

1. Introduction

Many global models have long-standing biases in top-of-the-atmosphere shortwave radiation over the subtropical oceans. The cloud albedo is typically too low over the maritime stratocumulus decks and too high over the regions of trade cumulus, with consequences for the model's radiation budget and response of the ocean in coupled models. The cause of radiation biases in the stratocumulus areas is generally consistent with a lack of cloud cover and liquid water path and is reasonably straight-forward to interpret (though by no means easy to remedy). However, in the trade cumulus regime the clouds in many models are still too reflective, yet comparison against satellite observations indicates that cloud fraction and/or liquid water path (LWP) are underestimated.

Here, we draw on ground-based observations from the MAGIC campaign to complement the view from satellite, and gain a better understanding of the underlying causes for trade cumulus clouds to be “too bright”, despite the apparent lack of cloud cover and water compared to satellite observations.

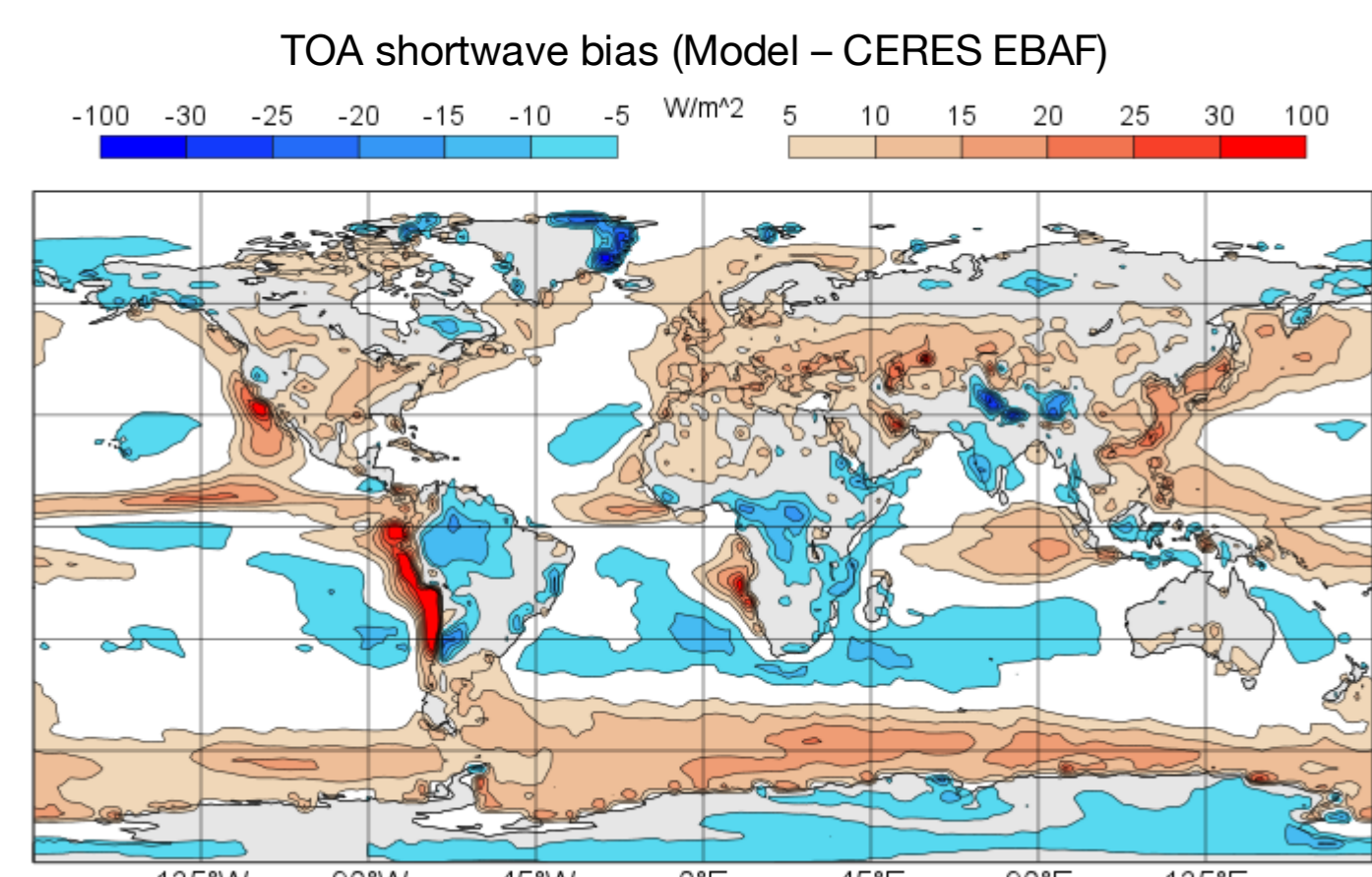


Figure 1 Model bias in TOA shortwave radiation compared to CERES EBAF observations for an ensemble of long-range runs covering 30 years. Blue shades indicate areas where too much sunlight is reflected back to space.

The map above illustrates the global scale of the problem, and that apart from MAGIC, few permanent observational sites are ideally located to tackle it.

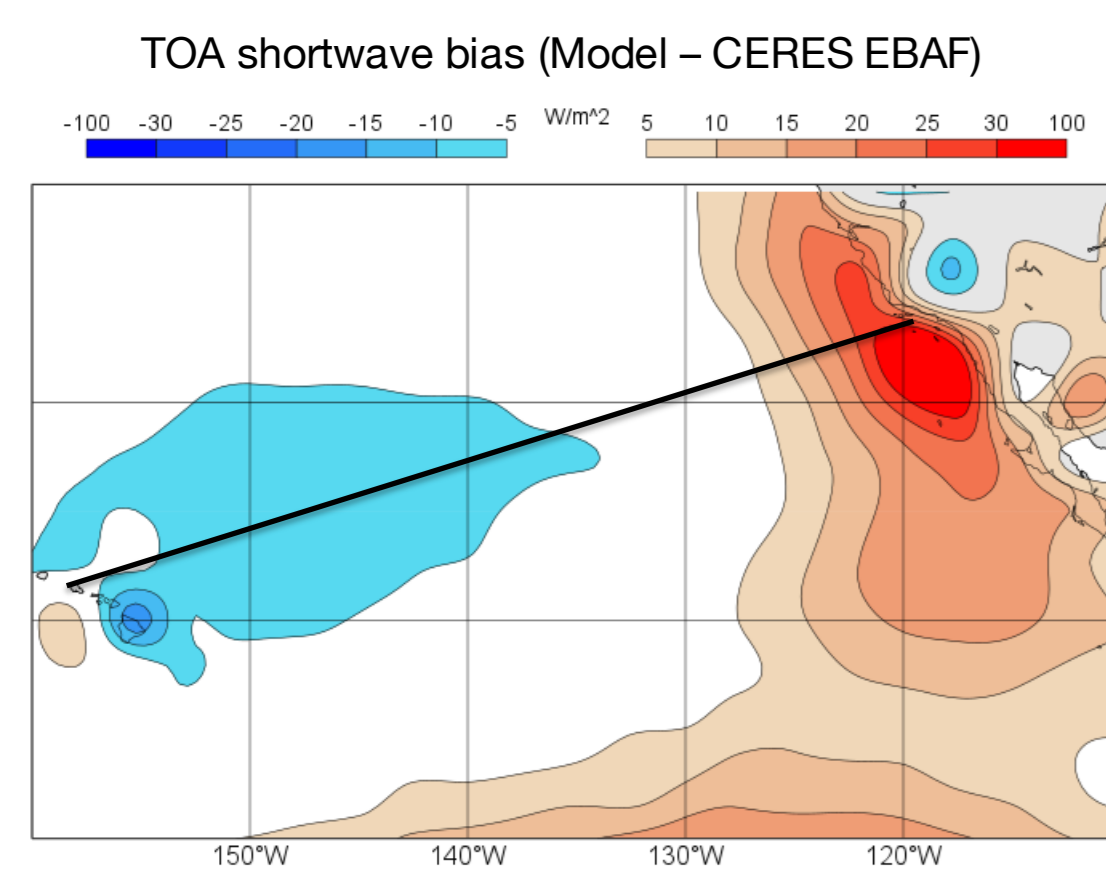


Figure 2 Highlighting the area of interest in the North East Pacific. The approximate MAGIC ship track between LA and Hawaii is marked.

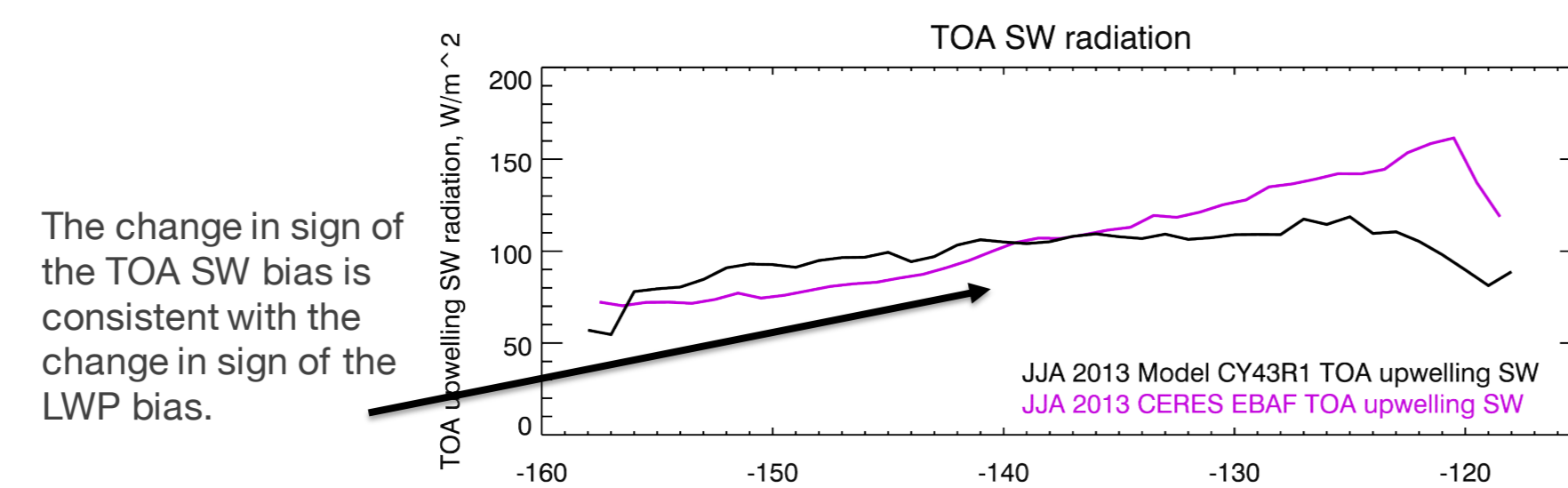
References:
 Ahlgrimm, M. and Forbes, R.M., 2016. Regime dependence of cloud condensate variability observed at the Atmospheric Radiation Measurement Sites. *JGRMS*, 142(697), pp.1605-1617.
 McGibbon, J and C Bretherton, 2017. Skill of ship-following large-eddy simulations in reproducing MAGIC observations across the Northeast Pacific stratocumulus to cumulus transition region, submitted to JAMES.
 Menon, S., Genio, A.D.D., Koch, D. and Tseloudis, G., 2002. GCM simulations of the aerosol indirect effect: Sensitivity to cloud parameterization and aerosol burden. *J. Atmos. Sci.*, 59(3), pp.692-713.
 Morcrette, C. J. (2012). Improvements to a prognostic cloud scheme through changes to its cloud erosion parameterization. *Atmosph. Sci. Lett.*, 13: 95-102. doi:10.1002/asl.374
 Painemal, D., Greenwald, T., Cadeddu, M. and Minnis, P., 2016. First extended validation of satellite microwave liquid water path with ship-based observations of marine low clouds. *GRL*, 43(12), pp.6563-6570.

Acknowledgements:
 Many thanks to Maria Cadeddu and David Painemal for their help in using the 3-channel microwave radiometer and MODIS/CERES data products, and to Jeremy McGibbon and Chris Bretherton for supplying the ship-following forcing.

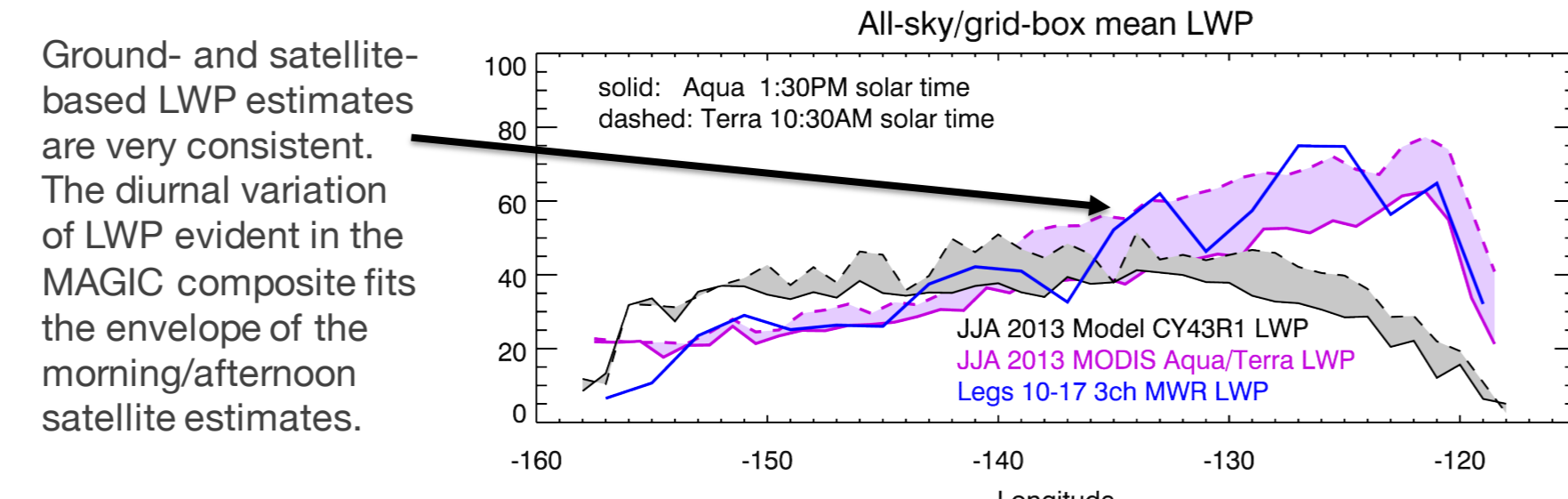
2. What do the observations tell us?

Satellite observations

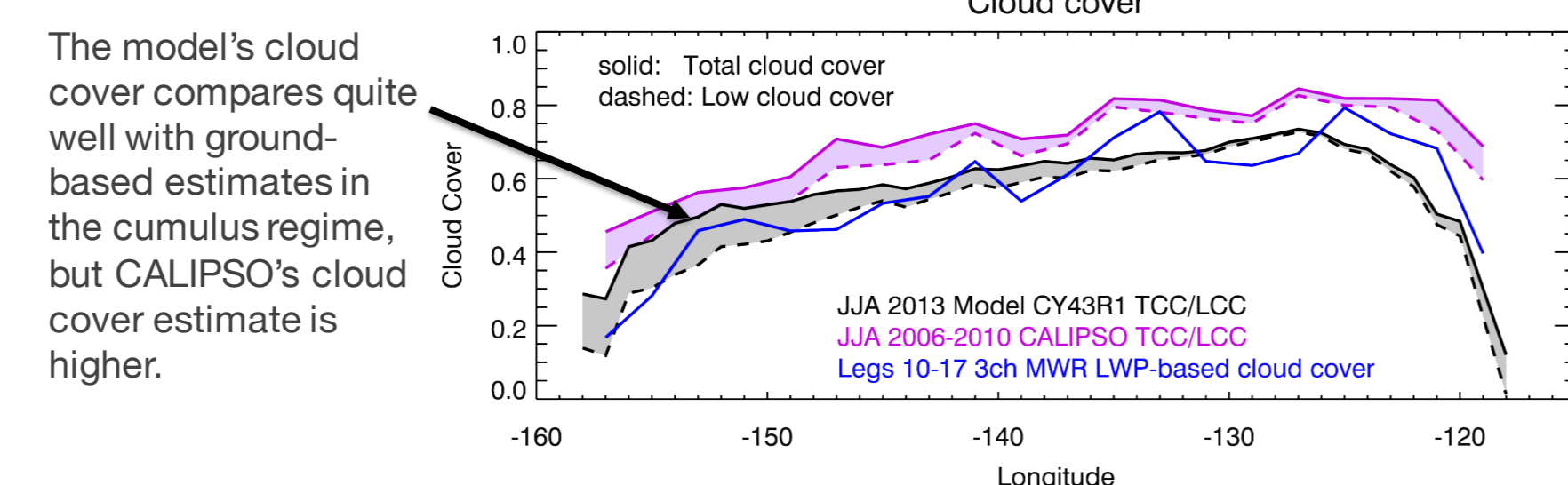
Gridded and averaged for JJA 2013, shown along average MAGIC ship track, and corresponding JJA 2013 average model data



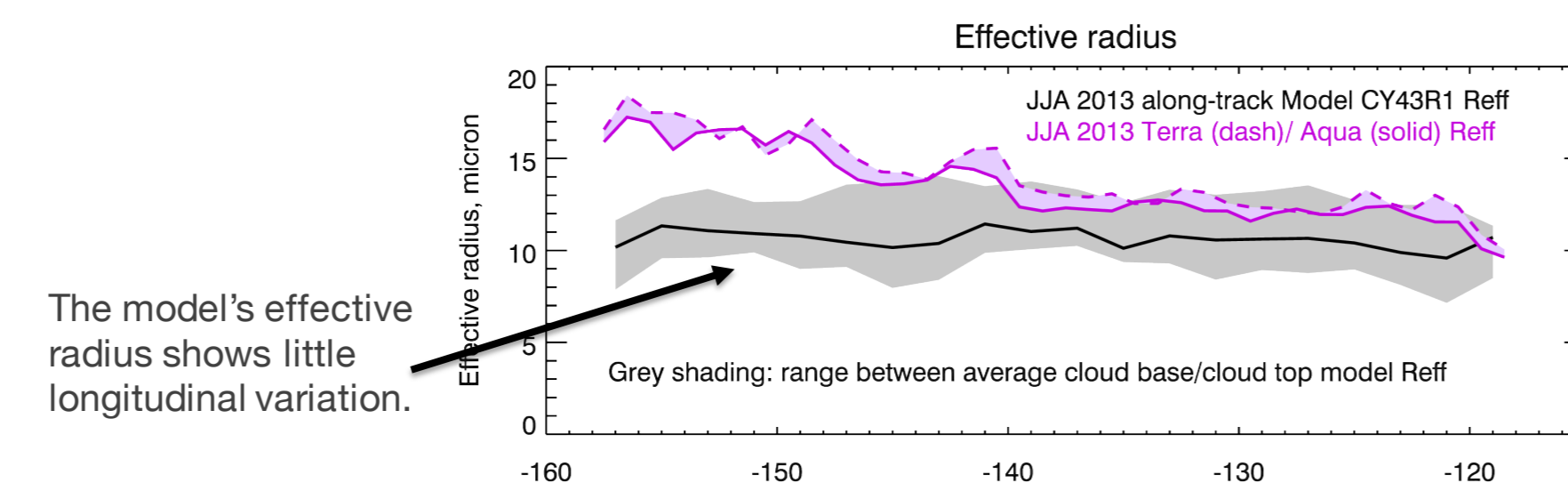
The change in sign of the TOA SW bias is consistent with the change in sign of the LWP bias.



Ground- and satellite-based LWP estimates are very consistent. The diurnal variation of LWP evident in the MAGIC composite fits the envelope of the morning/afternoon satellite estimates.



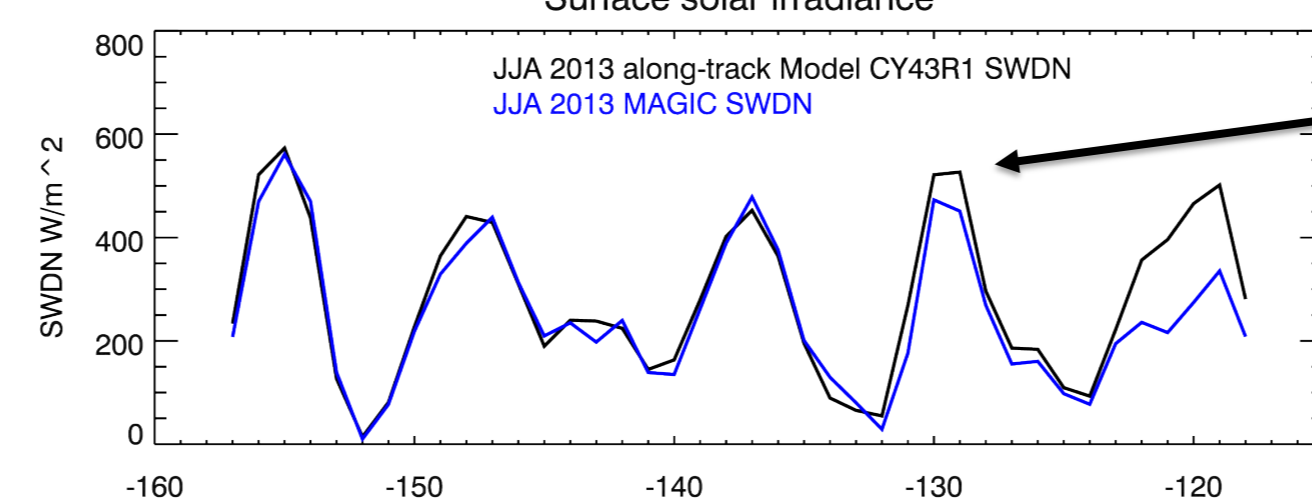
The model's cloud cover compares quite well with ground-based estimates in the cumulus regime, but CALIPSO's cloud cover estimate is higher.



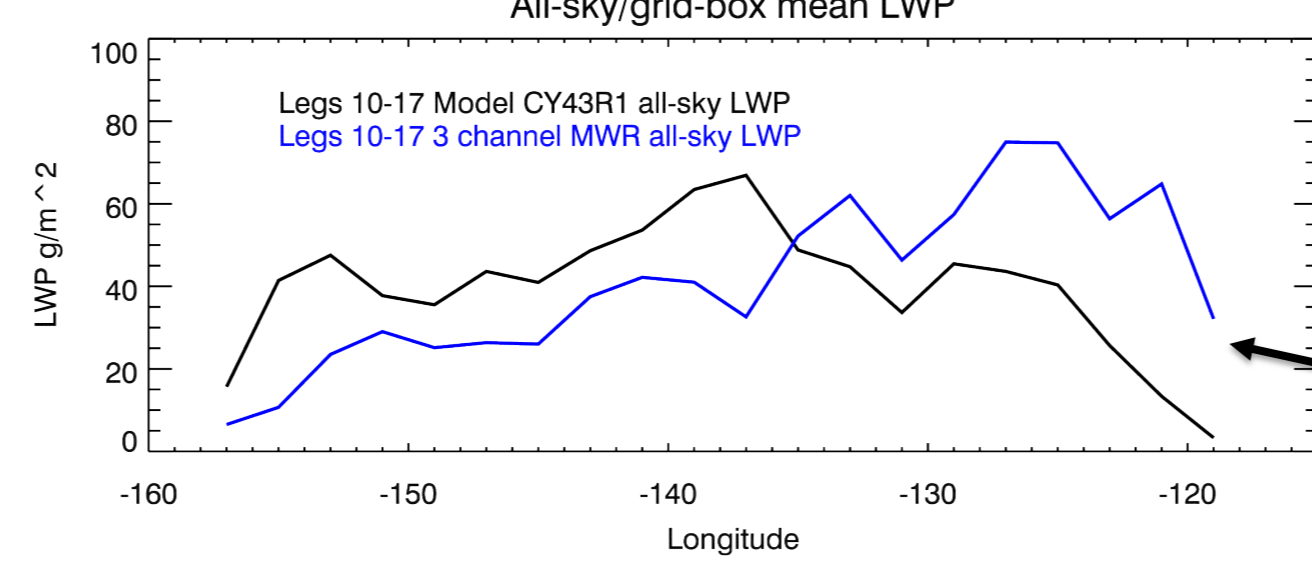
The model's effective radius shows little longitudinal variation.

MAGIC ship-based observations

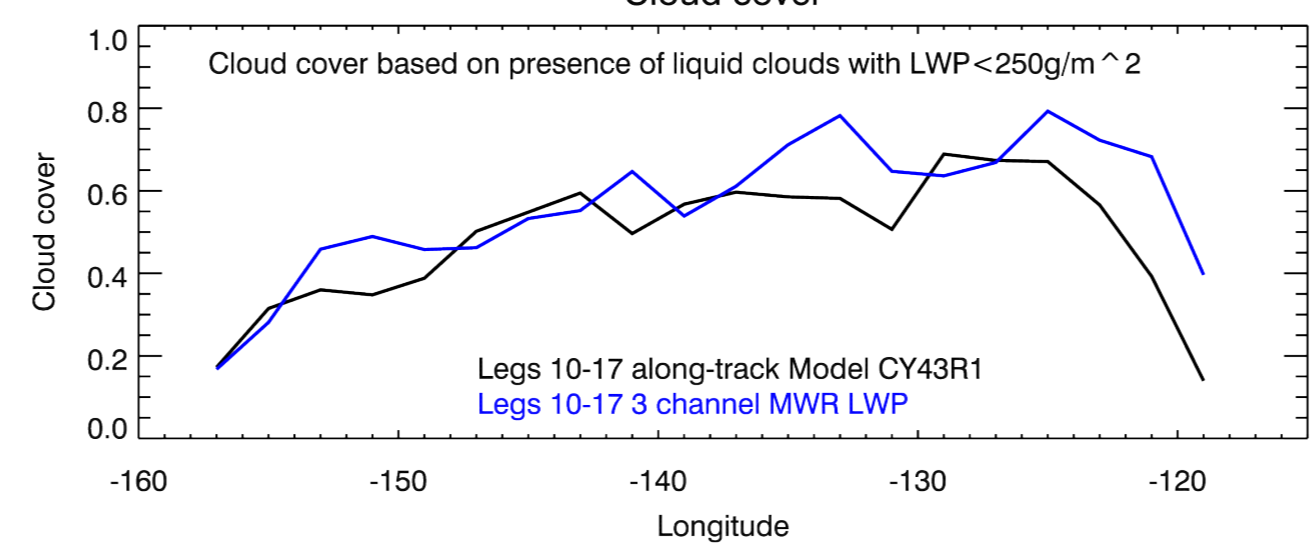
Averaged over legs 10-17 (mid-May through August 2013), model data matched in time and space to each ship track, then averaged over legs 10-17



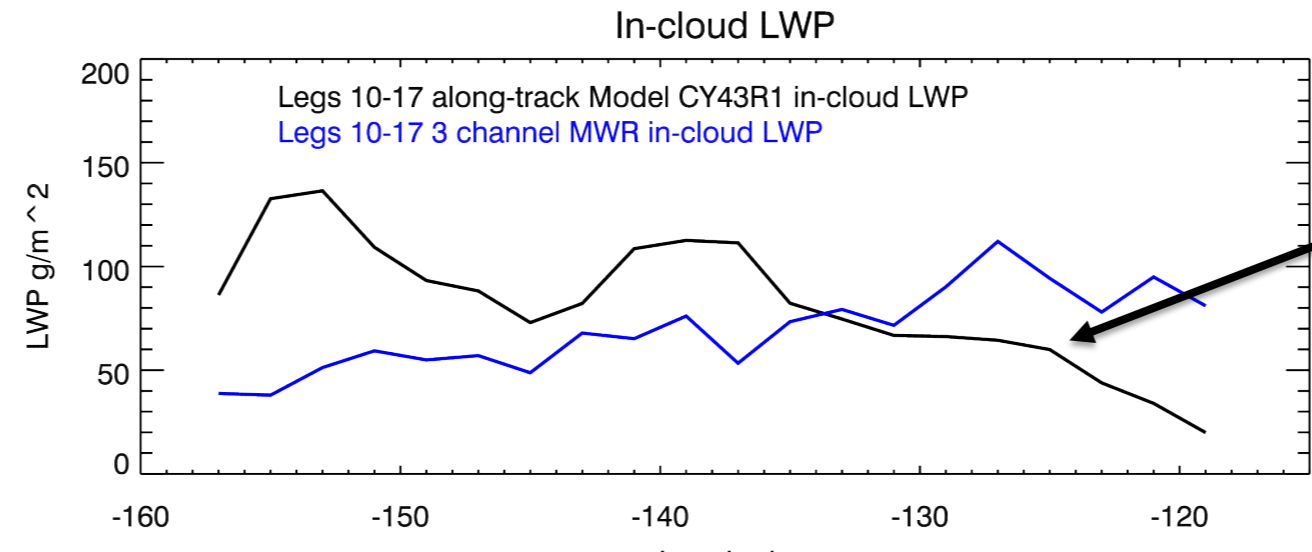
Due to consistent ship departure times, the 8-track composite of surface shortwave radiation still shows a pronounced diurnal cycle, making it more difficult to illustrate the model bias.



The model overestimates the all-sky LWP in the cumulus regime, and underestimates it in the stratocumulus regime.



The radiation scheme calculates cloud albedo based on the in-cloud condensate amount. The in-cloud LWP shows even the discrepancy between model and observations even more clearly.



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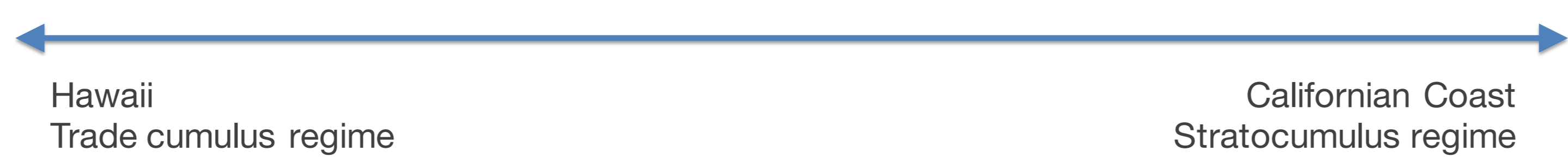


Figure 3: (Left) JJA 2013 mean satellite observations compared to the model along the average MAGIC ship track. (Right) Ship-based observations for MAGIC Legs 10-17 compared to model output matched in time and space.

3. Key insights

- The all-sky LWP retrieved from ARM's 3-channel microwave radiometer agrees well with LWP estimates retrieved from MODIS (using the 3.7µm channel) in combination with CERES. Compared with these observations, **the model overestimates LWP in the trade cumulus regime**, consistent with the sign of the shortwave radiation bias. Previous model evaluations relied on LWP retrieved from SSM/I. However, the SSM/I estimates are biased high in the cumulus regime (Painemal et al. 2016) due to ambiguities in radiance attribution to clouds vs. water vapour vs. the ocean surface. This likely explains the apparent inconsistency between radiation and LWP biases.
- Cloud cover derived from CALIPSO is around 5-10% greater than cloud cover estimates from ground-based instruments (ceilometer and MWR). The **model's cloud cover is in reasonable agreement** with the MAGIC observations in the transition and cumulus regimes, but underestimated near the Californian coast. The model's apparent underestimate of cloud cover compared to CALIPSO is of the same order as the disagreement between ground- and satellite based observations.
- To **improve the model's cloud albedo, LWP must be reduced in the cumulus regime while maintaining cloud cover.**
- A more realistic representation of **cloud effective radius could enhance** the contrast in cloud albedo between the two cloud regimes.

4. SCM experiments – tackling the problem

The LES forcing dataset developed by McGibbon and Bretherton (2017) for MAGIC Leg 15a has been adapted to drive the ECMWF single column model, and provides the opportunity to test model changes that address the challenges identified to the left.

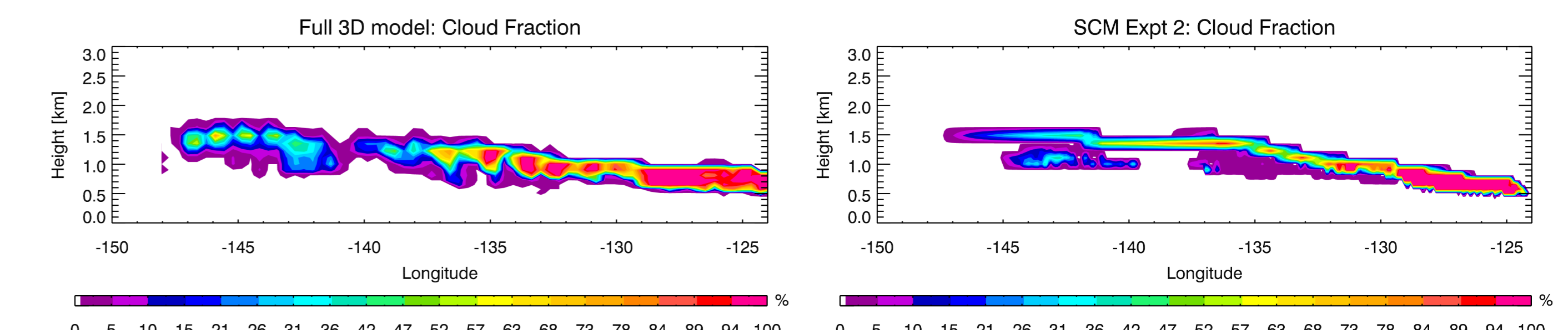


Figure 4: Model cloud fraction along MAGIC Leg 15a from (left) the full forecast model, and (right) the single column model.

1) Warm rain processes and cloud erosion

The amount of liquid retained in the cloud is the result of a delicate balance between processes that act as sources of liquid water, and those that deplete the cloud. We test a package of changes that impacts this balance by introducing a more realistic rain fall speed (depending on DSD) to replace a constant value. Together with some numerical changes, this alters the residence time of precipitation in a given model layer, and the rate of rain evaporation. We also test a new parameterization for cloud erosion (Morcrette 2012) which takes into account an estimate of how much cloud edge is exposed to the environment.

2) Cloud microphysical properties

The model's droplet number concentration over ocean is solely dependent on wind speed. This accounts for variation in N_d from sea spray, but does not account for variation in other aerosol species due to e.g. pollution near the coast. The CAMS climatology offers an alternative way to parameterise N_d based on climatological sea salt, organic matter and sulfate concentrations (following Menon et al. 2002), which may better capture the observed longitudinal dependence of the effective radius.

3) Cloud heterogeneity

The model does not resolve how cloud condensate varies at subgrid scales. Instead, this subgrid-scale heterogeneity is parameterized. Observations from ARM sites around the world show that overcast boundary layer clouds, such as stratocumulus, tend to be more homogeneous than cumulus clouds (Ahlgrimm and Forbes 2016). A new regime-dependent heterogeneity parameterization replaces a constant value, and impacts the in-cloud condensate amount used in the radiation scheme.

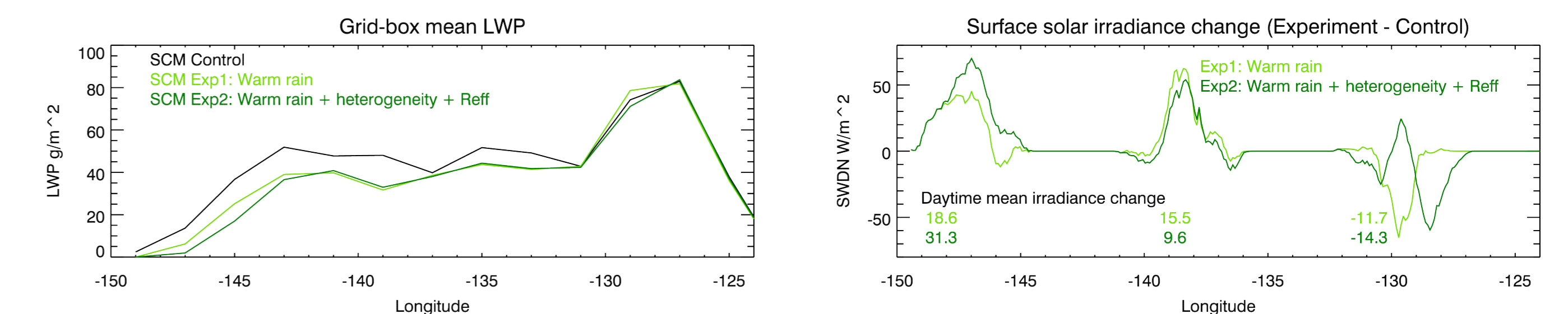


Figure 5: (Left) Grid-box mean LWP for three SCM experiments. (Right) Difference in surface solar irradiance between two SCM experiments and control. A positive difference indicates that more shortwave radiation is reaching the surface.

5. Results

The new fall speed, evaporation and cloud erosion changes **reduce LWP in the cumulus regime**, while maintaining it in the stratocumulus regime. The decreased LWP in the cumulus regime allows more solar radiation to reach the surface. Adding the regime-dependent heterogeneity parameterization and climatology-based droplet number concentration **enhances the contrast in cloud albedo** between the stratocumulus and cumulus regimes further.