

# Large-eddy simulations of airflow dynamics and physics over the island of Graciosa

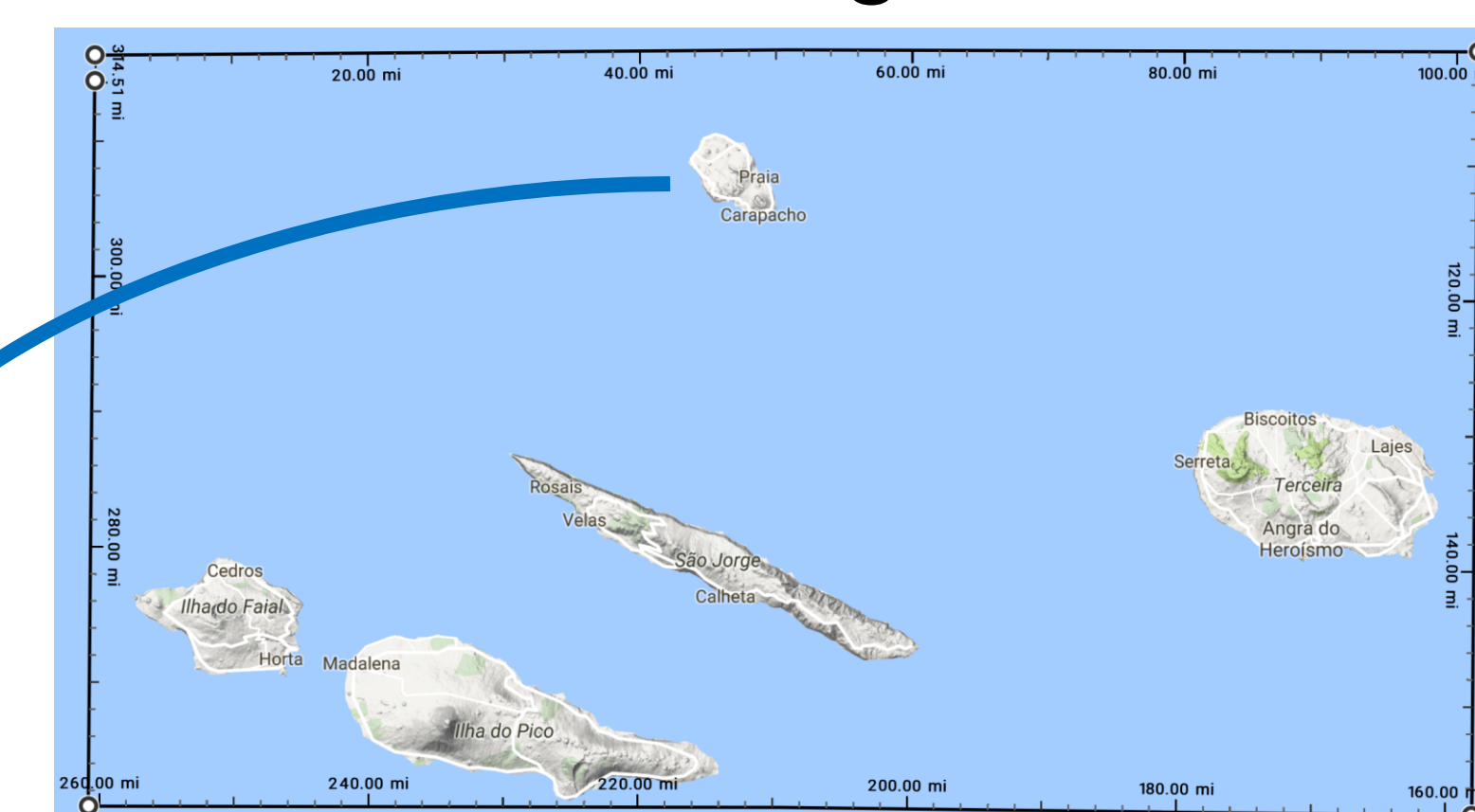
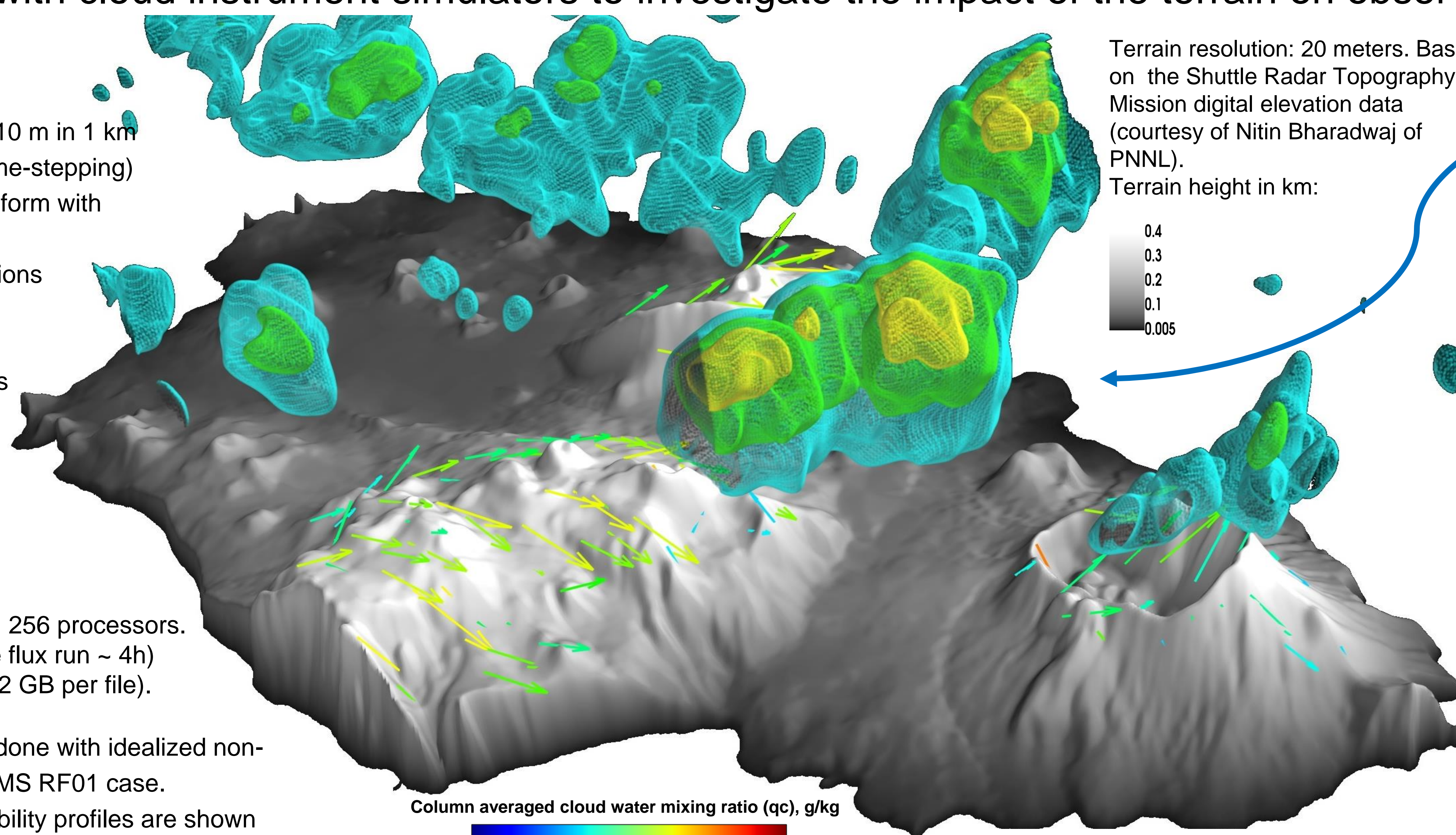
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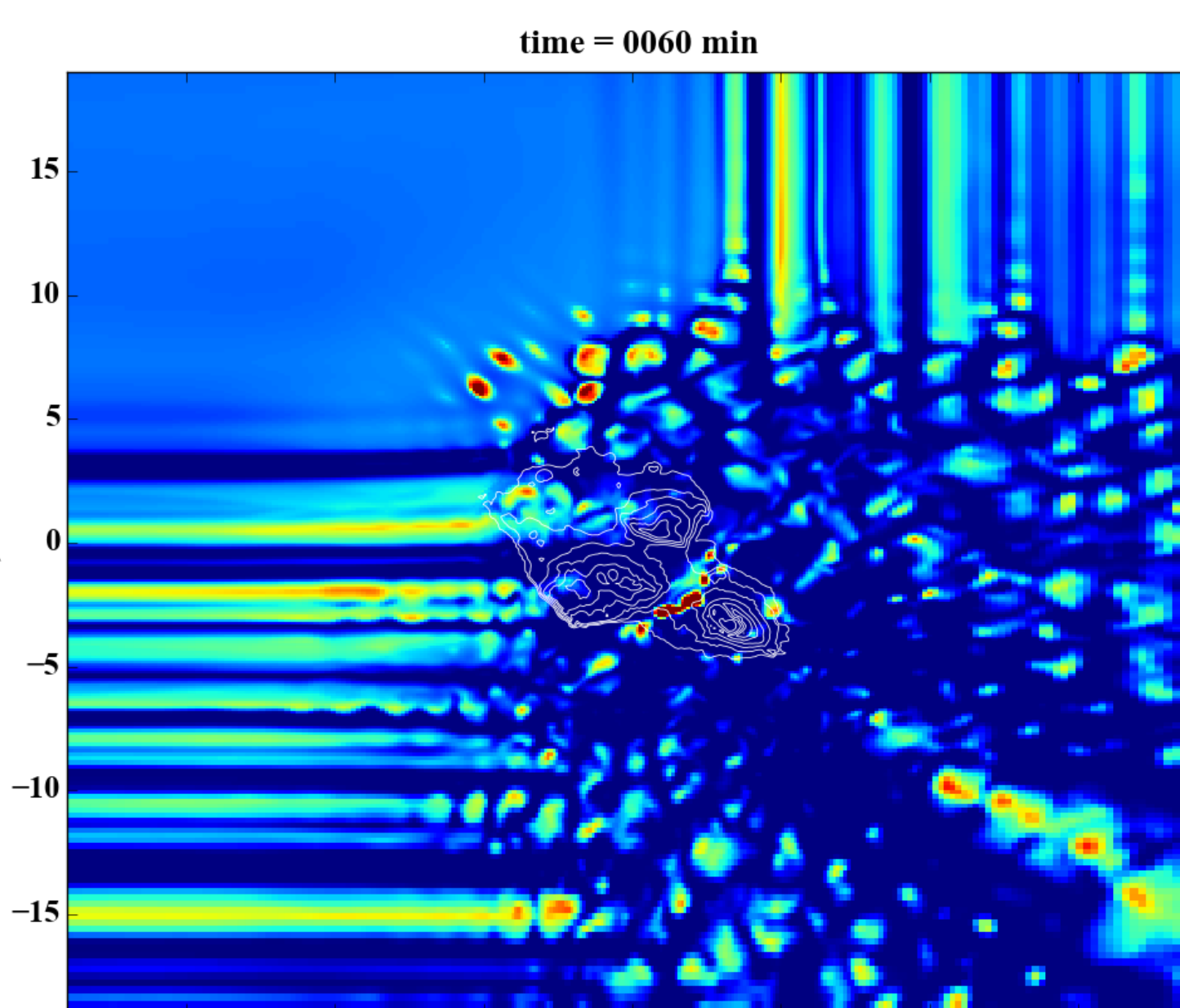
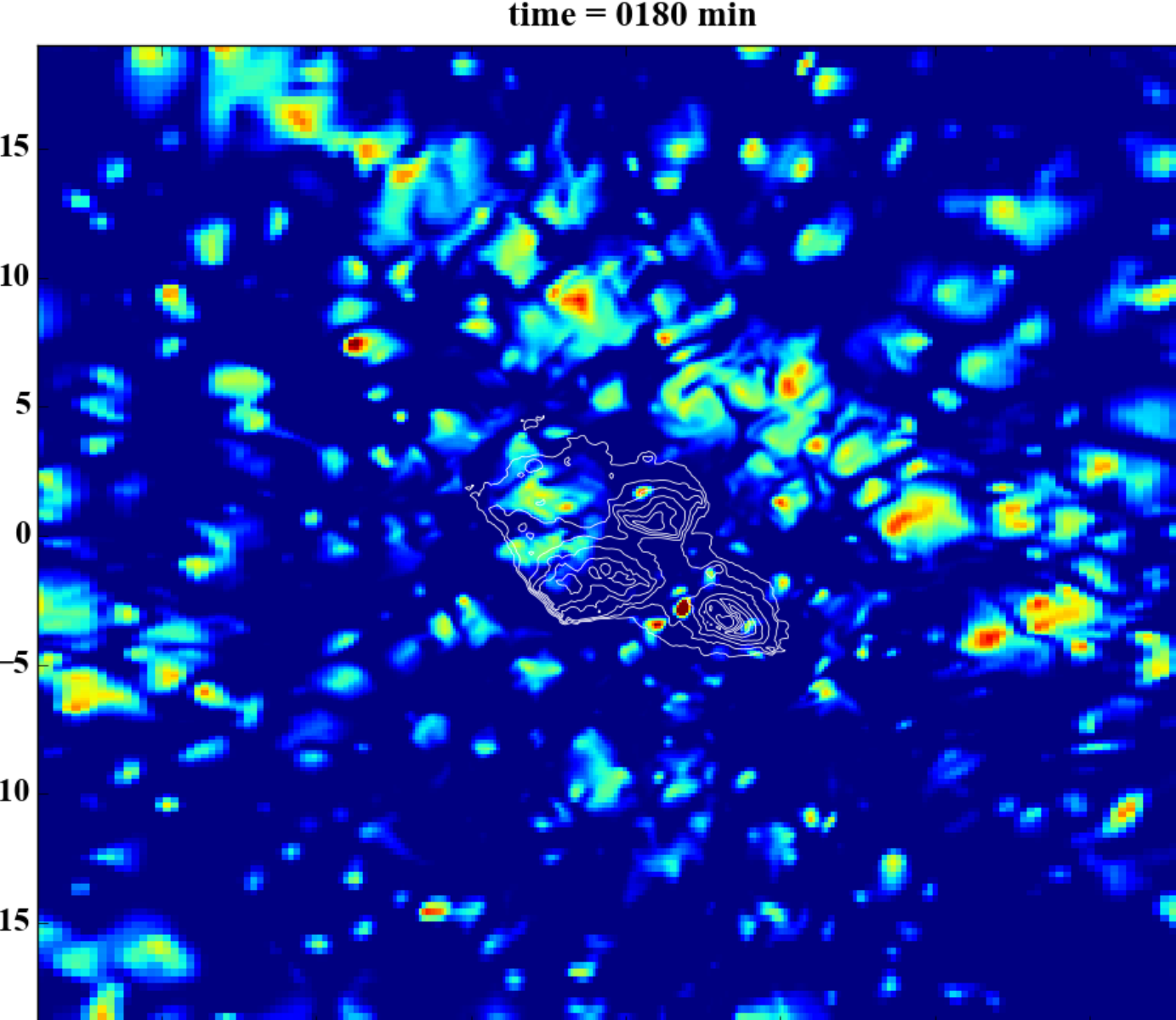
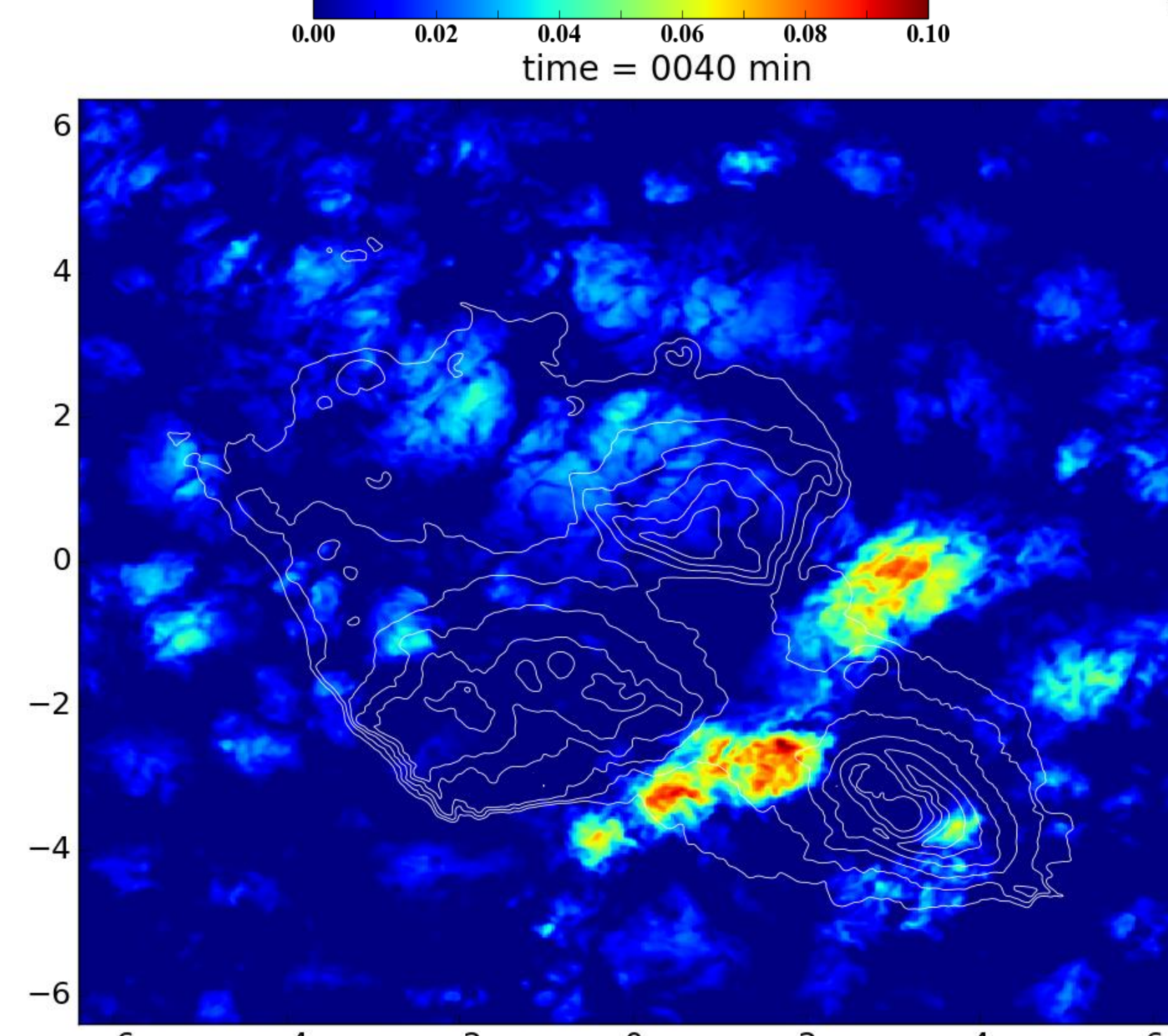
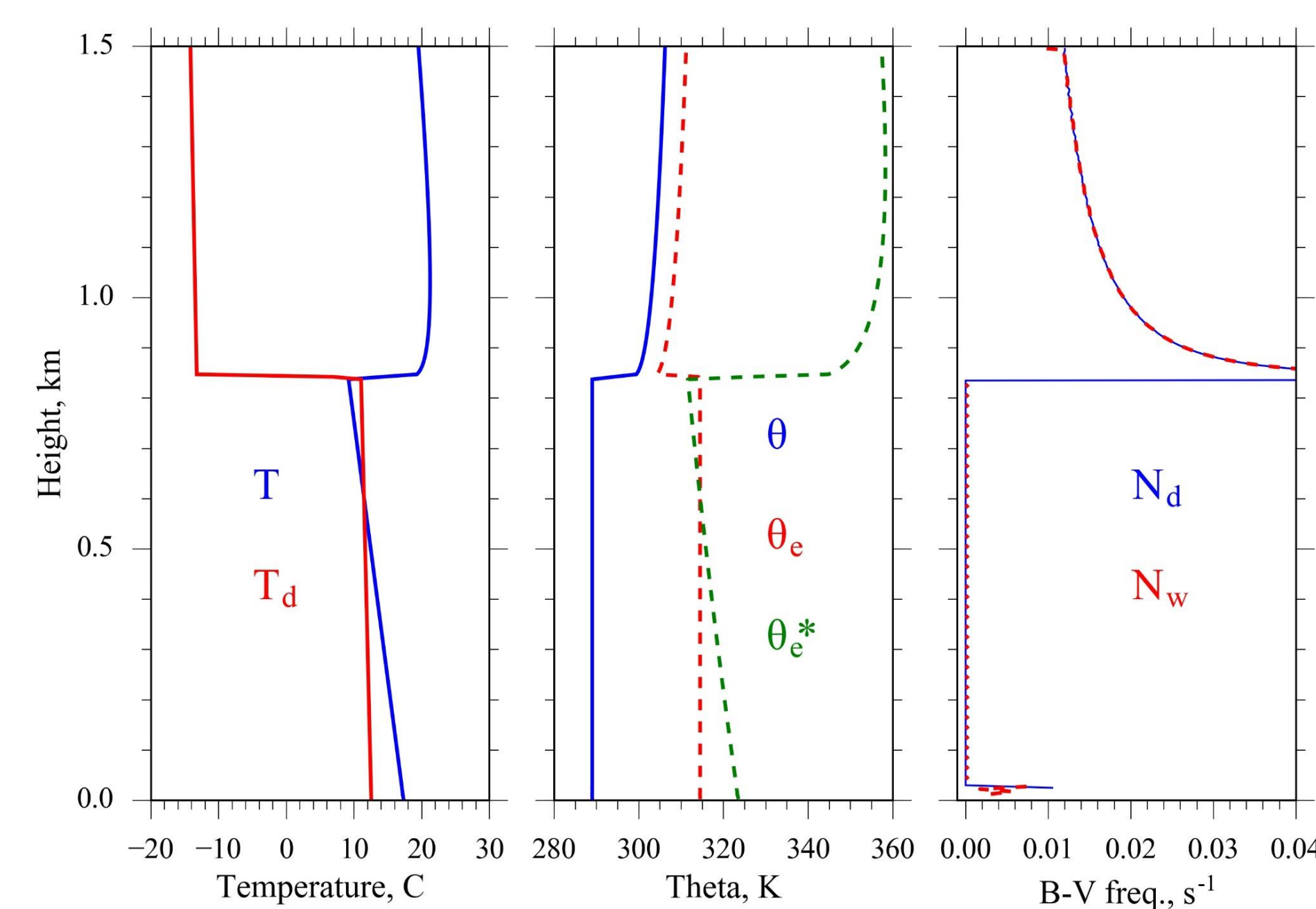
Three-dimensional Cloud Model 1 (CM1) experiments are performed to explore the mechanical and thermal impacts of Graciosa Island on the sampling of oceanic airflow and cloud evolution. The aim of the project is to create a modeling framework at true measurement resolutions to couple with cloud instrument simulators to investigate the impact of the terrain on observed cloud climatologies.

## Moist simulation setup:

- Grid size: 12.8 x 12.8 x 1.5 km
- $\Delta x, \Delta y = 20$  m;  $\Delta z$  (stretched) = 10 m in 1 km
- $\Delta t = 0.1$  s (0.2 s with adaptive time-stepping)
- Initial wind profile is vertically uniform with the values of  $u = 7, v = -5.5$  m/s.
- Doubly periodic boundary conditions
- Kessler microphysics
- Turbulence is initiated with  $[-0.1, 0.1]$  K random perturbations in base theta field.
- Surface fluxes of heat and moisture are included.
- Monin-Obukhov theory based land-surface scheme.
- 1 hour integration takes ~7h with 256 processors. (Open boundary and no surface flux run ~ 4h)
- Data output at every 10 mins (~2 GB per file).
- Single sounding initialization is done with idealized non-precipitating profile from DYCOMS RF01 case. Temperature, moisture, and stability profiles are shown below. Note that the sounding has a strong capping inversion (~10 K) at around 850 meters, which is conducive to development of marine stratocumulus below this level.

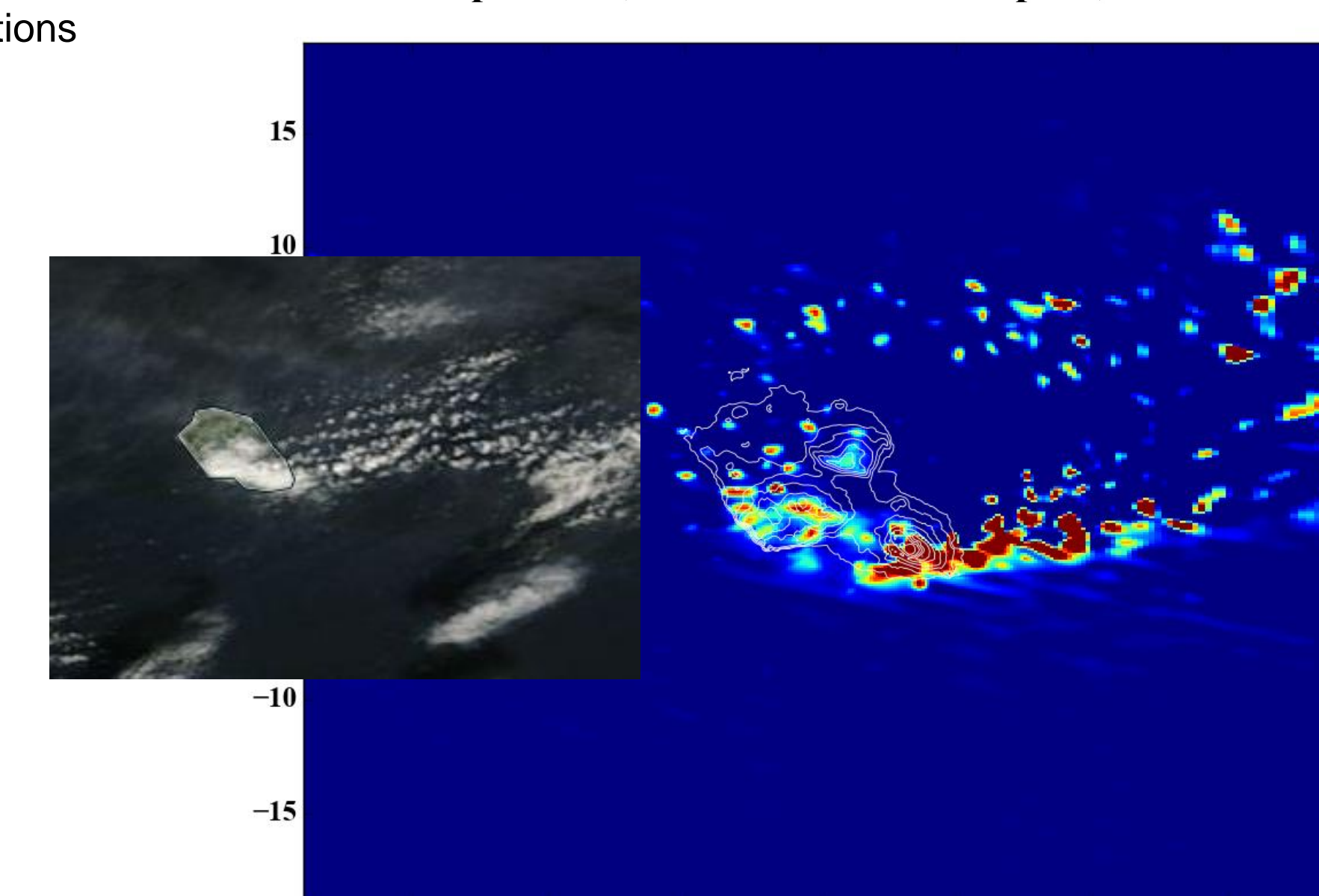
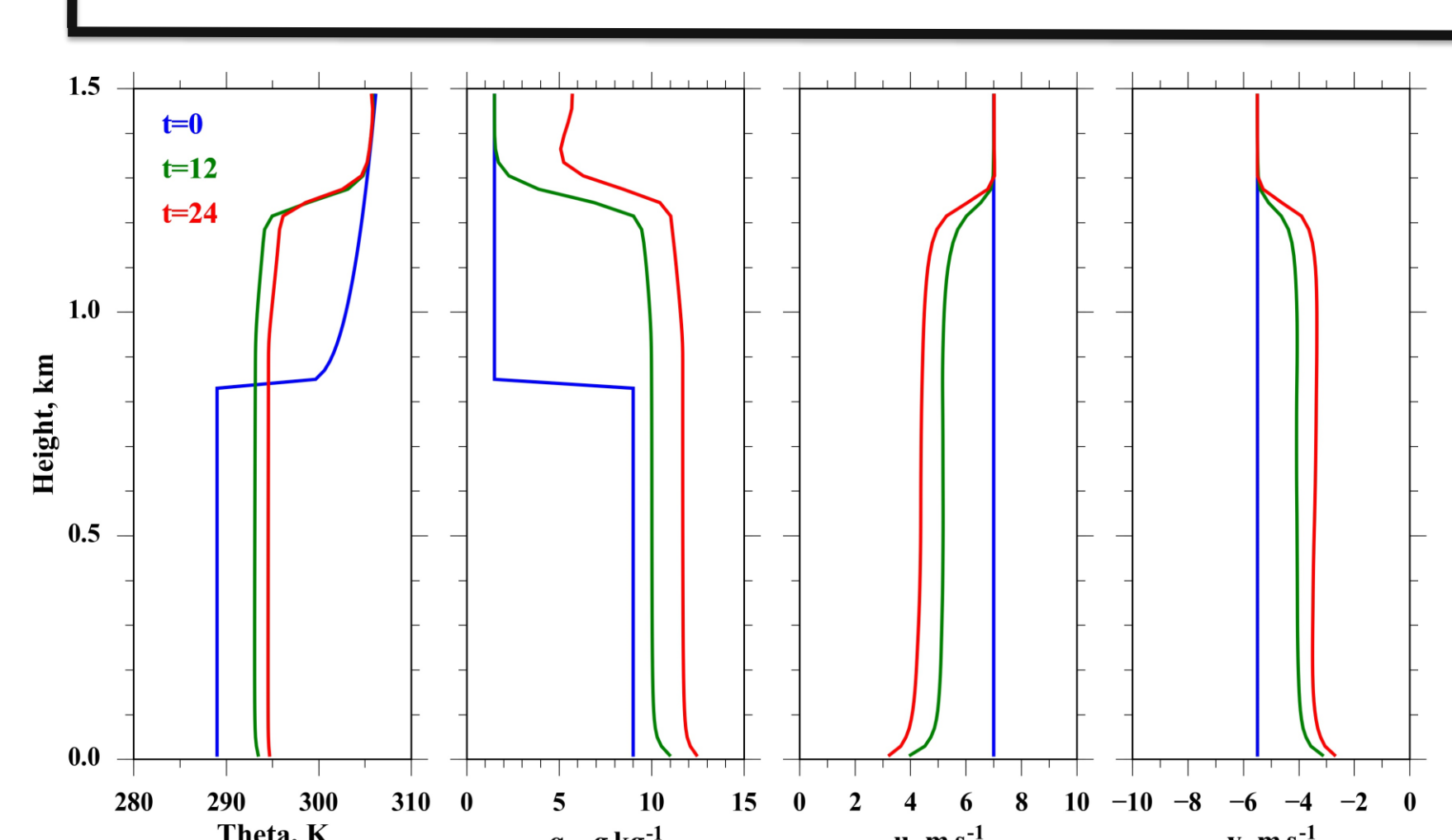
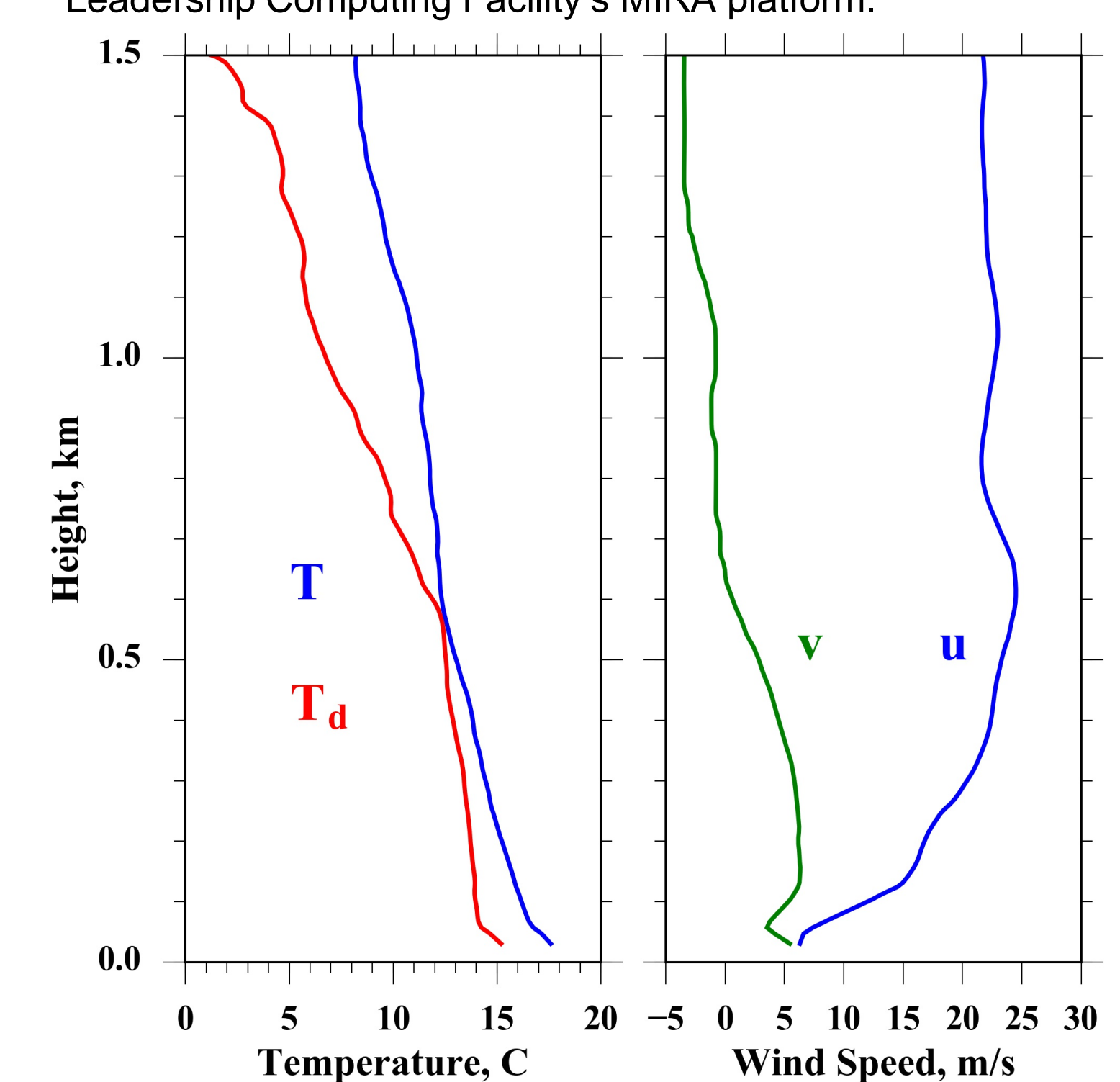
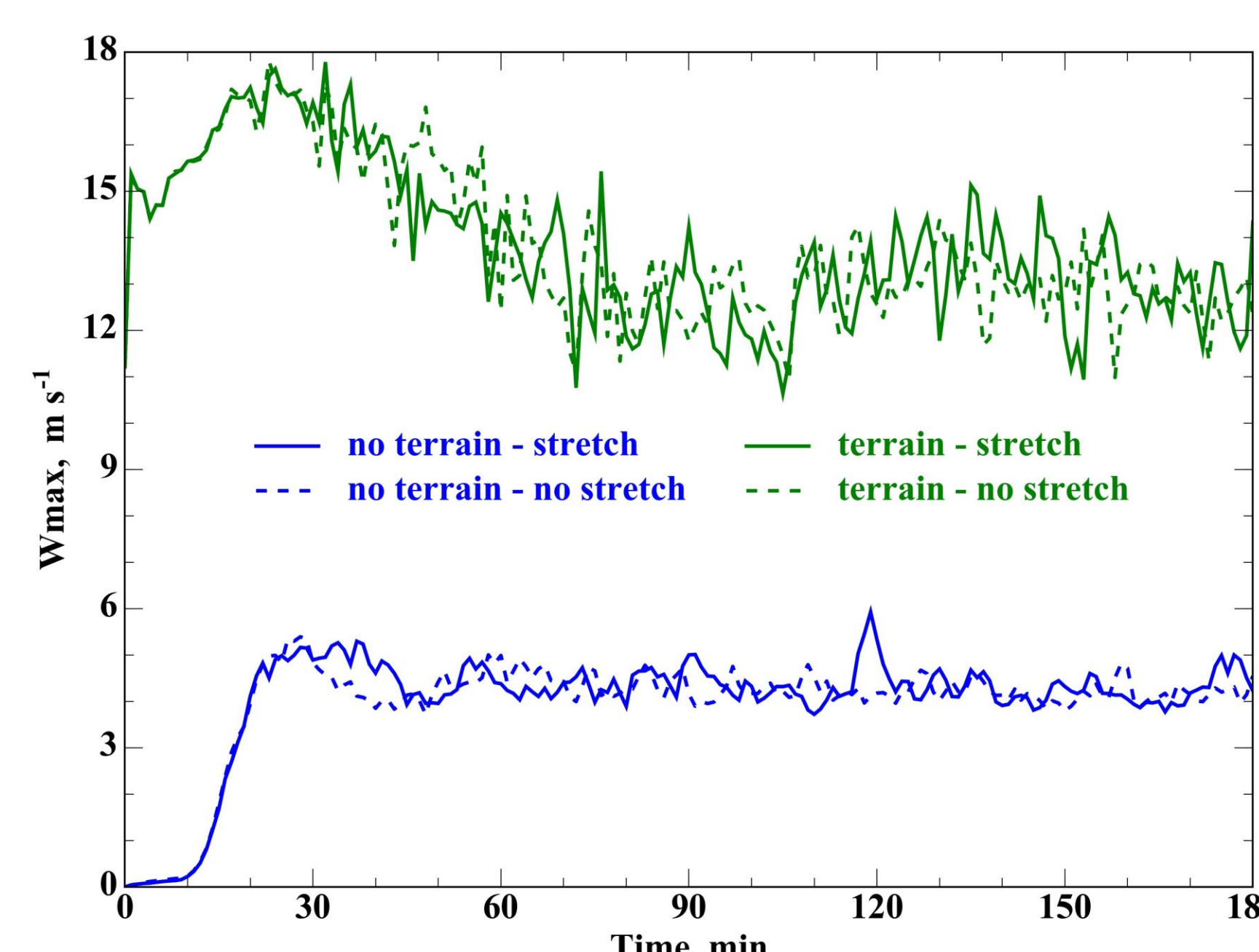
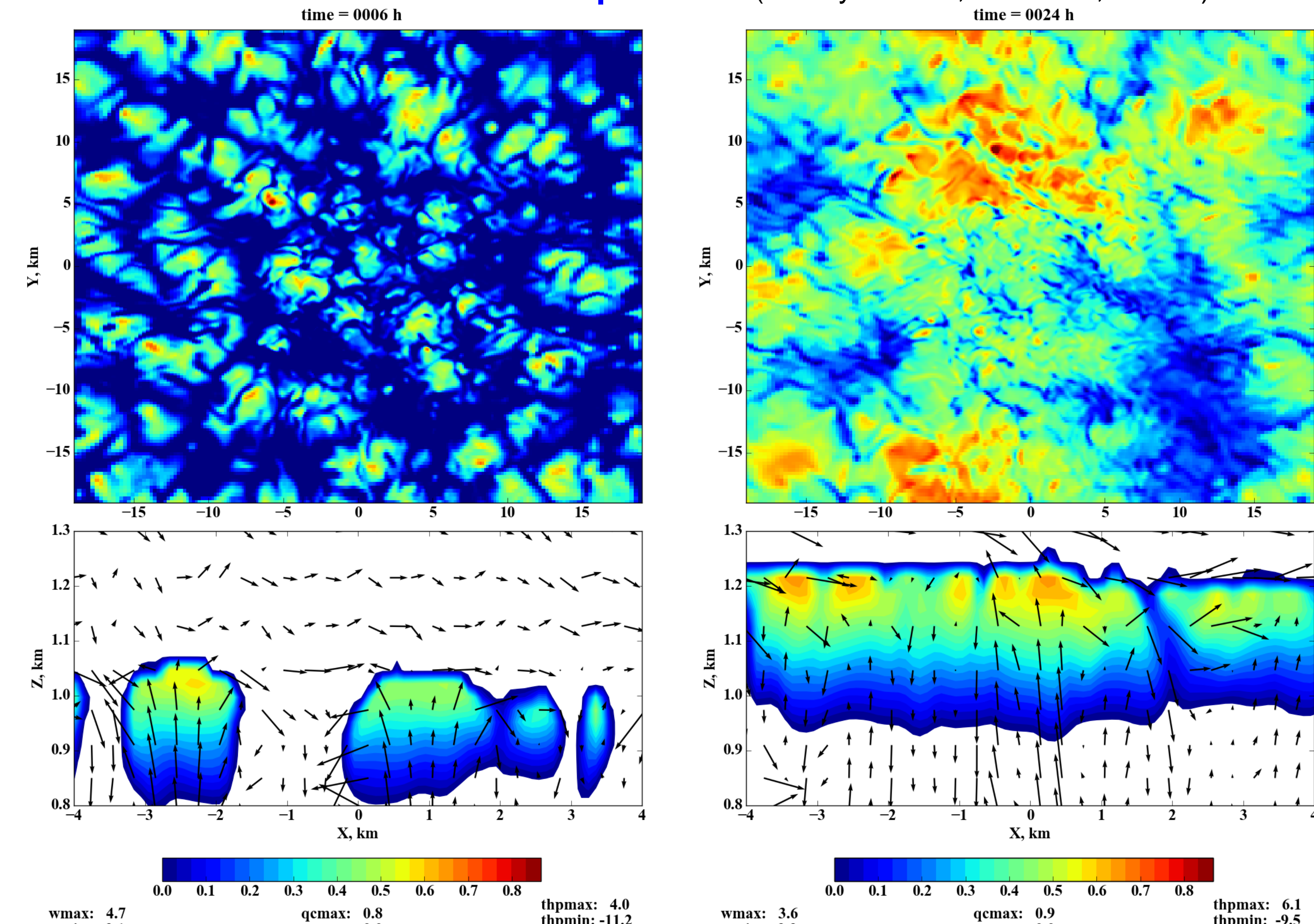


- 3D view of topography and iso-surfaces of cloud water mixing ratio (qc; 0-1 g/kg) along with select wind vectors at the surface at  $t = 40$  min.
- CM1 employs a terrain-following grid system in the vertical which is reflected on the placement of wind field. The maximum upward motion ( $\sim w_{max} = 20$  m/s) is located near the steepest upslopes as indicated from large vectors at the surface.
- The relatively deeper cloud layer over the island is produced as a result of mountain waves that are triggered by orographic forcing.



12 UTC February 03, 2017 sounding from the ENA site. Note the high (> 20 m/s) westerly winds above the cloud base. Such large advection speeds worsen the viability of periodic boundary conditions in a relatively small single domain. A computational domain larger than 100 km might alleviate this problem, yet such grid setup requires a supercomputer on the scale of Argonne Leadership Computing Facility's MIRA platform.

## Flat-terrain with radiation experiment ( $\Delta x = \Delta y = 100$ m; $\Delta z = 20$ m; $\Delta t = 1$ s)



- Horizontal distribution of column averaged qc showing the cumulus ( $t=6$ h) to stratocumulus ( $t=24$ h) transition at top panels. Cross-sections (at  $y=0$ ) of qc show near 200 m growth of boundary layer between the snapshots along with the circulation patterns at bottom panels. Note that the system reaches to convective-radiative equilibrium at about 12 h as shown in the vertical profiles.
- Radiation scheme is based on NASA-Goddard longwave and shortwave radiation codes, initialized for 22 February 2017 at 19:30 and called at every 5 mins. (Skin temperature = 26 C). Morrison microphysics (double-moment) is used to explicitly resolve the cloud structure. Initial cloud droplet number concentration ( $N_d$ ) is 100 cm<sup>-3</sup>. A sensitivity simulation with  $N_d=250$  cm<sup>-3</sup> shows a relatively small change in boundary layer depth, which implies that flow dynamics dominates over microphysical processes.
- The equilibrium cloud field can be used as an input to terrain simulations to study the island effects further.

- Overall, sub-cloud column is warmed and moistened; and kinetic energy is reduced as wind magnitude is lowered about a factor of two within 24h period. Note that, evident near the inversion transitions, the boundary layer growth is stabilized at 12 h into the simulation.