

The thermodynamic structure and evolution of the Mixed Layer (ML) must be accurately represented in numerical models, as errors can lead to significant biases in many atmospheric processes, including the radiative fluxes, cloud properties and processes, precipitation processes, aerosol and chemical processes, and dispersion. We explore the ability of the SGP remote sensing instruments to capture the diurnal behavior of the ML and determine the ML height. The instruments and measurements include:

- Doppler lidar (DL) measurements of wind velocity
- AERI (AERI) retrievals of temperature

We examine the measurements acquired during the Land-Atmosphere Feedback Experiment (LAFE) (August 2017).

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### Method 2 Automated technique to identify sharp gradients at the top of the ML

using various parameters: Aerosol backscatter or extinction (RL-355 nm, UW-HSRL-532 nm)

Aerosol depolarization (UW-HSRL) Water vapor mixing ratio (RL, WV-DIAL)



hour of day

**UW-HSRL Backscatter 532nm** 

2.5

3

LAFE MLH compariso

ML height (km)

# **Remote Sensing Measurements of the CBL Structure during LAFE**

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# Background

Raman lidar (RLID) measurements of water vapor mixing ratio, and temperature

NCAR Water Vapor DIAL (WV-DIAL) measurements of water vapor mixing ratio

University of Wisconsin High Spectral Resolution Lidar (HSRL) measurements of aerosol backscatter and aerosol depolarization

- Potential Temperature (RL, AERI)
  - techniques:
- 20170808 SGP/LAFE 5











# Method 3

ML heights were derived from potential temperature profiles derived from a combination of AERI+RL temperature profiles. AERI temperature profiles are used below about 1 km and RL temperature profiles are used above 1 km. This takes advantage of better AERI performance near the surface and the higher resolution RL profiles farther away from the surface. ML heights are derived using two

1) Heffter technique (Heffter, 1980; Della Monache et al., 2004) seeks the lowest height where the potential temperature lapse rate exceeds 5°K/km and the inversion strength exceeds 2°K. • 2) Liu-Liang technique (Liu and Liang 2010) uses potential temperature differences and gradients, somewhat similar to the Heffter technique, for convective and neutral layers. For stable layers, the ML is

### Method 4

The bulk Richardson number (Ri<sub>b</sub>) represents the ratio of thermally produced turbulence to mechanically produced turbulence. ML heights were derived at the height at where (Ri<sub>b</sub> > 0.5). Ri<sub>b</sub> is computed using

## Findings

- Obtaining ML height from aerosol and water vapor gradients worked well during morning through late afternoon but did not work well at night due to the presence of elevated aerosol and water vapor layers. This is a common situation when trying to use these parameters to derive ML height
- ML height from RL+AERI potential temperature files provide a more realistic representation of the nighttime behavior of the ML height than the RL water vapor measurements or the HSRL aerosol measurements. During the day, these methods provide consistent results.
- At night, ML heights derived from RL+AERI potential temperature files using the Heffter and Liu-Liang algorithms compare well with ML heights derived from radiosonde potential temperature profiles in a similar manner. During the day, ML heights derived from RL+AERI potential temperature files using the Heffter and Liu-Liang methods are slightly lower, due possibly to some difficulty in calibrating the RL temperature profiles in cloudy conditions.
- ML heights derived from DL vertical velocity variance measurements during the day compare well with ML heights derived from aerosol, water vapor, and temperature gradients.
- During the day, ML heights derived from bulk Richardson number thresholds do not match well with ML heights derived from gradients in aerosols, water vapor, or potential temperature. At night, ML heights from bulk Richardson number compare better with those derived from potential temperature gradients.



