

System



## Objectives

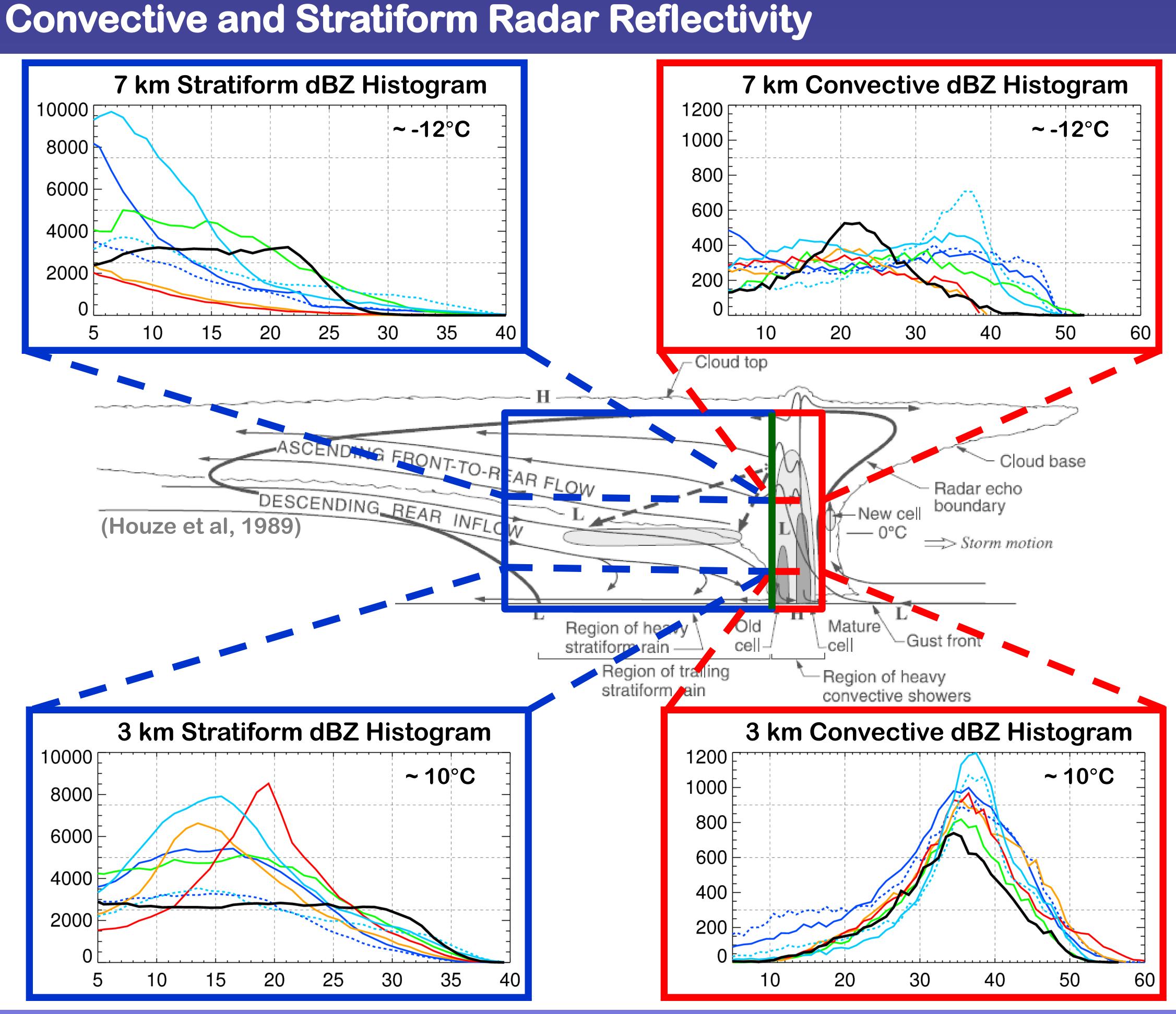
>Carefully compare observations with cloudresolving and limited area model simulations of **TWP-ICE tropical oceanic convection (Section 1 of** the comparison framework presented here).

>Investigate discrepancies between model output and observations via model intercomparison and sensitivity tests to determine reasons for differences (Section 2 in progress, not shown).

>Based on section 2 findings, improve model parameterizations, especially with regard to bulk microphysics schemes in models covering all scales (Section 3).

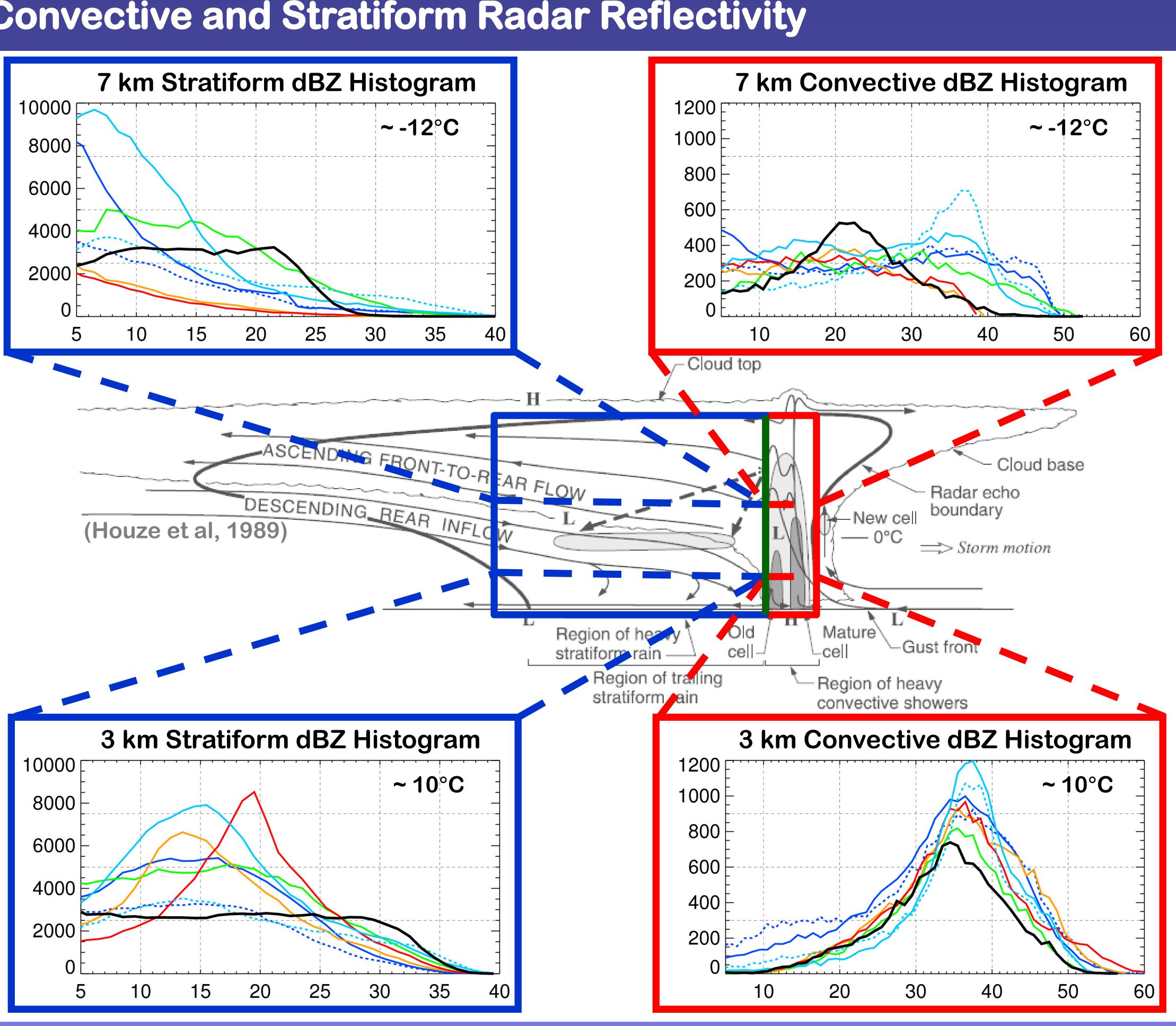
### For the active monsoon period, some stark differences are visible when comparing model with observed radar reflectivity in convective and stratiform regions at 3 and 7 km. Model updraft vertical velocity statistics are not grossly different from those observed (not shown), suggesting that the microphysics schemes are the primary reason for the differences in radar reflectivity.





Legend

**CPOL Radar Observations DHARMA 1-Moment Base DHARMA 1-Moment** Sensitivity **UKMO 1-Moment Base MESONH 1-Moment Base MESONH 2-Moment Base** SAM 2-Moment Base SAM 2-Moment Sensitivity



# **A Comparison Framework to Evaluate TWP-ICE Cloud-resolving Simulations with Observations**

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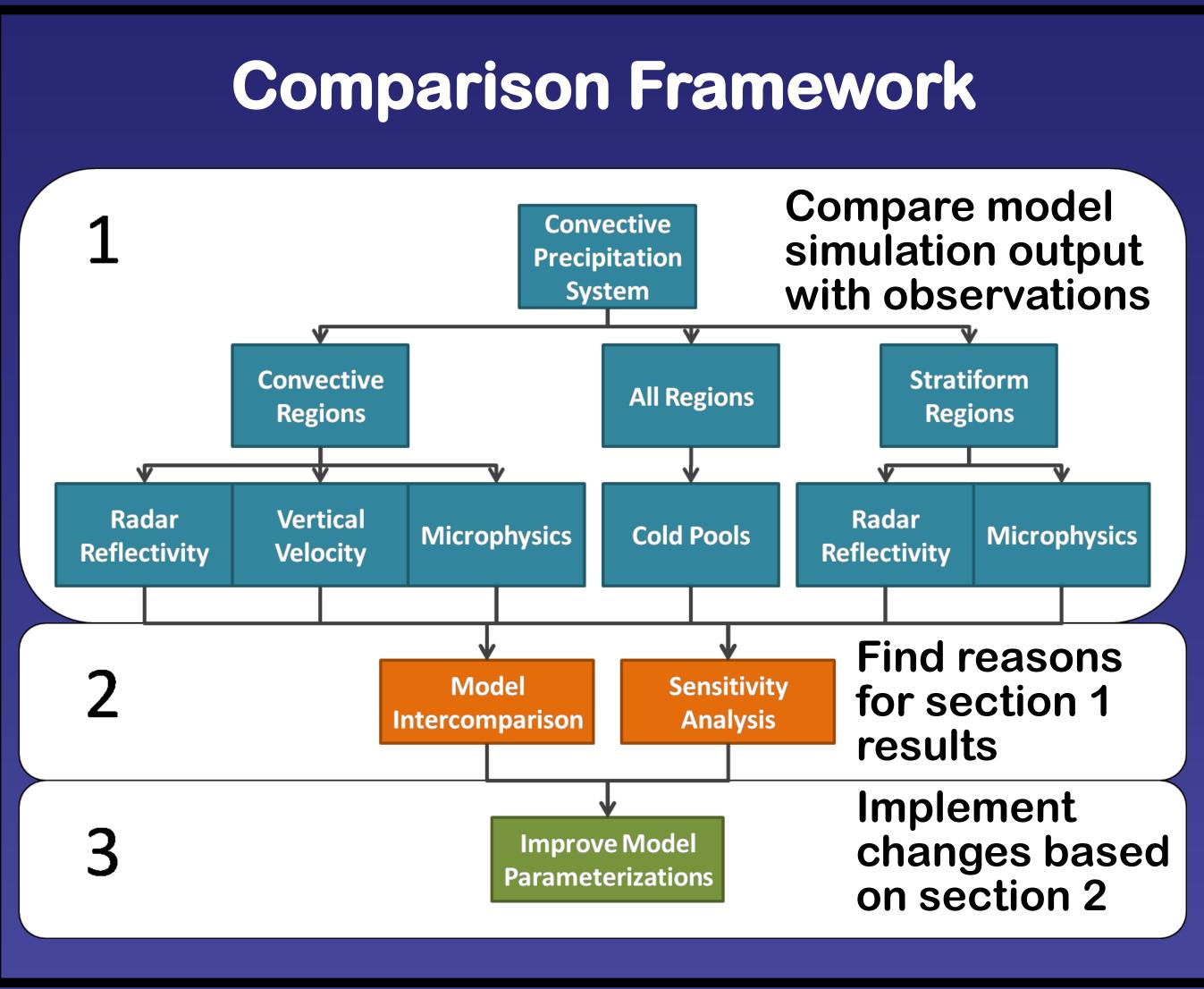
### **Data and Methods**

>The primary instrument used for observational data is the C-band polarimetric Doppler radar (CPOL). It is used for radar reflectivity, dual **Doppler velocity, and DSD retrievals.** 

>9 model simulations produced from 4 cloud-resolving models (CRMs, periodic lateral boundary conditions) and 1 land-area model (LAM, nested with open lateral boundary conditions) are being evaluated.

> Model output is degraded to CPOL horizontal resolution of 2.5 km. Convective and stratiform regions are identified using the Steiner et al. (1999) algorithm at the 3 km level.

>CPOL-CRM comparisons are performed within a pentagonal model forcing domain, whereas CPOL-LAM comparisons are performed within the 142.5 km CPOL range ring.



### **Conclusions and Future Work**

WSM6 and Thompson runs

differences.





- >The comparison framework is key to organizing numerous simulations and observations into results that are applicable to improvement of the models
- >There are some glaring differences between model simulated radar reflectivity and observations
  - >CRM convective area is too high
- >CRM convective dBZ in most models is too high in the ice and mixed phase regions
- >CRM stratiform dBZ in most models is too low
- >The two WRF (LAM) simulations (not shown) have convective dBZ that is too high, although the distributions aloft are vastly different between the
  - >LAM stratiform area is far too low, but the Thompson run has dBZ values that are far too high aloft whereas the WSM6 run is the opposite
- >Results from updraft vertical velocity comparisons suggest that microphysics schemes are likely the primary reason for the dBZ differences
- $\succ$  The next step is to dig deeper into the various model output results in an attempt to determine the degree to which the ice water content and/or the size distribution assumptions for each ice species are causing the dBZ