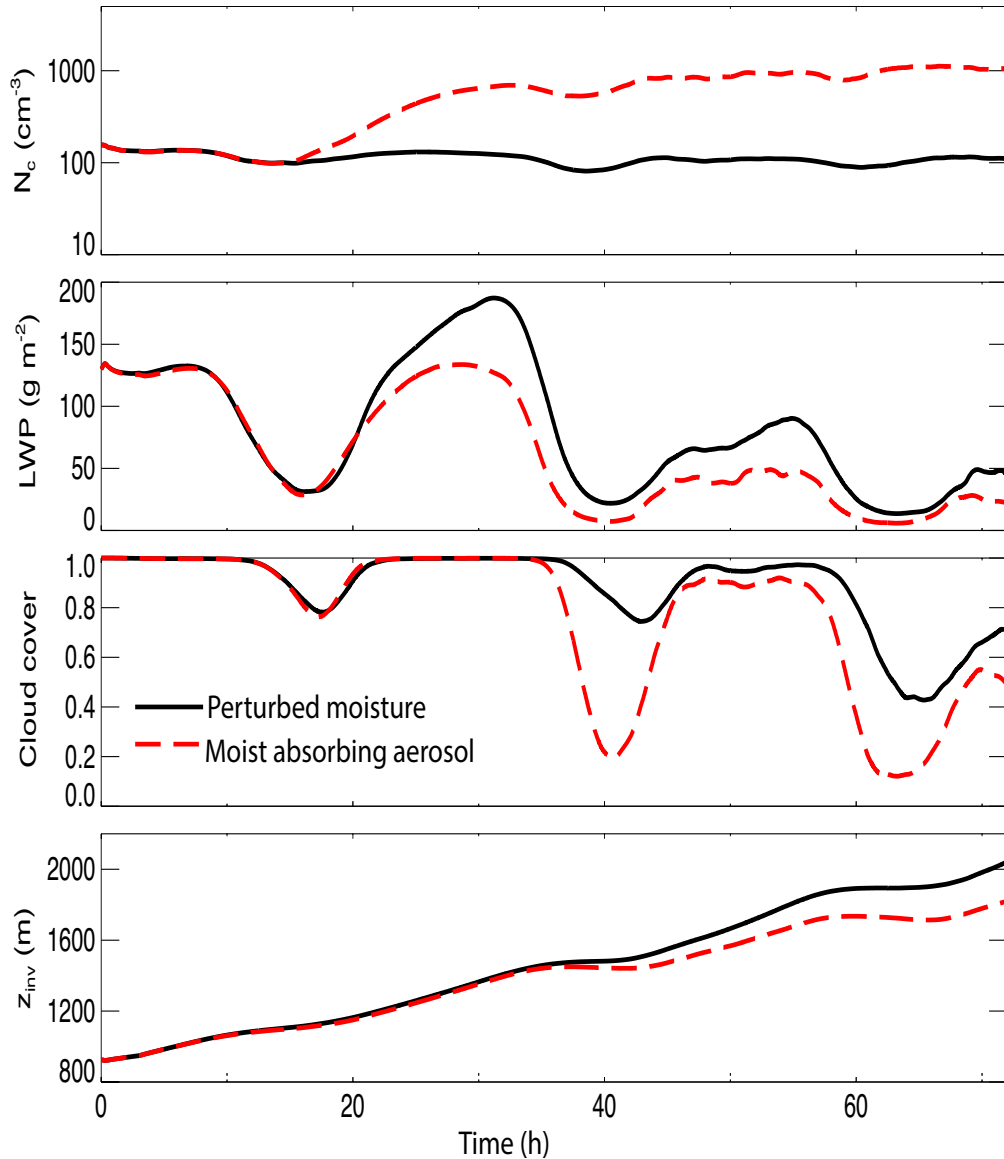
# Higher initial absorbing aerosol layer



- Still, transition hastened
- **Stronger net negative forcing**
  - SW: Twomey more dominant (greater CF)
  - LW: forcing reduced (greater CF), yet not to be ignored

	DF ( $\text{W m}^{-2}$ )		
	SW	LW	TOTAL
Day 2	<b>-11.2</b>	<b>-1.9</b>	<b>-13.1</b>
Day 3	<b>-5.0</b>	<b>-4.7</b>	<b>-9.7</b>

# A moisture perturbation aloft

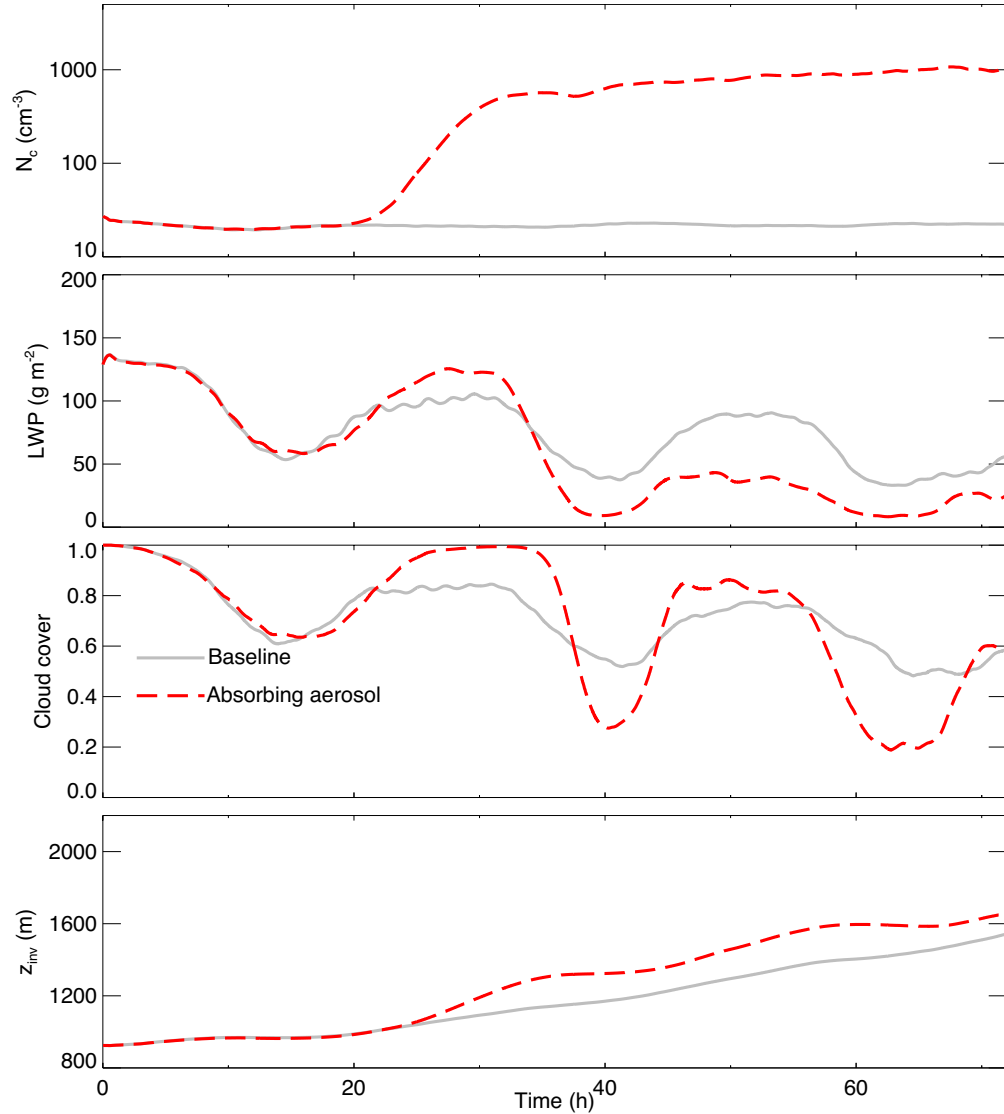


- Adebisi et al. (2015) report plume moister than environment by  $\sim 1 \text{ g/kg}$
- Relative to dry baseline: less entrainment drying  $\rightarrow$  greater LWP and CF
- A thicker cloud layer has more to lose  $\rightarrow$  greater reduction in LWP and CF
- SW: positive forcing ( $-\Delta\text{CF}$  overcomes Twomey)
- LW: forcing slightly increased (greater  $-\Delta\text{CF}$ )

$\rightarrow$  **weaker net negative forcing**

	$DF$ ( $\text{W m}^{-2}$ )		
	SW	LW	TOTAL
Day 2	<b>3.0</b>	<b>-2.2</b>	<b>0.8</b>
Day 3	<b>2.8</b>	<b>-6.8</b>	<b>-4.0</b>

# Heavily drizzling baseline, dry aerosol plume



- Background sulfate: 150  $\rightarrow$  25/mg
- Baseline: drizzle reduces diurnal cycle of LWP, CF and reduces entrainment
- Entrainment of aerosol inhibits drizzle, reversing those changes
- SW: stronger Twomey effect (much lower baseline)
- LW forcing *reversed* (deeper PBL)
- Transition still not delayed

	$DF$ ( $\text{W m}^{-2}$ )		
	SW	LW	TOTAL
Day 2	<b>-52.0</b>	<b>6.3</b>	<b>-45.7</b>
Day 3	<b>-9.4</b>	<b>3.4</b>	<b>-6.0</b>



# Summary

- TOA indirect and semi-direct radiative forcings consistently **-ive** (greater than or comparable to +ive direct forcing)
  - LW forcings not to be ignored (subtropical troposphere is dry): substantial and **-ive**, induced by reduction in CF and PBL height, reversed in drizzling case
  - Higher aerosol layer **enhances**
  - Presence of additional moisture **reduces**
  - Presence of drizzle **enhances**
- } total aerosol radiative forcing
- Unlike Y15, here Sc → TrCu transition never delayed: *why?*

## Differences with Y15

- **Possible differences in microphysics treatment:**

Omit cloud droplet sedimentation and fix relaxation time for diffusional growth in DHARMA

→ Transition *delayed*

- **Differences in set-up:**

Y15 used prognostic aerosol

→ Feedback of aerosol consumption and drizzle leads to catastrophic cloud breakup at end of control run

- **Differences in the LES dynamics:**

Sc breakup much greater for DHARMA than SAM (used in Y15) in intercomparison of same SCT case (de Roode et al. 2016)

