

# Aerosol effects on cloud field properties and precipitation of tropical convective clouds

Seoung-Soo Lee<sup>1,2</sup> and Graham Feingold<sup>1</sup>

<sup>1</sup> NOAA Earth System Research Laboratory

<sup>2</sup> CIRES, University of Colorado, Boulder



# Introduction

- The effect of aerosol on deep convective clouds is poorly understood
- Aerosol is thought to invigorate deep convective clouds (e.g., Koren et al. 2005; Rosenfeld et al. 2008; Li et al. 2012) but distinct changes in *total* precipitation have not been demonstrated
- There must have been microphysical compensation processes, considering the well-proven aerosol-induced autoconversion suppression (e.g, Lee et al., 2011)
- Compensation may affect the spatial/temporal distribution of precipitation

# Goals

- **Consider aerosol-cloud interactions for a *cloud system* over days**
- **Explore aerosol-induced changes in organization and precipitation in a cloud system comprising multiple clouds**

# Model Description

- **Goddard Cumulus Ensemble (GCE) model coupled with double-moment microphysics (Saleeby and Cotton, 2004)**
- **Interactive aerosol**

# Case

- **A mesoscale system of deep convective clouds (reaching the tropopause)**
- **Based on observations during TWP-ICE Darwin, Australia (Fridlind, 2009)**
- **Two-day simulations (most convective period)**
- **Conditions as prescribed by GCSS TWP-ICE case study**

# Simulations

**2-D domain: 256 x 20 km<sup>2</sup>**

**$\Delta x = 500$  m and  $\Delta z = 200$  m**

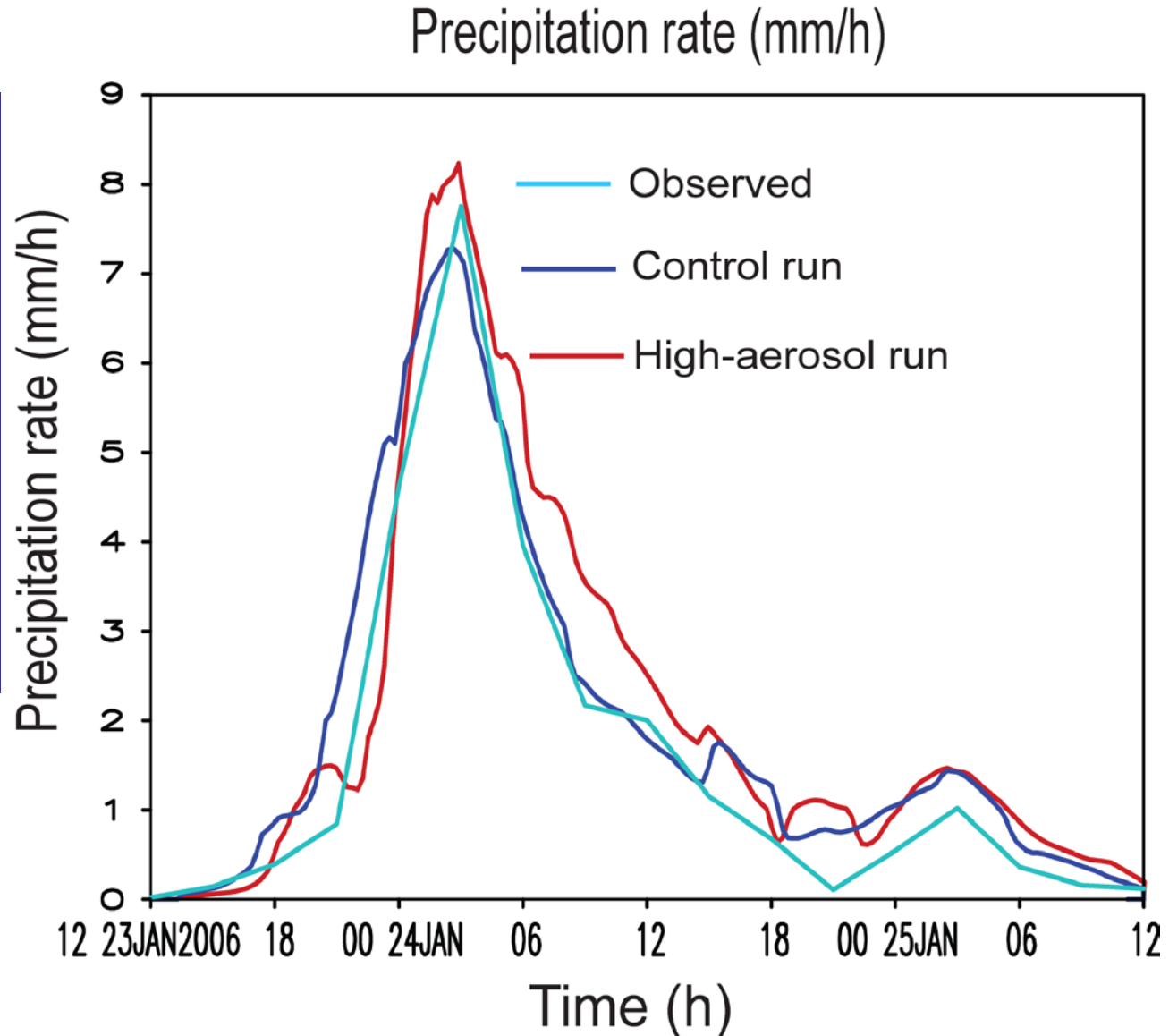
**PBL aerosol number concentration:**

- **Control run :  $\sim 400$  cm<sup>-3</sup> (Control)**
- **High-aerosol run:  $\sim 4000$  cm<sup>-3</sup> (High Aerosol)**

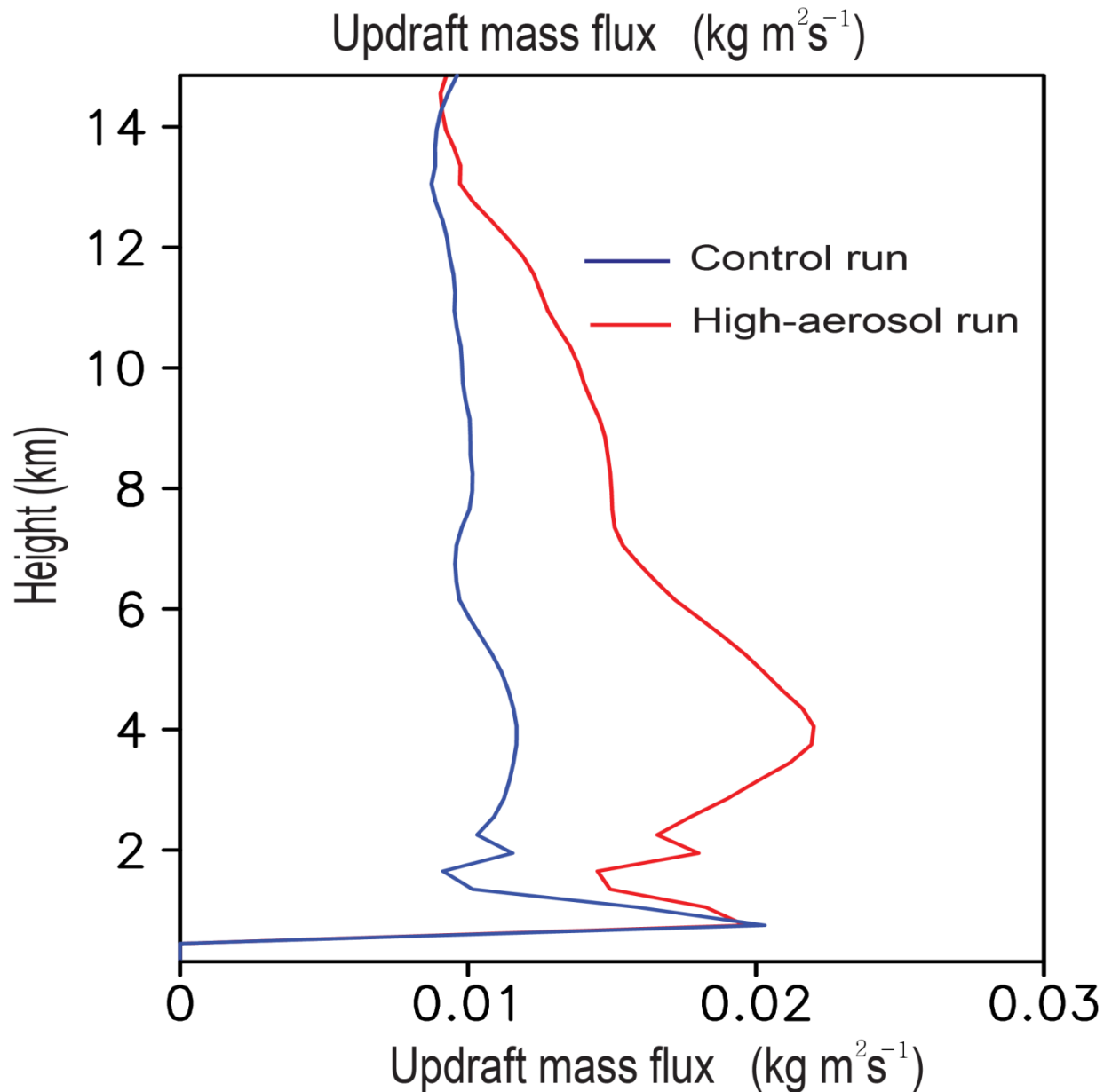
# Small Differences in Total Precipitation

**Cumulative  
Precipitation  
(mm)**

**Control: 88.6**  
**High-aerosol: 95.7**  
**(9% increase)**

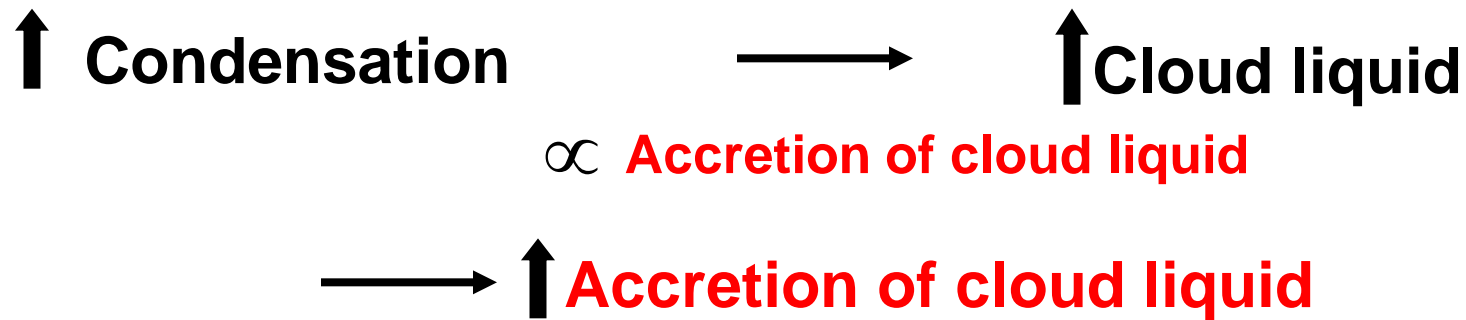


# Significant Increase in Updraft Mass Flux





# Cloud Response to an Increase in Aerosol



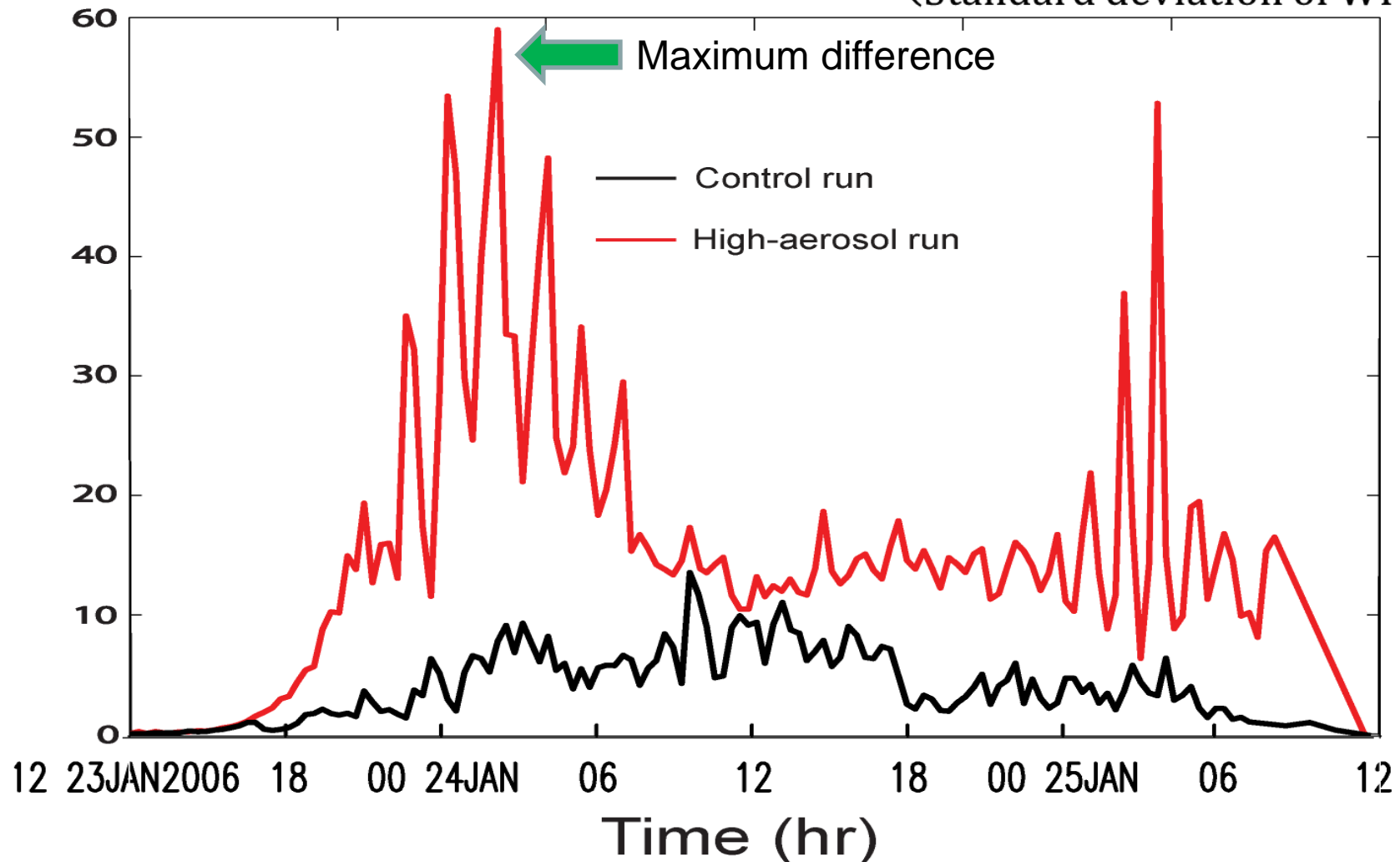
## ➤ Precipitation Budget (High - Control)

$\Delta$  **Precipitation**  $\approx$   $\Delta$  **Autoconversion +**  
**2.84 mm** **-21.08 mm**

**△ Accretion of cloud liquid  
23.91 mm**

# Distinct Differences in Cloud Field Properties

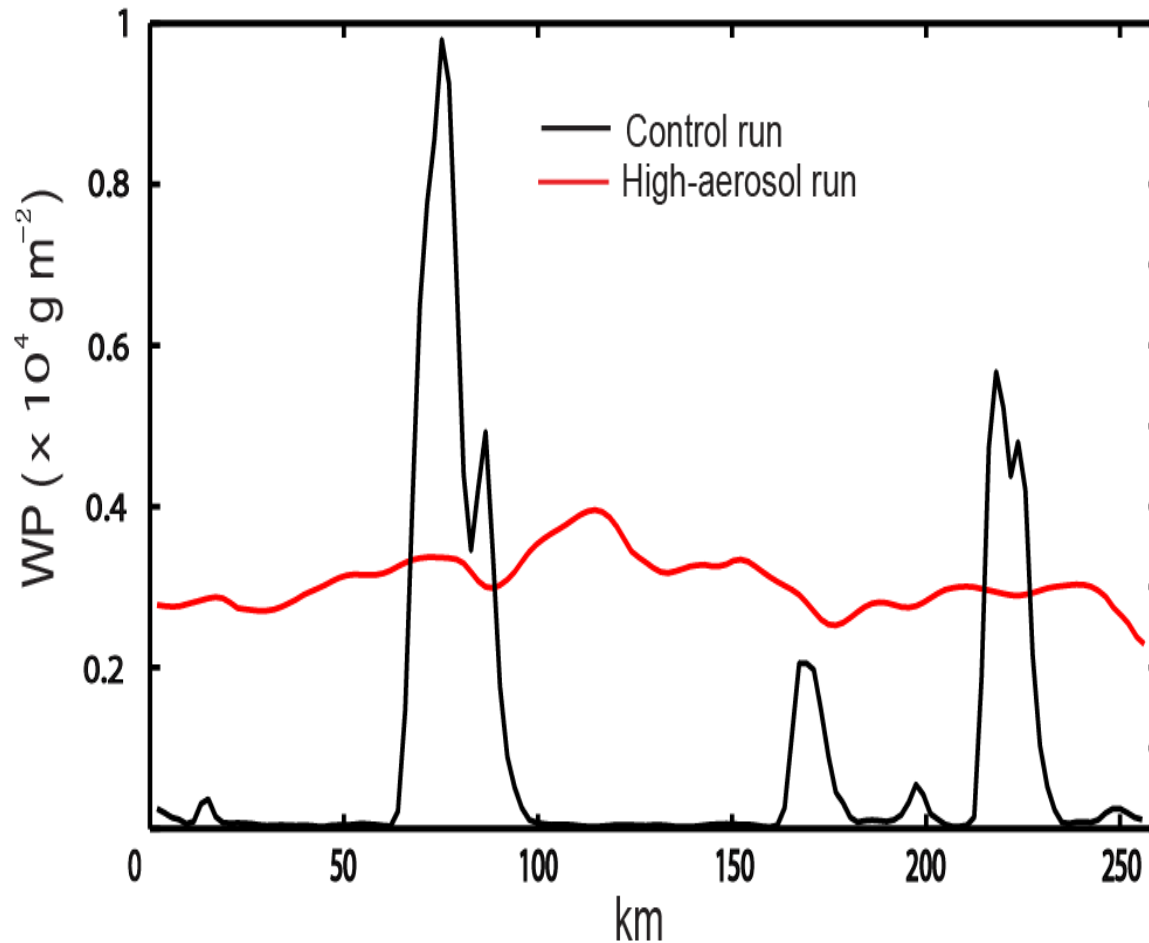
$$\text{WP (LWP+IWP) homogeneity} = \left( \frac{\text{Spatial mean of WP}}{\text{Standard deviation of WP}} \right)^2$$



# Distinct Differences in WP Spatial Distribution

At the time of the maximum difference in WP homogeneity

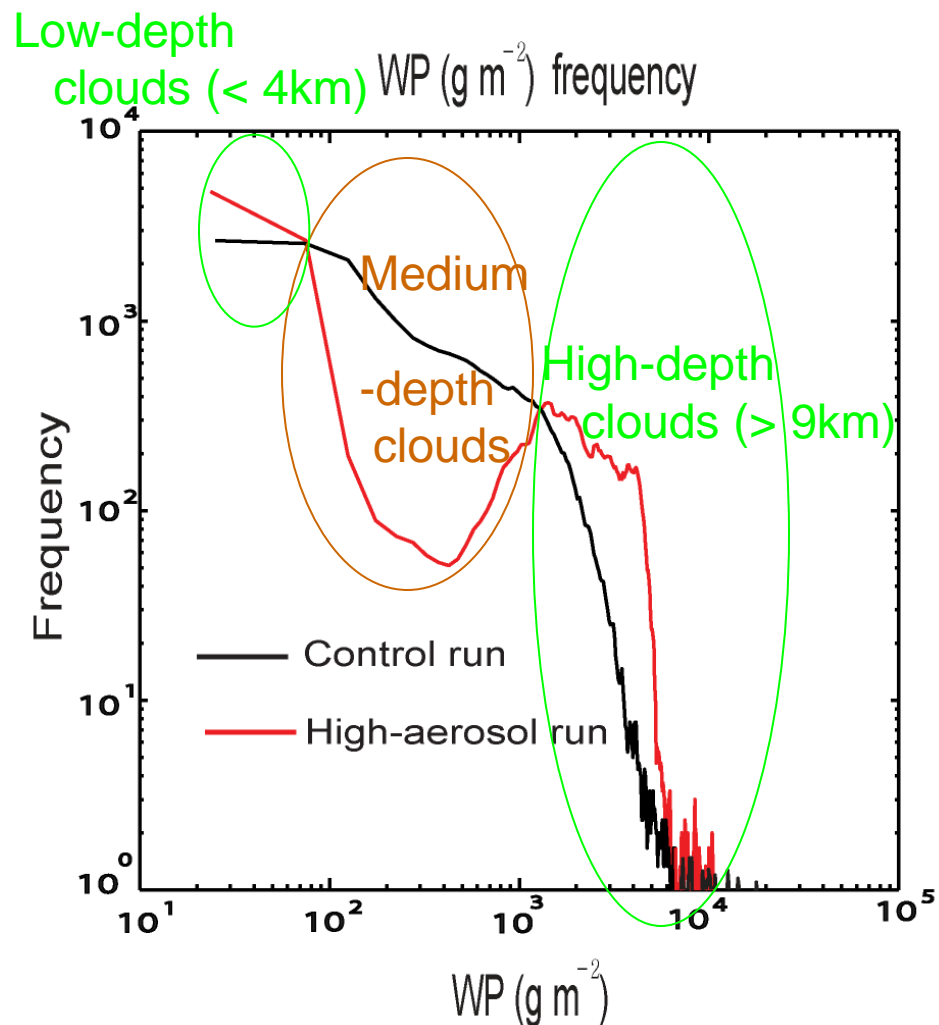
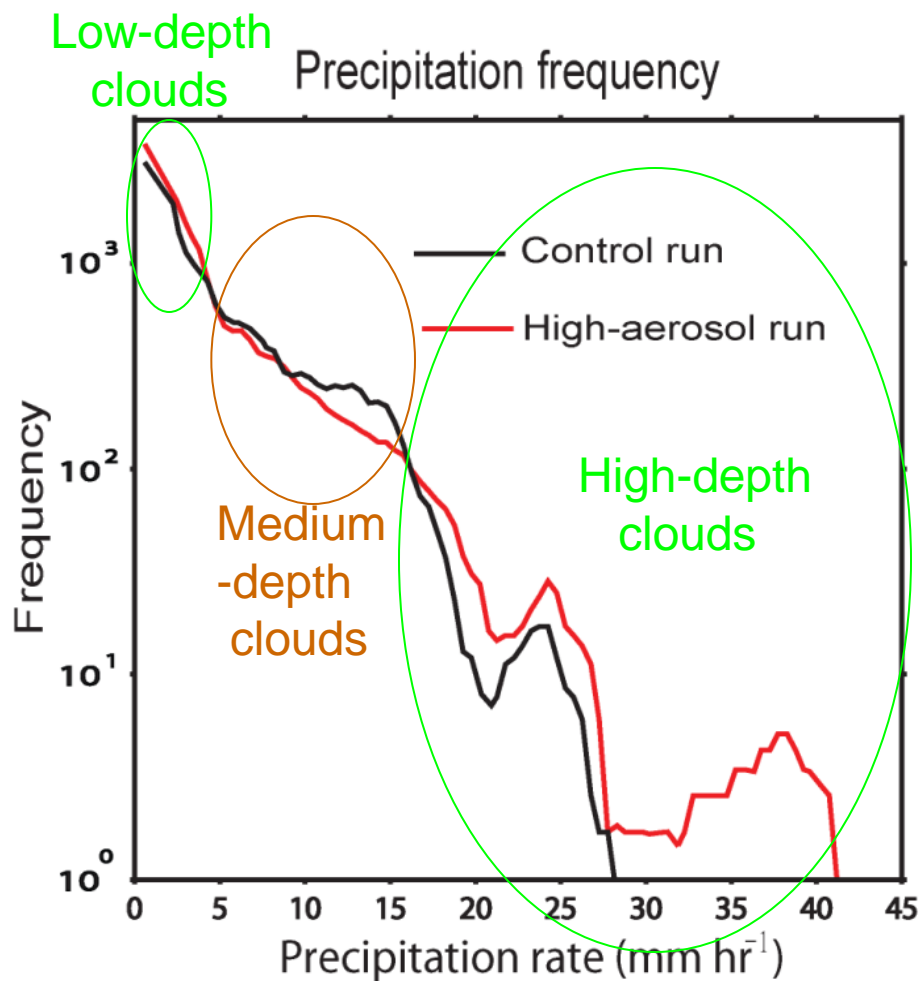
WP ( $\text{g m}^{-2}$ ) spatial distribution



**Total  
cloud population**

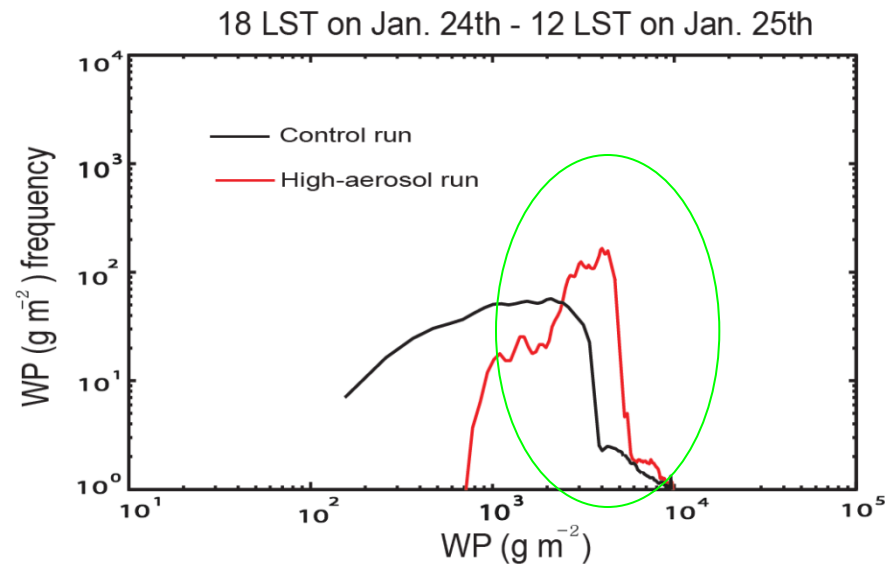
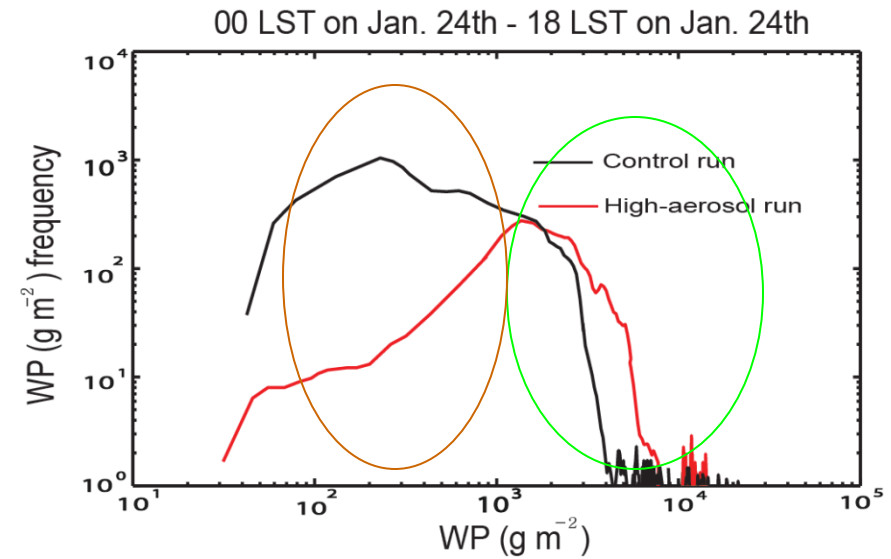
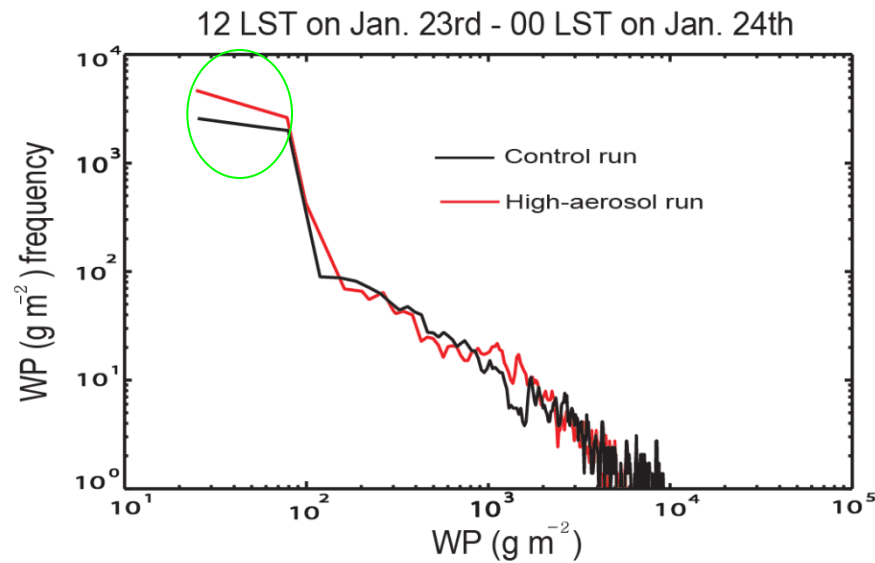
**Control run: 7  
High-aerosol run: 22**

# Correlation between Precipitation and WP Frequency



# Evolution of the Differences in WP Frequency

WP ( $\text{g m}^{-2}$ ) frequency



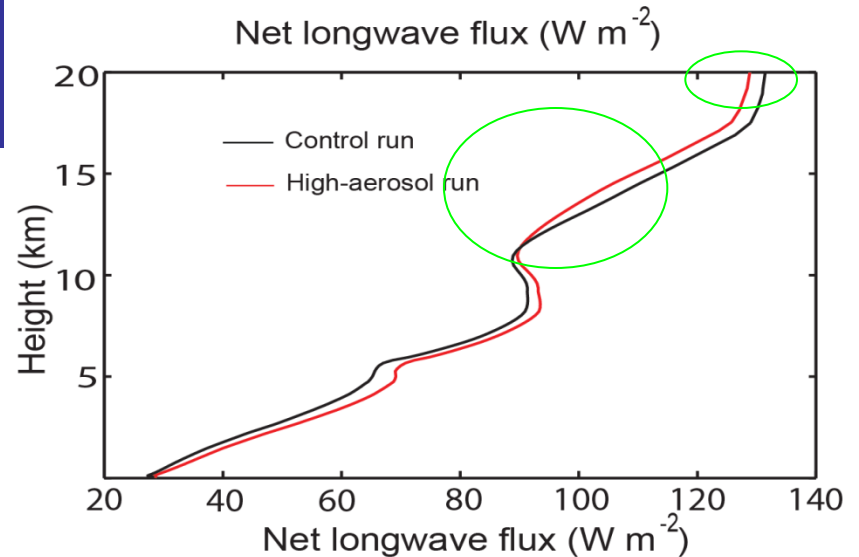
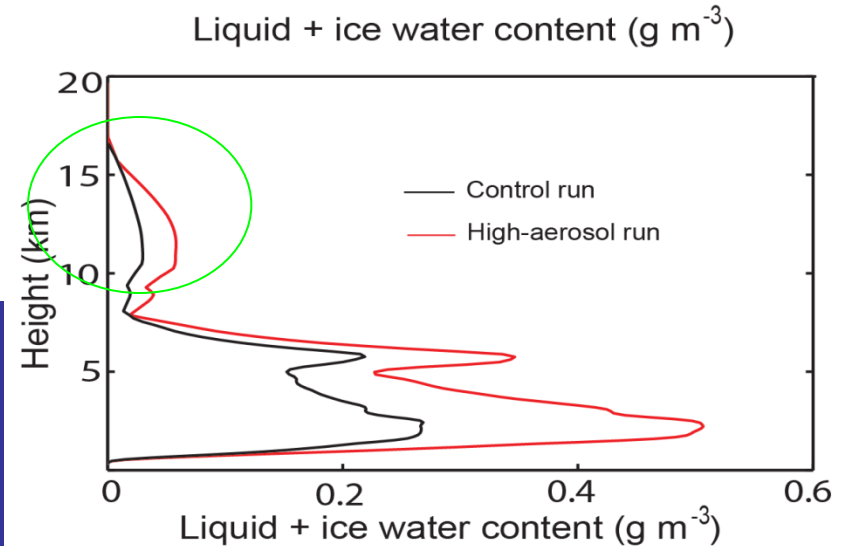
# Discussions and Conclusions

- For 2-day TWP-ICE simulations a 10-fold aerosol perturbation has a small effect on total precipitation (+9%)
- Substantial aerosol-induced enhancement in updrafts and cloud mass
- This enhancement accompanies significant changes in cloud field properties
  - Increase in WP homogeneity and in high and low WP; decrease in moderate WP
  - Increase in light and heavy rain; decrease in moderate rain
  - Increase in cloud population
- A meteorologically constrained cloud system achieves the approximately same amount of integrated precipitation

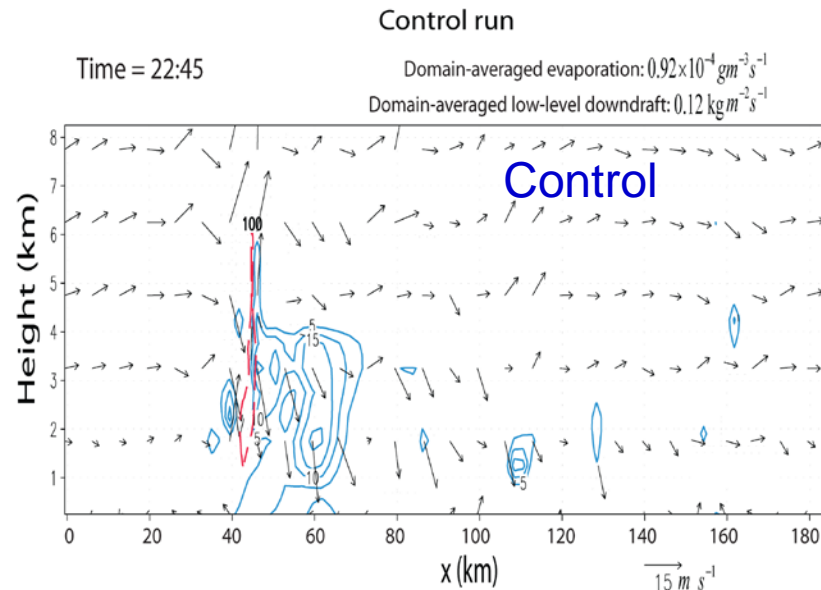
# Substantial Offset of Aerosol-Induced Change in Solar Radiative Fluxes by That in Longwave Fluxes

**Solar (longwave) radiative fluxes at the top ( $\text{W m}^{-2}$ )**

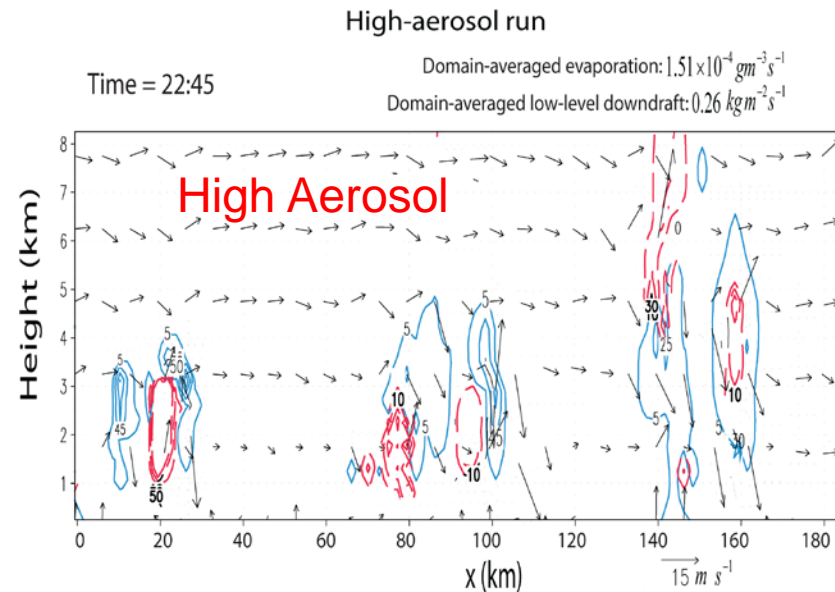
**Control run: -118.5 (133.2)**  
**High-aerosol run: -109.2 (129.5)**



# Distinct Differences in Cloud Field Properties

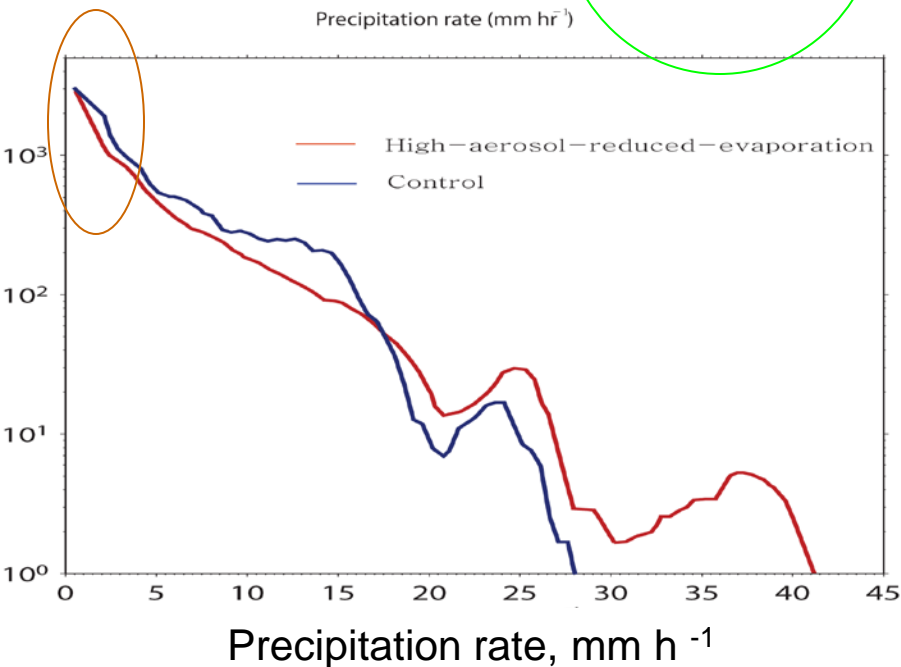
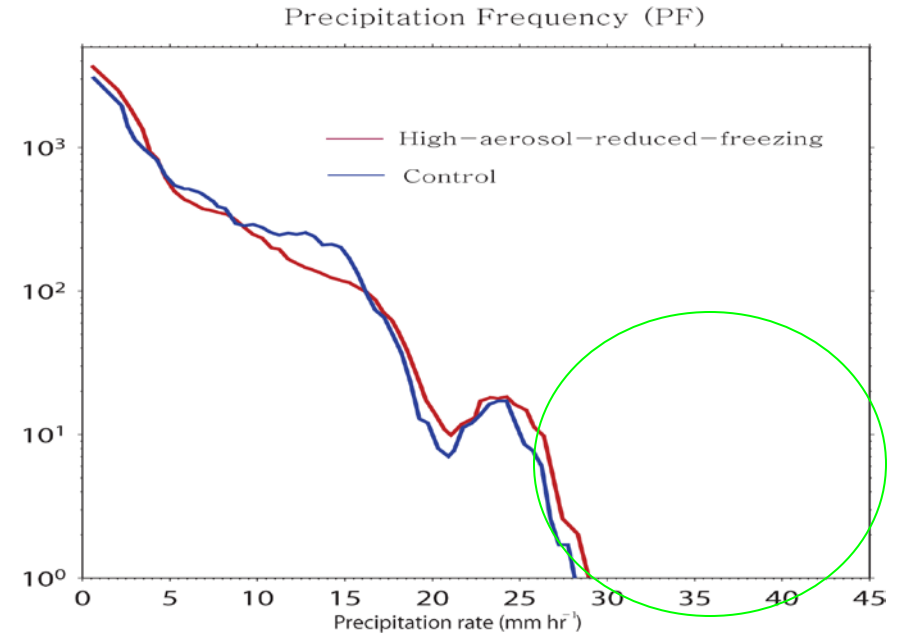
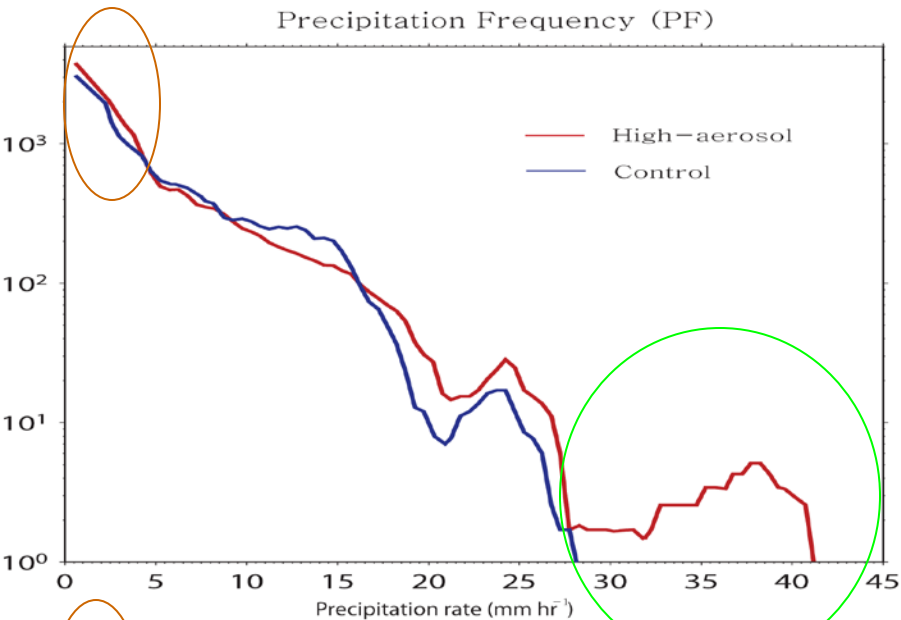


**Condensation/evaporation  
rates and wind flow**





# Precipitation Frequency



## Cumulative precip (mm) (cloud radius, km)

- Control: 88.6 (7.1 km)
- High-aerosol: 95.7 (5.6 km)
- High-aerosol-reduced-: 80.7 (7.5 km) evaporation
- High-aerosol-reduced-: 94.2 (5.3 km) freezing

# Discussions and Conclusions (I)

- For 2-day TWP-ICE simulations a 10-fold aerosol perturbation has a small effect on total precipitation (+9%)
- More significant changes to cloud system organization and the frequency distribution of rain rates
  - High aerosol simulations have larger number of small clouds (delayed autoconversion, more evaporation, stronger gustiness)
  - Aerosol-enhanced evaporation creates smaller clouds with lower rainrates
  - Aerosol-induced increase in freezing causes intermittent heavy precipitation, however, its impact on cloud-system organization is negligible
- We stress the importance of considering aerosol-precipitation interactions in cloud systems of long duration

- For 2-day TWP-ICE simulations a 10-fold aerosol perturbation has a small effect on total precipitation (+9%)
- Substantial aerosol-induced enhancement in updrafts and cloud mass
- This enhancement accompanies significant changes in cloud field properties
  - Increase in WP homogeneity and in high and low WP; decrease in moderate WP
  - Increase in light and heavy rain; decrease in moderate rain
  - Increase in cloud population
- A meteorologically constrained cloud system achieves the approximately same amount of integrated precipitation
- By-products of this system's effort for the achievement are substantial changes in updrafts, cloud mass and cloud field properties