Aerosol-Cloud Droplet Number Closure from ISDAC Observations

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Overview

- Energy balance of Arctic region particularly sensitive to enhancements of natural and anthropogenic aerosol
- Associated aerosol indirect effects modulate formation and properties of clouds significant impact on climate predictions
- Conducting aerosol-cloud droplet number closure study using aircraft/ground measurements in Arctic from DOE Indirect and Semi-Direct Aerosol Campaign (ISDAC)
 - Conducted in vicinity of Barrow, Alaska, in April, 2008
- Emphasis on improving representations of vertical (updraft) velocity for parameterizations of aerosol activation into cloud droplets and evaluating existing parameterizations (Abdul-Razzak and Ghan)



- Simulate cloud droplet activation using aircraft measurements in an adiabatic parcel model
- Compare measured and simulated cloud droplet number concentrations to assess representations of cloud droplet activation

Aircraft Measurements

Good agreement in Aerosol size distribution with NOAA-P3 and NSA ground measurements
Sub-linear relationship between PCASP and cloud droplet number concentrations



independent measurements from NOAA P-3 aircraft and ARM North Slope of Alaska (NSA) ground site during a missed approach on April 20, 2008.



Cloud droplet number concentrations versus below-cloud aerosol (PCASP) concentrations for 7 identified cases from 6 flights during ISDAC.

Case Study: ISDAC Flight 31 (April 26, 2008)





Selection of below- and in-cloud cases for parcel model analysis from time series of altitude, cloud droplet concentration (N₂), and aerosol particle concentration (N₂), during ISDAC flight 31 on April 26, 2008.



Vertical velocity histogram for cloud interval 1. A Gaussian fit yields a standard deviation, σ , of 36 cm s⁻¹. Similar analysis for cloud interval 2 (not shown) yields $\sigma = 41$ cm s⁻¹.

Preliminary Results

- Parcel model analysis of this case indicates that updraft velocities of > 0.6σ give the same droplet number concentration
- For aerosol distributions such as these that are dominated by larger particles, the droplet number concentration is relatively more sensitive to the cloudbase aerosol concentrations than to the updraft speed



Aerosol particle composition determined using SPLAT for below-cloud intervals B1 and B2.

Case	W [m s ⁻¹]	N _{calc} [cm ⁻³]	N _{avg,obs} [cm ⁻³]	% Diff. (avg)	N _{max,obs} [cm ⁻³]	% Diff. (max)
1	>0.6σ	185	191	3	228	19
2	>0.6σ	160	190	16	215	26

Preliminary parcel model results for selected cases (above) during ISDAC flight 31. All simulations assume a mass accommodation coefficient, a, equal to 1. Calculated droplet concentrations are compared against average observed values N_{argubes} and 95th percentile observed values, N_{maxobs}.

Below-Cloud Aerosol Size Distributions for "Golden" Days



Date	Flights	[µm]	Ug	[cm-3]	
April 8	16	0.188	1.40	165	
April 26	30,31	0.194	1.48	199	

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