

# Regional and Cloud-Scale Modeling of Tropical West Pacific Clouds Using WRF

Jimy Dudhia, Ming Chen, Bill Kuo, Hui Liu, Jeff Anderson and Steven Cavallo  
National Center for Atmospheric Research, Boulder, CO

## Introduction

The Tropical West Pacific is now a well-observed region of tropical convection, and ARM data is being used as verification for model parameterization studies. Cloud-scale modeling in this region will serve as a guide for parameterization of convection that is needed in climate models.

In this poster we focus on three improvements being made in WRF physics for applications in regional climate and cloud-scale modeling, (1) model-top clear-sky downward longwave radiation, (2) microphysics-radiation assumptions, (3) convective parameterization sensitivity

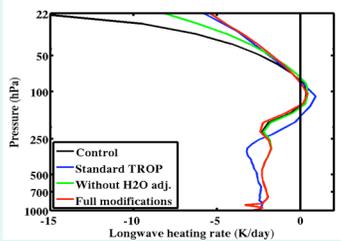
## 1. Model-top Radiation Improvement

Mesoscale model tops are usually at 10-50 hPa. A noticeable erroneous cooling occurs near the model top level as seen in 1D tests using the RRTM longwave scheme (below left).

This has been traced to the usual isothermal assumption for downward longwave radiation above the model top.

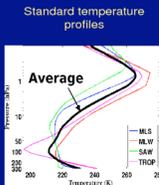
Here we improve on this assumption to remove the bias using a method of adding a more realistic lapse rate at high levels, that is applied in the RRTM longwave scheme (black and green curves below).

An added improvement comes from reducing the relative-humidity assumption for levels with missing data (green and red curves below).



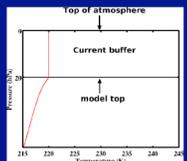
How are we correcting it?

- Tests indicate that cooling is most sensitive to assumed temperature and CO<sub>2</sub> profile
- Currently, one level is added above the model top at top of atmosphere (TOA), and temperature is assumed to be isothermal
- Bias can be corrected by assuming a mean temperature profile with several levels between the model top and TOA



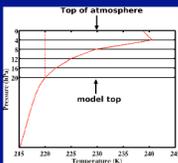
How are we correcting it?

- Currently, one level is added from model top to TOA at  $p = 0$  hPa
- Temperature is assumed isothermal
- Mixing ratios are held constant, except for O<sub>3</sub> which is reduced to  $0.6 \times O_3(\text{model top})$



How are we correcting it?

- New method adds several "buffer" layers from model top to TOA with a constant pressure interval between levels of  $\Delta p = -4$  hPa
- Results sensitive to  $\Delta p$
- Temperature profile interpolated from a new table to WRF column pressure with buffer layers appended
- Because of the stratopause located  $\sim 1$  hPa, method is not suitable for model tops under 5 hPa.



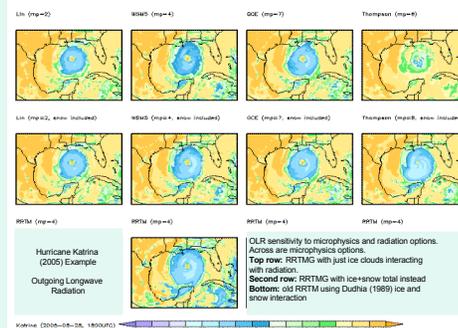
## 2. Microphysics-Radiation Interaction

The interaction between microphysics and radiation parameterizations is sensitive to the particle assumptions.

In the below examples, we see that for some microphysics options, ice clouds are not suitable for OLR calculations.

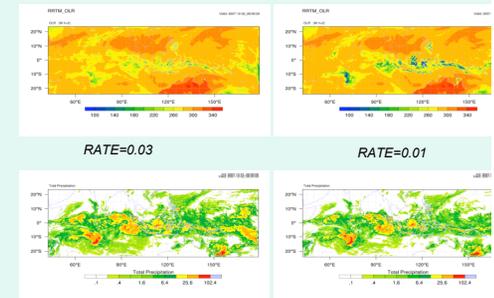
A fix for this is to consider ice+snow for radiation purposes, but ideally the particle assumptions should be consistent between microphysics and radiation.

More work is ongoing in this area and this project has a microphysical observation component (Heymsfield and Mace) that will be helpful in guiding this development.



## 3. Cumulus Parameterization

The Kain-Fritsch convective scheme is a popular scheme in regional climate simulations. One parameter in this scheme is *RATE* which represents the cloud to precipitation conversion rate in parameterized updrafts. We find sensitivity to this parameter that affects the partitioning of condensate into upper clouds and precipitation. The default value of 0.03 is modified to 0.01 to give better OLR and precipitation for tropical convection.



The above figures show that the OLR is significantly reduced and precipitation is slightly reduced when  $RATE = 0.01$  consistent with more cloud, less precipitation.

*Synergistic Activities: TWP-ICE LAM Intercomparison study. Domains 9/3/1 km shown below with Case 1 and Case 2 sample results. See also Varble et al., Zhu et al., Hagos.*

