🐨 SOA: Potential in Ambient Air at GoAmazon & Partitioning Quantification in a Chamber University

J.L. Jimenez^{1,2}, B.B. Palm^{1,2}, J. Krechmer^{1,2}, X. Liu^{1,2}, S.S. de Sá³, P. Campuzano-Jost^{1,2}, D.A. Day^{1,2}, W. Hu^{1,2}, D. Pagonis^{1,2}, P. Ziemann^{1,2}, Y. Liu³, R. Seco⁴, J.H. Park⁵, S. Sjostedt⁴, A. Guenther⁴, S. Kim⁴, of Colorado J. Brito⁷, F. Wurm⁷, P. Artaxo⁷, L. Yee⁸, G. Isaacman⁸, A. Goldstein⁸, R. Wernis⁸, R. Souza⁹, A.O. Manzi¹⁰, O. Vega¹¹, J. Tota¹², M.K. Newburn¹³, M.L. Alexander¹³, S. Springston¹⁴, S. Martin³, W.H. Brune¹⁵ Boulder ¹CIRES and ²Dept. of Chemistry, Univ. of Colorado, Boulder; ³Harvard University, Cambridge, MA; ⁴University of São Paulo, Brazil; ⁸University of California, Berkeley; ⁹University of a lifornia, Irvine; ⁵National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder, CO; ⁶Pacific Northwest National Center for Atmospheric Research, Boulder; ⁹University of California, Berkeley; ⁹University, California, Berkeley; ⁹University of California, Berkeley; ⁹University, Californi

the State of Amazonas, Manaus, Brazil; ¹²Universidade Federal do Oeste do Pará, Santarém, Brazil; ¹³EMSL, PNNL, Richland, WA;¹⁴ Brookhaven National Laboratory; ¹⁵Penn State Univ.

Conclusions

- OFR approx. reproduces and extends atmospheric oxidation
- Yields: standard addition to ambient are similar to chambers
- SOA formation correlates with ambient VOC concentrations Oxidation of ambient air produces x2-10 more SOA than **predicted** from traditional VOC precursors, suggesting another source such as unmeasured oxidation products of VOCs
- Multilinear regression suggests that **partially oxidized biogenic** gases were the largest contributors during both seasons
- Biomass burning precursors more important in the dry season

Motivation: SOA Sources, Aging, and Lifetime • Secondary Organic Aerosols (SOA) have significant effects on climate, air quality, and human health, but its sources, formation and aging in the atmosphere is poorly understood

<u>Case Study: Covariance of SOA Formation with Ambient VOCs</u>



Condensable Gases Partitioning vs. Seed Surface Area



Conclusions

- Quantification of fundamental SOA formation processes in a Teflon chamber is possible using fast-product-burst method
- Simple chemical system rapidly produces low-volatility gas products taken up by liquid OA seed and chamber walls
- Clear differences in kinetics vs. OA seed amounts allow us to separate and quantify the influence of wall and particle uptake
- We reproduce gas- and aerosol-phases w/ kinetic box model
- Aerosol mass accommodation coefficient (α): average ~0.6, near 1 for low-volatility species, decreases as c* increases
- The experimentally-constrained model framework can be used to derive a correction factor (Φ~1-4) for SOA mass yields due to vapor losses to walls as a function of species properties.
- Experiments underway or planned for solid and viscous seeds
- Model suggests no difference in yields between batch and continuous-flow chambers



Chamber studies:

o Cannot study ambient air in near real-time

 Have difficulty producing OA as oxidized as in the atmosphere • A new method is used to study OA formation and aging in real-time \rightarrow oxidation flow reactor

 Allows studying real-time variations of SOA-forming potential in the field



• SOA formation after approx. 3 days of OH aging follows the availability of biogenic precursor gases • Substantially higher SOA potential at night

SOA Formed from Ambient Air Compared to Amount Predicted from VOCs in GoAmazon IOP1 Wet Season IOP₂ Dry Season



• SOA from OH oxidation often >> than predicted from measured ambient precursors • Consistent with previous OH oxidation measurements (Palm et al. 2016) • Suggests that SOA is dominated by unmeasured primary and/or secondary S/IVOCs. The most likely explanation are S/IVOCs formed from O₂ or NO₂ oxidation of VOCs at night. The GECKO-A model will be used to evaluate this possibility.



Biogenic

1.07 µg m[∹] 61%

IOP1: 1.78 µg m⁻³

- Max. SOA in OFR

Sum of pred. SOA

---- NO_v---- Biogenic

-- BB tracers

IOP1



IOP2

(s)

- Wall loss timescale (no aerosol) measured at ~1000 s (consistent with Ziemann, Donahue results; Krechmer et al., 2016)
- Can directly measure the strong gas loss timescale dependence on aerosol surface area



- Equilibrium time scales for each gas-phase compound at different seed aerosol condensational sink values
- Combined gas loss timescales consistent with model predictions
- Compounds are colored according to SIMPOL-estimated c*

Motivation: GPP laboratory experiments

- Gas-particle partitioning (GPP) is a process that describes gasphase species uptake into particles depending on equilibrium
- Most aerosol models use gas-particle equilibrium partitioning theory as default treatment of gas-aerosol transfer, despite questions about its validity and applicability
- Remaining questions about how gas-phase wall losses affect SOA yields

<u>Measurement of α , the mass accommodation coefficient</u>



- This has been a remarkably difficult parameter to quantify, with some literature studies proposing or using α = 0.001 Average for all compounds $\alpha \sim 0.6$
- On the same range as values measured in several previous studies of liquid droplets or aerosols
- Points from are averages for each compound across multiple



 Measured decay of ambient VOCs and injected CO verifies that modeled OH exposure is accurate within approx. a factor of 2-3 (consistent w/ Li et al. 2010, Peng et al. 2015)

Method: Oxidation Flow Reactor • 13 L Potential Aerosol Mass (PAM) oxidation flow reactor (Kang et al., 2007; 2011; Lambe et al., 2011,; 2012; Li et al., 2013; 2015; Ortega et al., 2013; 2015; Palm et al., 2016, 2017; Peng et al., 2015; 2016; 2017) • Larger range of oxidant exposures than chambers with reduced wall effects • Oxidant concentrations in the reactor are elevated in

one of several ways:

OFR185: $O_{2} + hv (185nm) \rightarrow O_{2}$

decay: <u>kinetics</u>) and *c*^{*} (dependent on ending equilibrium 0.38 µg m⁻³ concentration: <u>thermodynamics</u>) Low volatility compounds (right panel above) have very well constrained α

Higher volatility compounds (left panel above) have $\alpha < 1$, but harder to quantify at < 0.1 from single experiments

- Combining results of multiple experiments reduces uncertainty
- Results are robust against uncertainties on C_{wall}

Experimental Method

experiments Error bars are standard error of the mean

Wall-loss SOA yield correction model



- Matrix providing an aerosol mass yield correction factor (ϕ) for SOA formation from single condensable gas w/ given c* and CS
- Factors for each VBS bin can be derived and applied
- Assumes that SVOC production is rapid
- Does not account for oligomerization, or heterogeneous particle-phase chemistry (but it can be used as a component of a more complex model that includes those processes)

Model set-up



• Aerosols and gases formed in the reactor are analyzed w/ HR-ToF-AMS, SMPS, and PTR-TOF-MS, CCN, etc. Photochemically-processed aerosols show SOA formation, oxidation, hygroscopicity, and cloud activation similar to the atmosphere (e.g., Kang et al., 2011; Lambe et al., 2011) • A model has been developed and verified to correct for losses of condensable gases in the OFR (Palm et al. 2016)





3.0 H

2.0

Biogenic

IOP2: 2.35 µg m⁻¹

Max. SOA in OFR

- Sum of pred. SOA

---- NO---- Biogenic

Biomass burning

BB tracers

1.61 µg m



— C₁₂H₂₃NO₆ meas

Model traces (α, c)

Best fit model:

- 1 0.26

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