

Relationships between Radar Reflectivity and Vertical Velocity Fluctuations in Convective Clouds



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Hannah C. Barnes*, Samson M. Hagos*, Zhe Feng*, Christopher R. Williams#, Alain Protat&

*Pacific Northwest National Laboratory, #CIRES/NOAA, &Australian Bureau of Meteorology

Introduction

- Dynamical and microphysical processes are known to be related
- Validation/development of cumulus and microphysical parameterizations require data describing the statistical characteristics of vertical velocity
 - Key areas of uncertainty include the relationship between vertical velocity variability and mass flux/microphysical processes.
- Radar reflectivity can be used as a proxy for the microphysical characteristics of convection

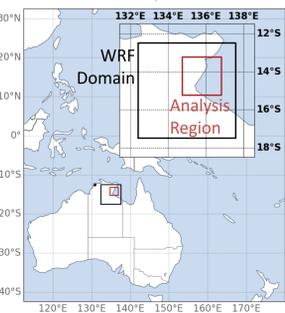
Objectives

- Explore whether vertical velocity variability can be characterized by substituting reflectivity into basic conservation equations
- Increase our ability to characterize vertical velocity and mass flux in convection

WRF Simulation

Time	18 -20 Jan 2006
Horizontal Resolution	500 m
Vertical Levels	100
Output Interval	1 min
Longwave / Shortwave Radiation Scheme	RRTMG
Microphysics Scheme	Thompson
Land Surface Scheme	Unified Noah
Surface Layer Scheme	Eta Similarity

WRF and Analysis Domains

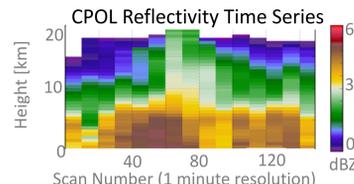
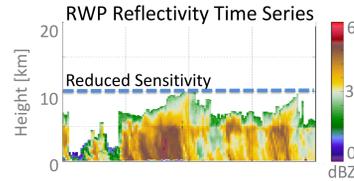
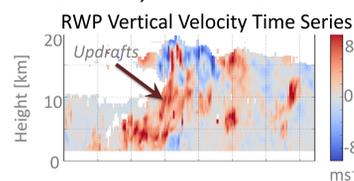


- Analysis conducted at 9 UTC 18 Jan 2006 for convective cores [Powell et al. 2016] in mixed phase regions (6-11 km)

Observational Data

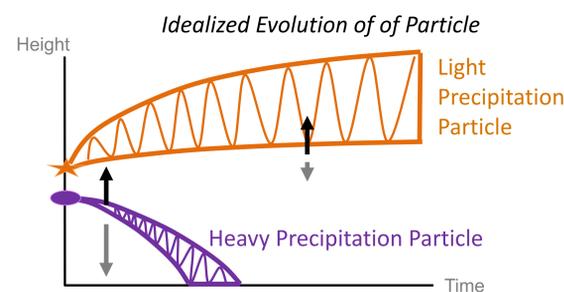
- Darwin, Australia
- 2005-06 & 2006-07 wet seasons
- Dual-frequency radar wind profiler (RWP)
 - Reflectivity
 - Vertical velocity [Williams 2012]
- CPOL Radar (Scanning C-band) [Keenan et al. 1998]
 - Reflectivity
 - Hydrometeor types [Keenan 2003; May and Keenan 2005]
- Data synchronized to 1 minute resolution over profiler

23 January 2006 1710-1930 UTC



Kinetic Theory (Diffusion + Drift)

Based on Fokker – Planck Equation

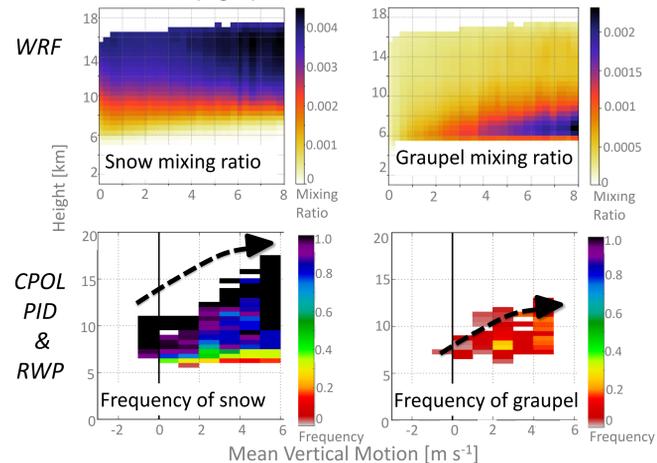


$$\frac{\partial P(x, t)}{\partial t} = -\frac{\partial [v(x)P(x, t)]}{\partial x} + D \frac{\partial^2 P(x, t)}{\partial x^2}$$

Change in distribution = Drift (Gravitational Balance) + Diffusion (Variability of Vertical Motion)

Drift

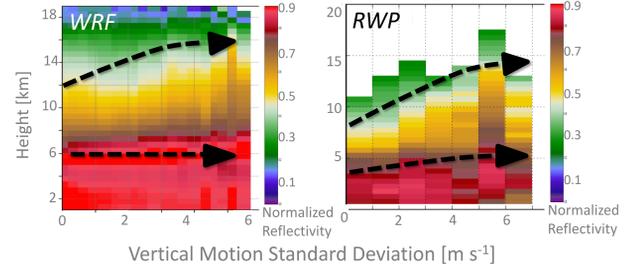
Microphysical Distribution in terms of Mean Vertical Velocity



- Higher vertical velocity magnitudes:
 - Suspends more mass
 - Lofts hydrometeors higher

Diffusion

Reflectivity in terms of Vertical Velocity Variability



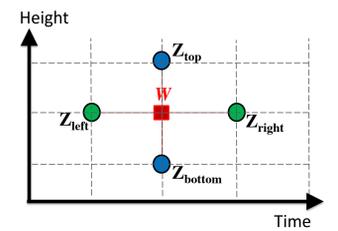
- Each profile normalized by maximum reflectivity
- Large vertical velocity variability associated with broad reflectivity distribution

Estimating Vertical Velocity

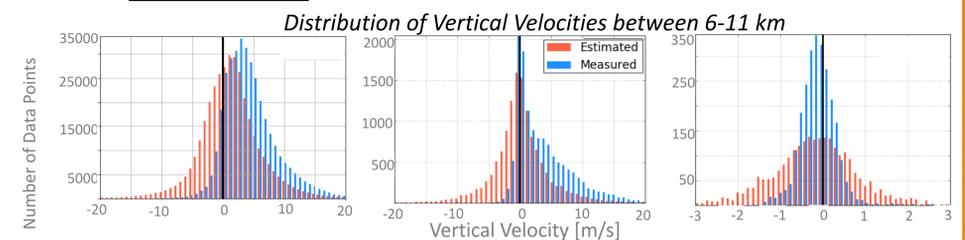
- Based on conservation of moisture and mass
- Reflectivity is used as a proxy for moisture and mass
- WRF based estimates assume no sources/sinks
- Profiler based estimates assume no sources/sinks and no horizontal advection
- Estimates exclude small vertical reflectivity derivatives

$$\frac{\partial Z_e}{\partial t} + \vec{U} \frac{\partial Z_e}{\partial \vec{x}} + W \frac{\partial Z_e}{\partial z} = \text{Source}$$

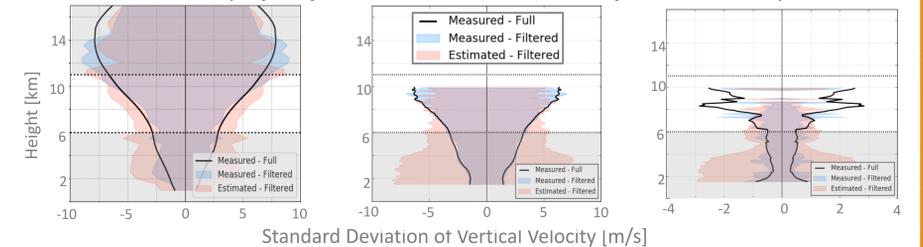
$$W = -\frac{\frac{\partial Z_e}{\partial t} + \vec{U} \frac{\partial Z_e}{\partial \vec{x}}}{\frac{\partial Z_e}{\partial z}}$$



WRF - Convection, RWP - Convection, RWP - Stratiform



Distribution of Vertical Velocities between 6-11 km



- The distribution of reflectivity can be used to describe a portion of the vertical velocity variability

Conclusions

- The statistical relationship between vertical velocity and hydrometeor distributions may be understood in terms of a kinetic theory based on:
 - Drift – Magnitude of vertical velocity determines the height and magnitude of the mass distribution
 - Diffusion – Vertical velocity variability influences the breadth of the hydrometeor distribution
- A portion of the vertical velocity variability in the mixed layer can be estimated by combining the conservation of moisture with the conservation of mass
 - Reflectivity with high spatial-temporal resolution can be used as a proxy for moisture and mass

Future Work

Our results motivate new observation strategies for scanning radars in order to improve our understanding of dynamical – microphysical interactions