The Eastern North Atlantic ARM site: Clouds, aerosols and their interactions in the remote marine boundary layer

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ENA Site Science



- macrophysical interactions across scales
- <u>Theme 1</u>. Acquire process-based understanding of cloud microphysical-
- <u>Theme 2</u>. Understand how microphysical-macrophysical interactions depend upon and influence the aerosol and meteorological environment
- <u>Theme 3</u>. Assess and improve process and climate model representations of clouds, aerosols and their interactions.

Cloud droplet concentration and CCN in the coupled and decoupled MBL

- A coupled marine boundary layer (MBL) is defined as one with liquid water potential temperature difference (surface- Sc cloud base) $\Delta \theta_1 < 0.5$ K and total water mixing ratio difference $\Delta q_t < 0.5$ g kg⁻¹ below the cloud base
- The marine boundary layer at Graciosa is decoupled more than 90% of the time (Rémillard et al. 2012)
- Coupled stratocumulus topped boundary layer (STBL) depths are shallow (typically < 1 km), while decoupled are thicker (typically > 1 km)
- Cloud droplet concentration is much more strongly correlated with surface CCN concentration under coupled conditions



 Schematic diagram showing the characteristics of coupled and decoupled boundary layer when stratocumulus cloud is present at the top of the MBL [from Dong et al. 2015]. The presence of a transition layer with jumps in moisture and potential temperature is characteristic of a decoupled

Constraining precipitation in process and large scale models

Using W-band Doppler spectra to confront and improve LES microphysics



► LES models produce too many large drizzle drops: Excessive negative skewness just below cloud base in LES model-derived Doppler spectra compared with W-band radar.



- Next Steps:
- Use aircraft DSD measurements (VOCALS, ACE-ENA) to forward model radar Doppler spectra
- Use 1D model framework with LES microphysics to improve DSD evolution



▲ Frequency with which inversions (solid line) and transition layers (dashed line) are found at Graciosa during CAP-MBL

► Cloud droplet concentration vs surface **CCN** concentration (0.2% supersaturation) for coupled (left) and decoupled (right) stratocumulus-topped boundary layers [from Dong et al. 2015]



- OBSERVATIONS GRW-WACR DHARMA260_nb35 — DHARMA130 nb35 - DHARMA130 nb60 ----- DHARMA065 nb60 – – SAMEX260
 - LES MODELS
- ◄ Simulated drizzle DSDs (esp. from DHARMA) produce spurious negative Doppler spectral skewness below cloud base compared with W-band data, indicating systematic errors in simulated DSD shape.

Improving drizzle occurrence statistics in the ECMWF IFS

Many global models overestimate the occurrence of light surface precipitation. Radar retrievals at ENA provide an observational estimate of the frequency of precipitation occurrence that are being used to constrain and improve the ECMWF model

	Precipitatio	n occurrence a t Operatior	t cloud base a nal Model	and the surface		
precip at cloud base				precip at surface		
				\frown		
		OBS N	MODEL			
Jul09	Oct09	Jan10	Apr10	Jul10	Oct10	
		SCM Co	ontrol			
		OBS N	IODEL			
Jul09	Oct09	Jan 10 SCM N	Apr10 New	Jul10	Octio	
		OBS N				
Jul09	Oct09	Jan10	Apr10	Jul10	Oct10	

 Monthly precipitation occurrence is shown for the CAP-MBL period (observations in red). Both the operational and SCM version of the ECMWF model (green) overestimate precipitation occurrence at the cloud base and the surface, indicating that too much drizzle is generated in-cloud, and not enough evaporates below cloud base.

Changes to the autoconversion, accretion and evaporation schemes lead to a better agreement with observations in a modified version of the SCM [Ahlgrimm and Forbes 2014].



Cloud Classification via Self-Organizing Map Algorithm

The ultimate goal is to objectively identify stratocumulus cloud types using WACR radar data

Low cloud map nodes identified as stratocumulus



- CAP-MBL data were divided into 1-hour-long segments and introduced into a modified unsupervised clustering algorithm known as a Self-Organizing Map (SOM).
- The second round of SOM (shown at left) produced refined map nodes and successful in differentiating stratocumulus from other cloud modes.
- SOM is a possible mechanism for automated cloud classification at ENA.



Jan Feb Mar Apr May Jun Jul Aug Sep Oct Nov Dec

▲ Seasonality of SOM-classified stratocumulus frequency showing clear preference for summer

Aerosol and Cloud Experiments in Eastern North Atlantic (ACE-ENA) 40 km X-band Scanning radar range

Science Themes

- Budget of MBL CCN and its Seasonal Variation
- Effects of Aerosol on Clouds and Precipitation
- Cloud Microphysical and Macrophysical
- Structures, and Entrainment Mixing
- Advancing Retrievals of Turbulence, Cloud, Drizzle • Model Evaluation and Processes Studies



PI: Jian Wang (BNL) DOE G-1 Aircraft 1st IOP (summer): 80 flight hours, June 15-July 25, 2017 2nd IOP (winter): 80 flight hours, January 11 – February 20, 2018

G-1 will complement ENA sampling of mesoscale structure **▼**





Flight Plans

Lagrangian drift to characterize the vertical structure and mesoscale variation of thermodynamics, aerosol, cloud, and precipitation

Repeated spirals near the ENA site and straight and level runs at multiple altitudes, \sim 20-30 km in length Aerosol profiling in FT - vertical profile up to 5000 m to identify potential elevated aerosol layers in the FT, which will be sampled using horizontal legs.



30-60 km

• Diurnal distribution of subcloud drizzle layer thickness in stratocumulus identified using SOM and normalized with respect to the height of cloud top (MBL Depth).

• Interpretation: For example, 0.4 means that 40% of the total MBL depth is occupied by evaporating drizzle in the subcloud layer.

- Depth of MBL through which falling drizzle evaporates is seasonal and diurna
- Largest diurnal difference is observed in stratocumulus during winter.



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