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Abstract: Mixing of environmental air with buoyant plumes has important implications on cloud lifecycle and subsequent impacts on atmospheric energy and water budgets. This mixing is notoriously difficult to measure particularly via remote sensing.

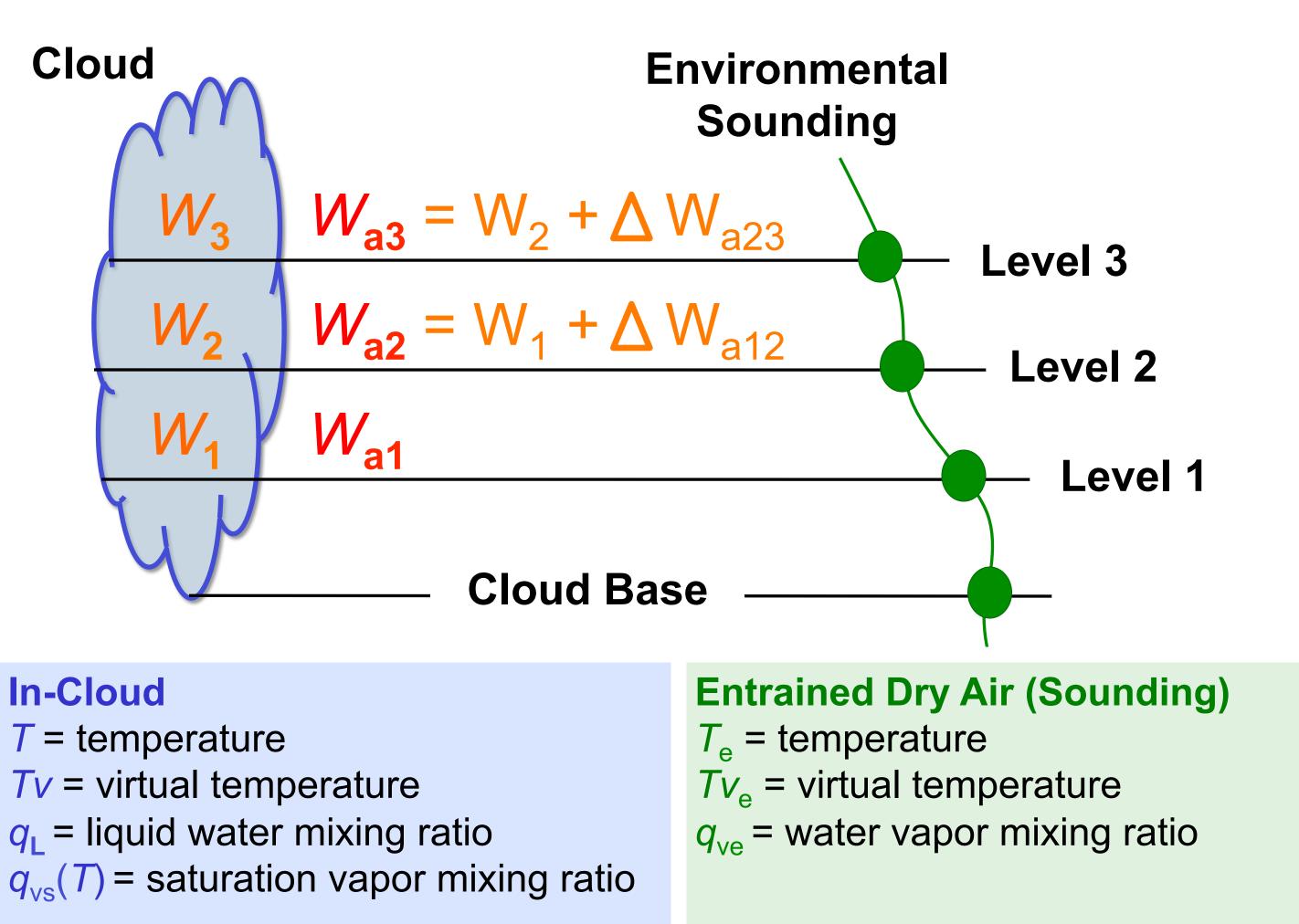
1. Method:

 The cloud grows adiabatically from cloud base and then experiences the first entrainment event and isobaric mixing at Level 1.

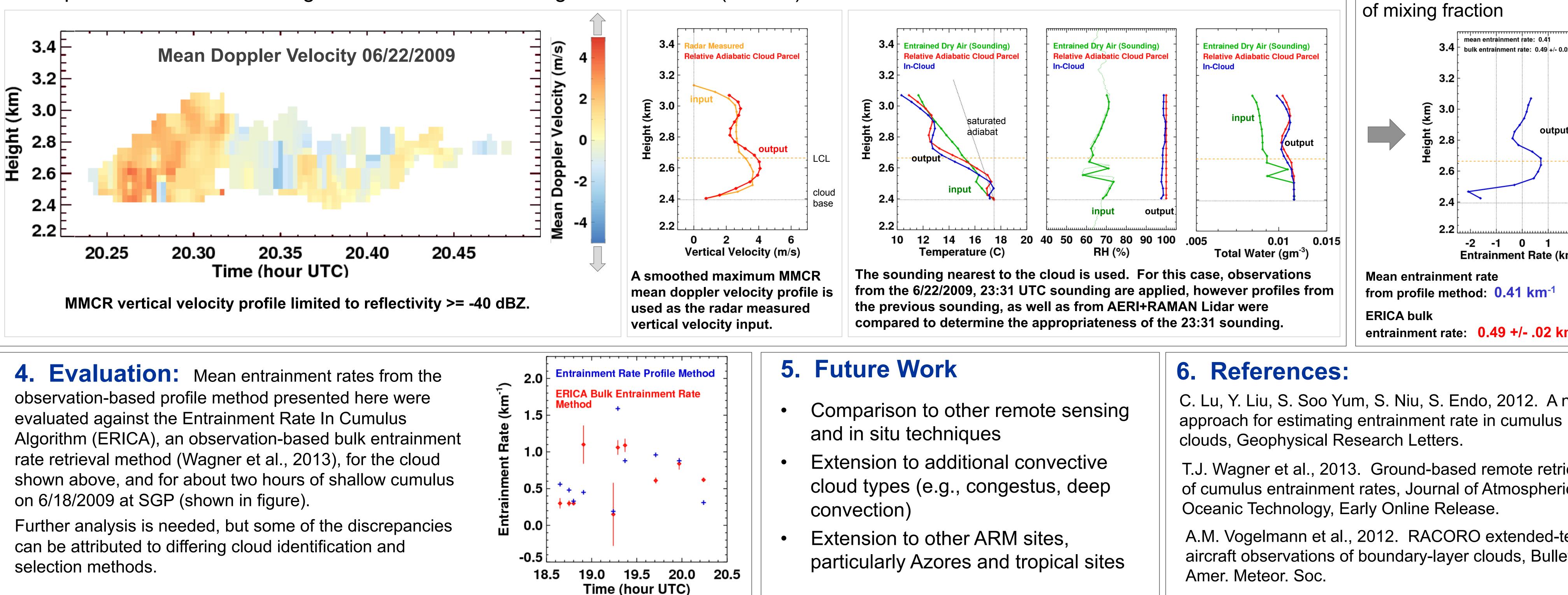
 After a new saturation is achieved during isobaric mixing, the cloud ascends adiabatically without entrainment from Level 1 to Level 2.

 It then experiences the second entrainment event and isobaric mixing at Level 2.

• The process is repeated for Level 3 and higher levels.



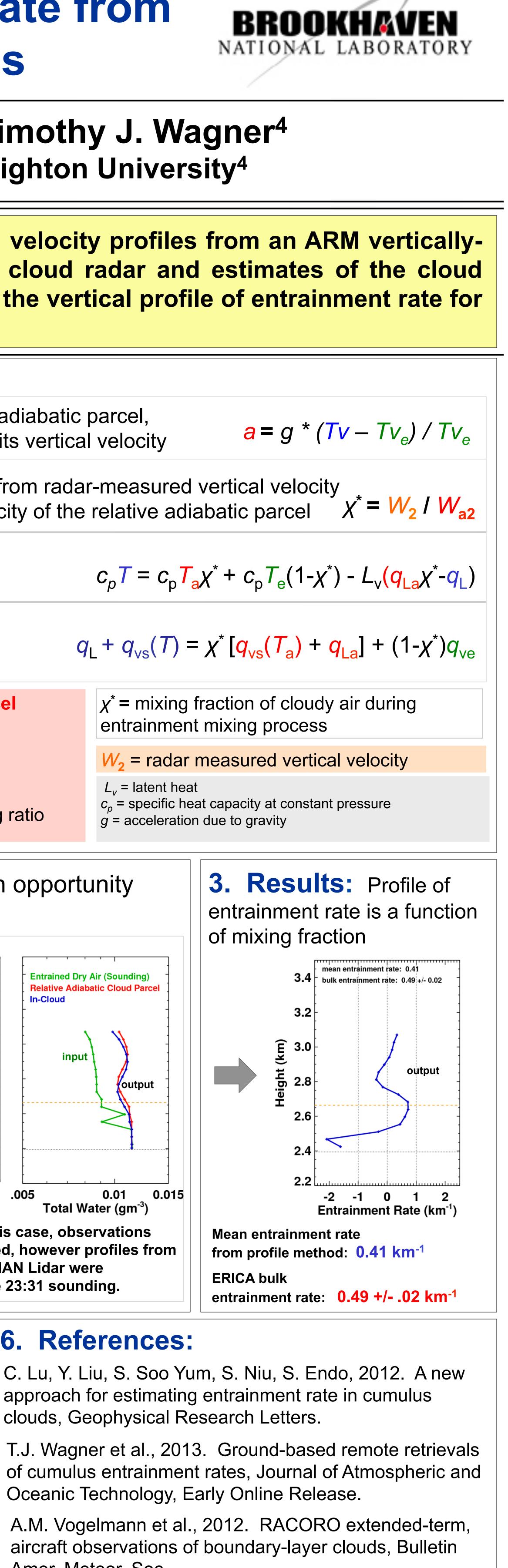
In-Cloud



Estimates of the vertical profile of entrainment rate from millimeter-wavelength radar observations

We introduce a technique using velocity profiles from an ARM verticallypointing, millimeter-wavelength cloud radar and estimates of the cloud adiabaticity in order to estimate the vertical profile of entrainment rate for shallow cumulus clouds.

2. Data: Shallow cumulus on June 22, 2009 at the Southern Great Plains, during the RACORO campaign, provided an opportunity to implement the method using ARM Millimeter-wavelength Cloud Radar (MMCR) and radiosonde observations.



Mixing Equations:

Acceleration of the relative adiabatic parcel, which is used to determine its vertical velocity

Mixing fraction determined from radar-measured vertical velocity and calculated vertical velocity of the relative adiabatic parcel

In-cloud temperature solved iteratively

Solve for in-cloud liquid water mixing ratio

Relative Adiabatic Cloud Parcel T_a = temperature a = acceleration W_a = vertical velocity q_{La} = liquid water mixing ratio

 $q_{vs}(T_a)$ = saturation vapor mixing ratio

entrainment mixing process

 L_{v} = latent heat g = acceleration due to gravity

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