

Quantifying Aerosol Influences on the Cloud Radiative Effect

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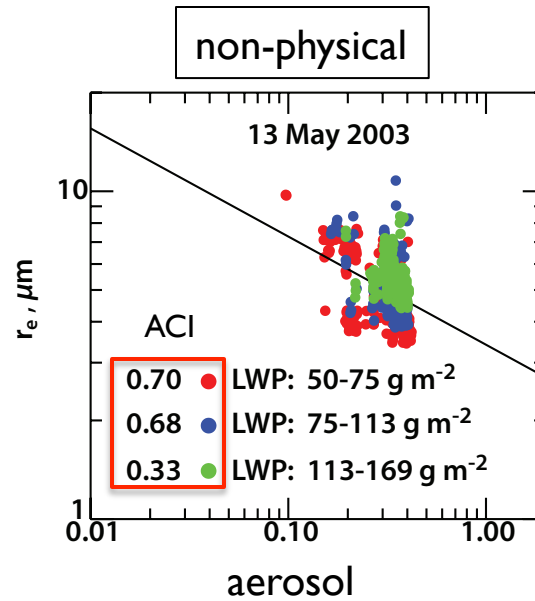
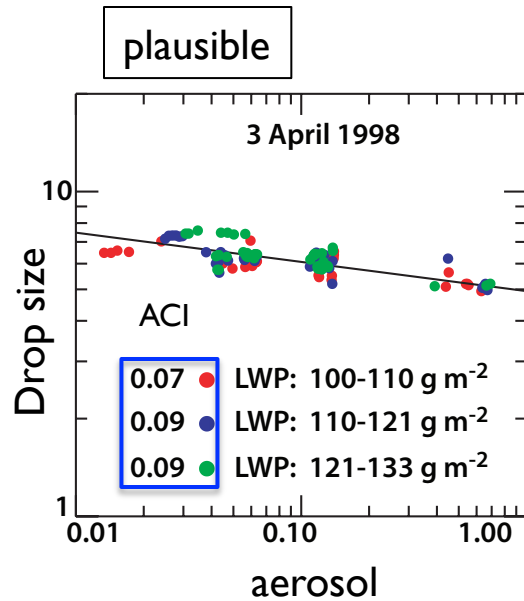
ASR, May 2016



Goal

Observational constraints on the Cloud Radiative Effect (CRE) / Aerosol Indirect Effect (AIE)

Approaches to quantifying Aerosol-Cloud Interactions (ACI)

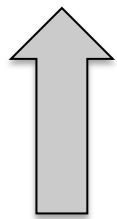


Quantifying these terms is a challenge

- Measurement errors, scale dependence, etc.

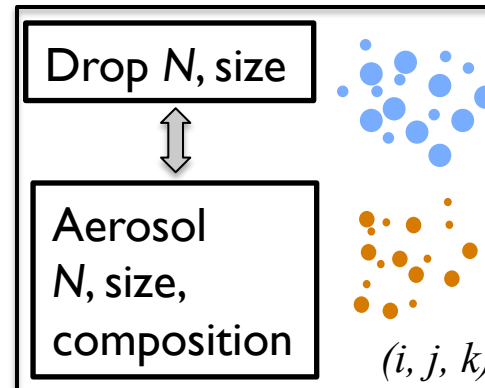
Small differences in ACI magnify to very large differences in CRE

→ measurements of microphysical ACI metrics do not constrain the CRE!



$$ACI = - \left. \frac{\partial r_e}{\partial N_a} \right|_L$$

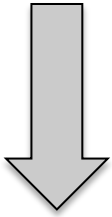
$$0 < ACI < 0.33$$



Bottom-up

Approaches to quantifying Aerosol-Cloud Interactions (ACI)

Top-down



$$A = A_c f_c + A_s(1 - f_c)$$



Cloud field Properties

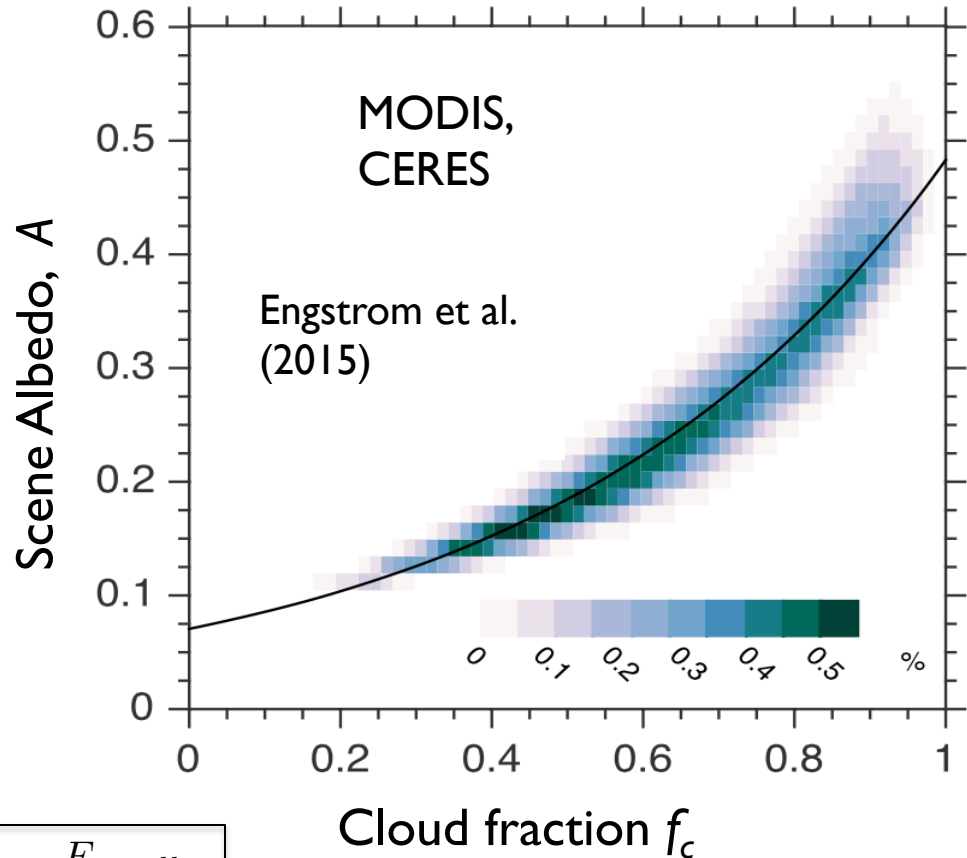
- Cloud fraction, f_c
- Liquid water path, L
- Optical depth, τ
- Cloud albedo, A_c
- Cloud depth, H

$$A_c = f(L, N)$$

$$\text{rCRE} = 1 - \frac{F_{sw,all}}{F_{sw,clr}}$$

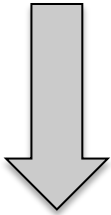
rCRE= relative cloud radiative effect

F_{sw} = downwelling shortwave flux



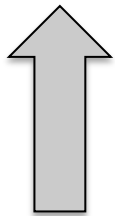
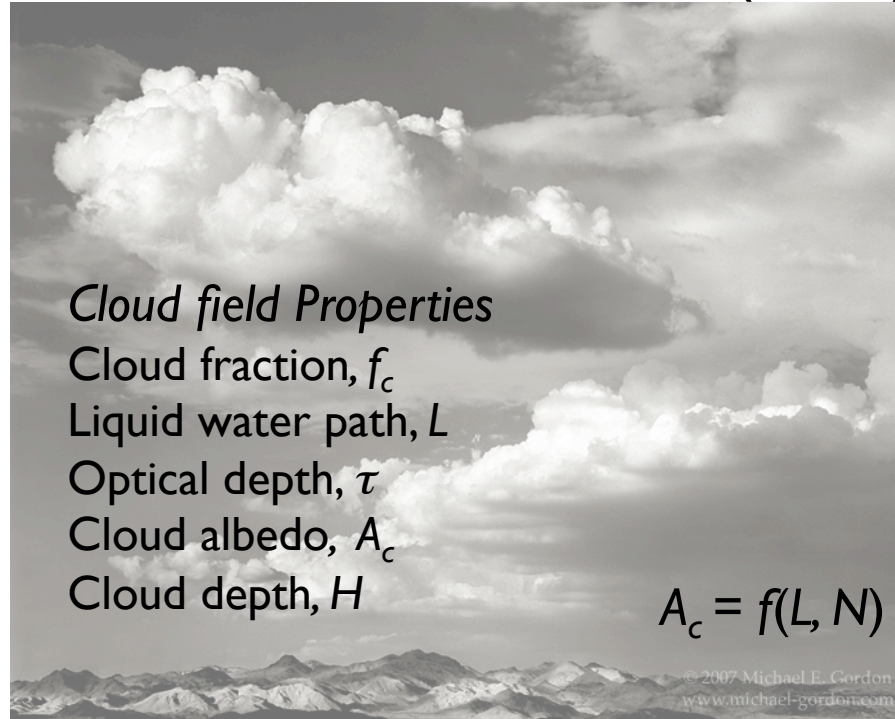
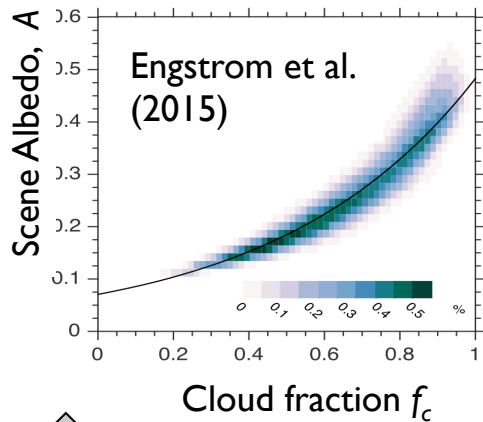
Approaches to quantifying Aerosol-Cloud Interactions (ACI)

Top-down



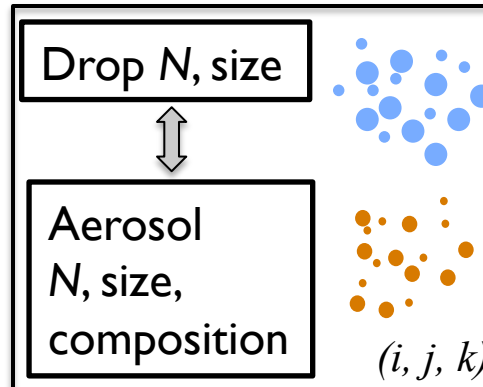
$$A = A_c f_c + A_s(1 - f_c)$$

MODIS, CERES



$$ACI = - \left. \frac{\partial r_e}{\partial N_a} \right|_L$$

$$0 < ACI < 0.33$$



Bottom-up

Approaches to quantifying ACI in cloud systems

Albedo

Emissions

$$\Delta R = R \frac{d \ln R}{d \ln \tau} \frac{d \ln \tau}{d \ln N_d} \frac{d \ln N_d}{d \ln CCN} \frac{d \ln CCN}{d \ln E} \Delta \ln E$$

Ghan et al. (PNAS 2016)

$$\frac{d \ln \tau}{d \ln N_d} = \frac{1}{3} \left[1 + 2 \frac{d \ln L}{d \ln N_d} + \frac{d \ln k}{d \ln N_d} + 3 \frac{d \ln H}{d \ln N_d} \right]$$

Ackerman et al. (2000)

$$\frac{d(\ln L)}{d(\ln N_d)} =$$

$$\frac{d(\ln k)}{d(\ln N_d)} =$$

etc..

**Caution!!
Uncertainties
compound**

Chen et al. (2014)
Goren and Rosenfeld (2014)

Feingold et al. (1997)
Liu and Daum (2000)

Demonstration of Difference Between Approaches

Example from LES

Two sets of simulations (~100 each)

differing only in co-variability of initial meteorology and aerosol conditions:

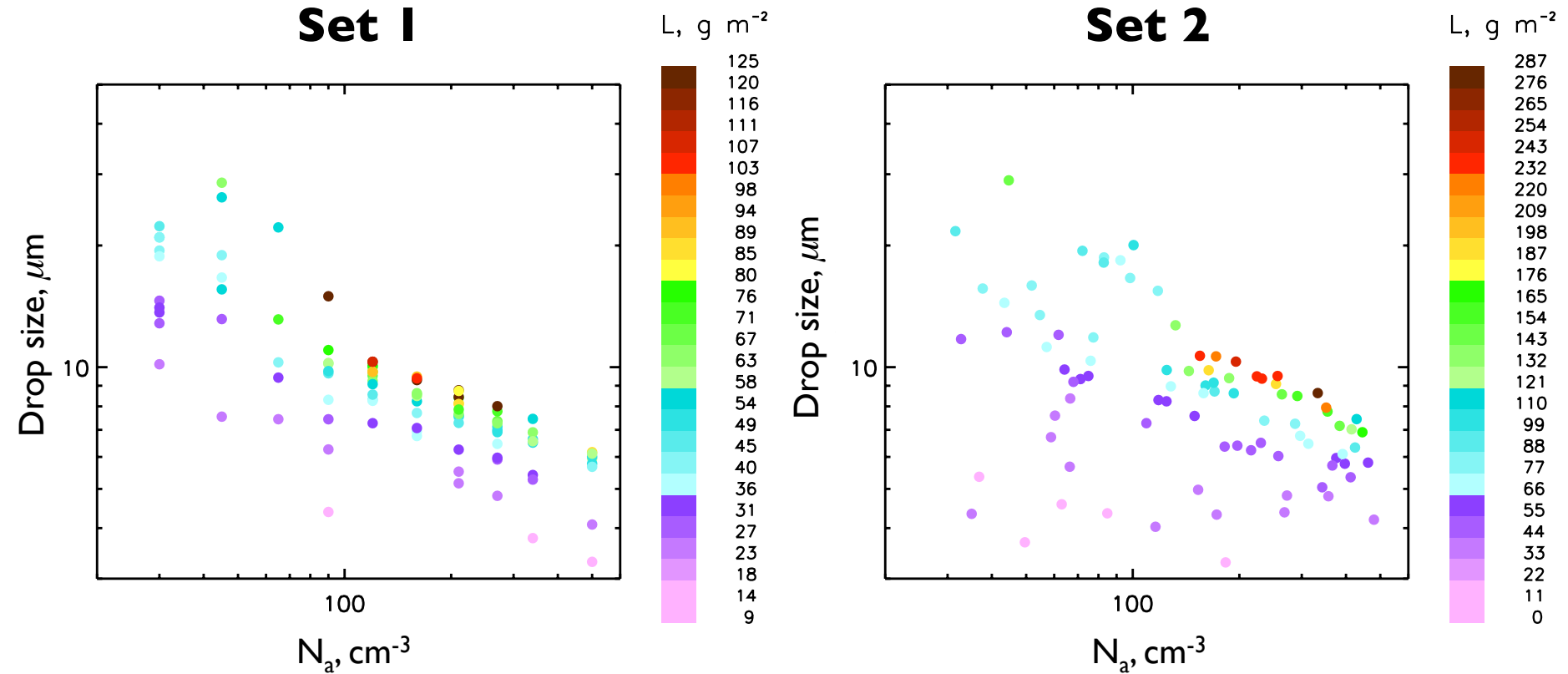
Set 1: regular grid spacings in 6-D parameter space; many runs vary N for fixed met conditions

Set 2: Latin Hypercube Sampling of 6-D parameter space; maximizes minimum distance between parameters for optimal coverage

Neither is realistic but differences will be illuminating

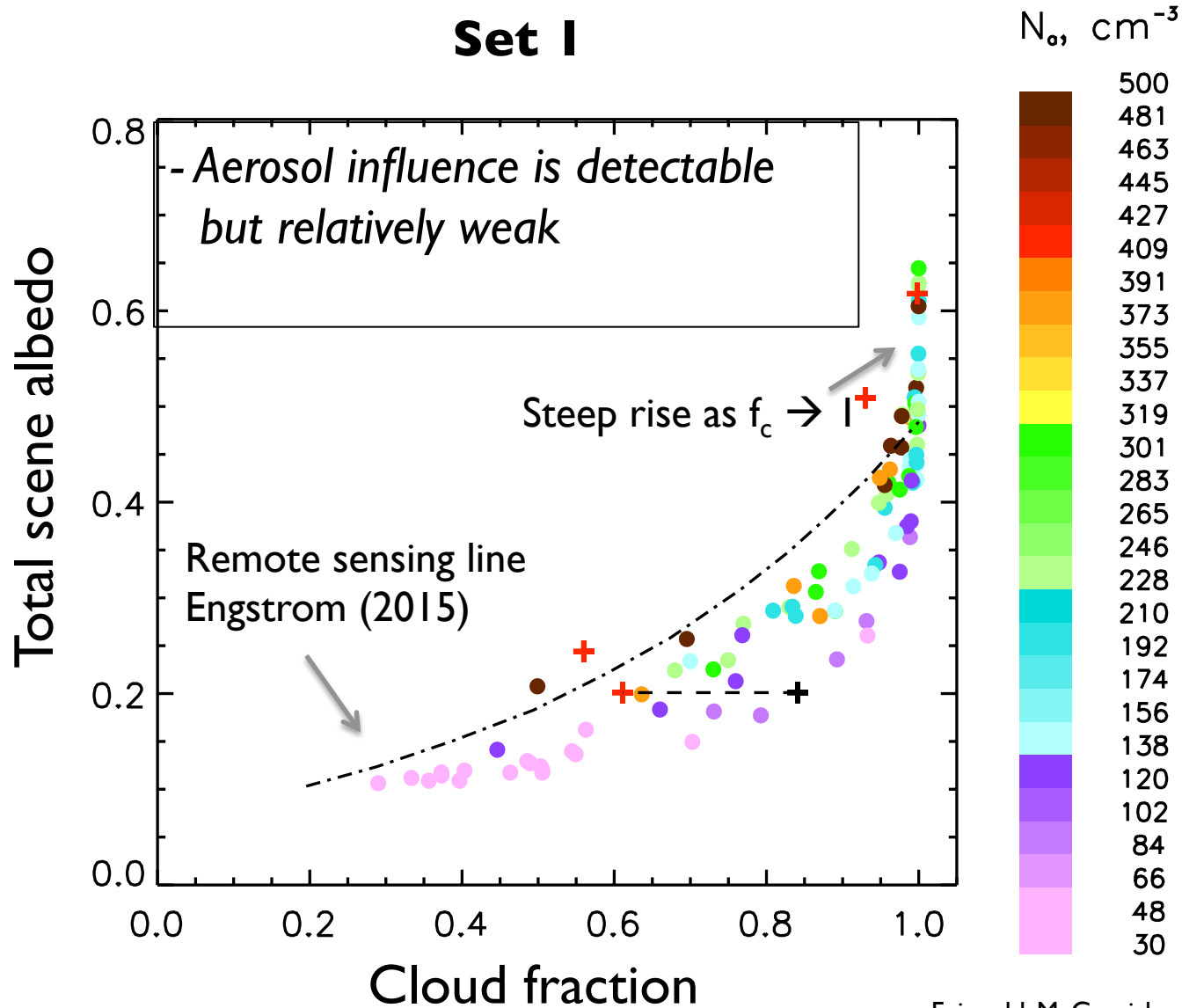
Microphysical Response (bottom-up)

Both Sets show expected reduction in r_e with increasing N_a



Albedo- f_c Response (top-down)

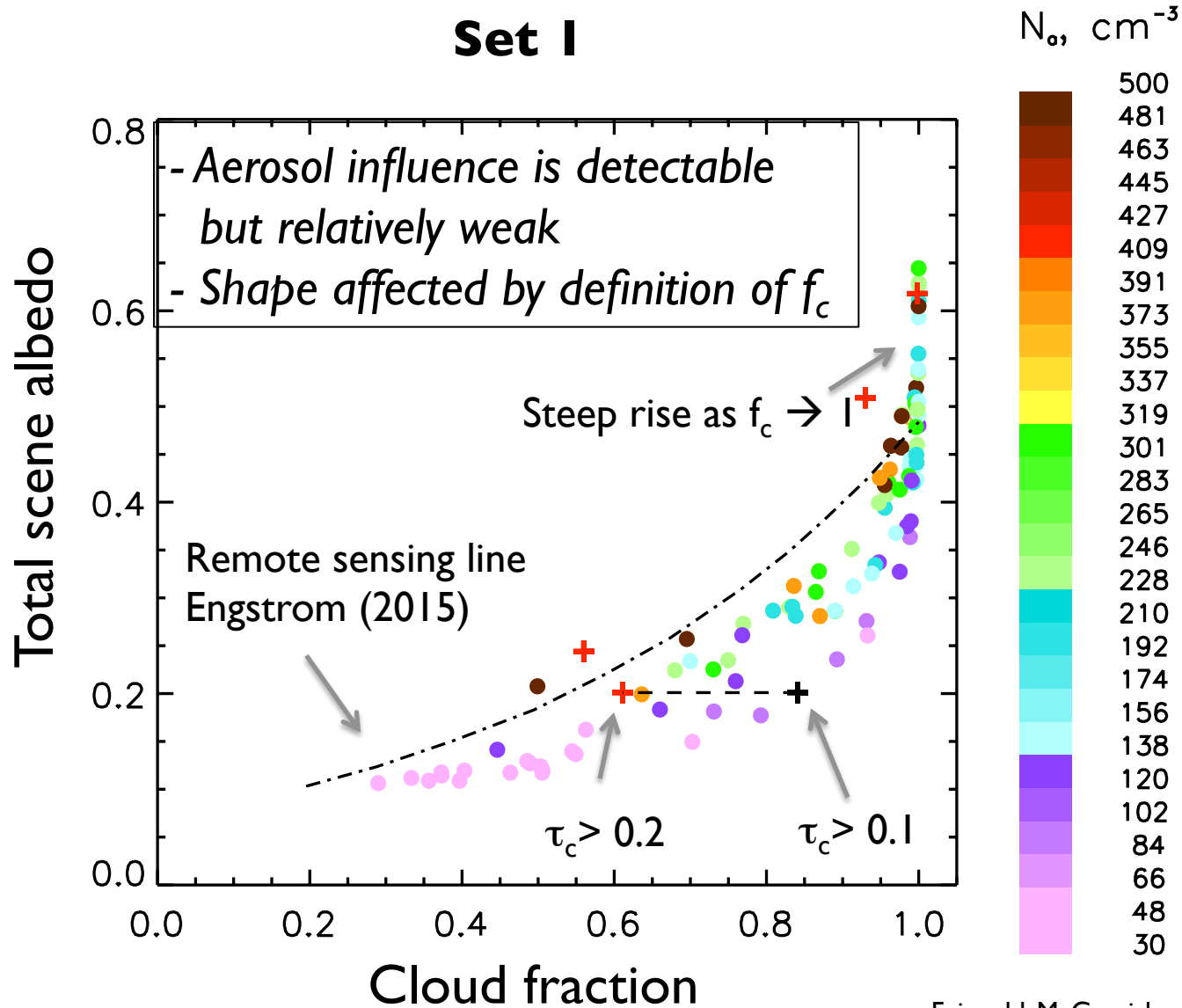
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Albedo- f_c Response (top-down)

$$A = A_c f_c + A_s(1 - f_c)$$

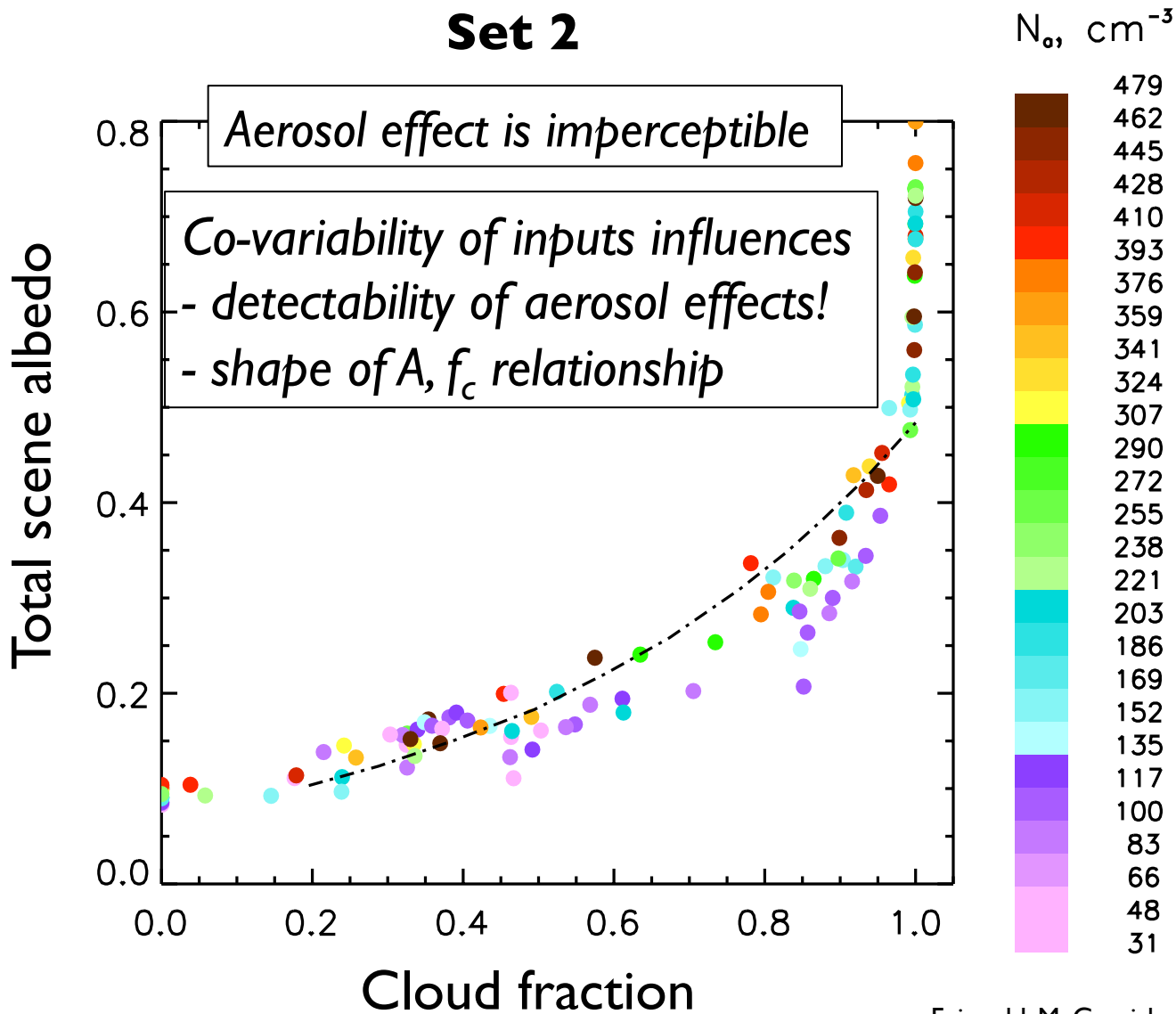
Set I



Albedo- f_c Response (top-down)

$$A = A_c f_c + A_s(1 - f_c)$$

Set 2



Data

Modeling

Routine,
long-term,
regime-based.

Natural
co-variability
of met. and N_a

$L, \tau, N_a,$
Radiation,
Surface Fluxes..

Albedo,
Cloud
fraction

CRE

Initial conditions
Met: NWP/VA/RGCM
Aerosol: observed

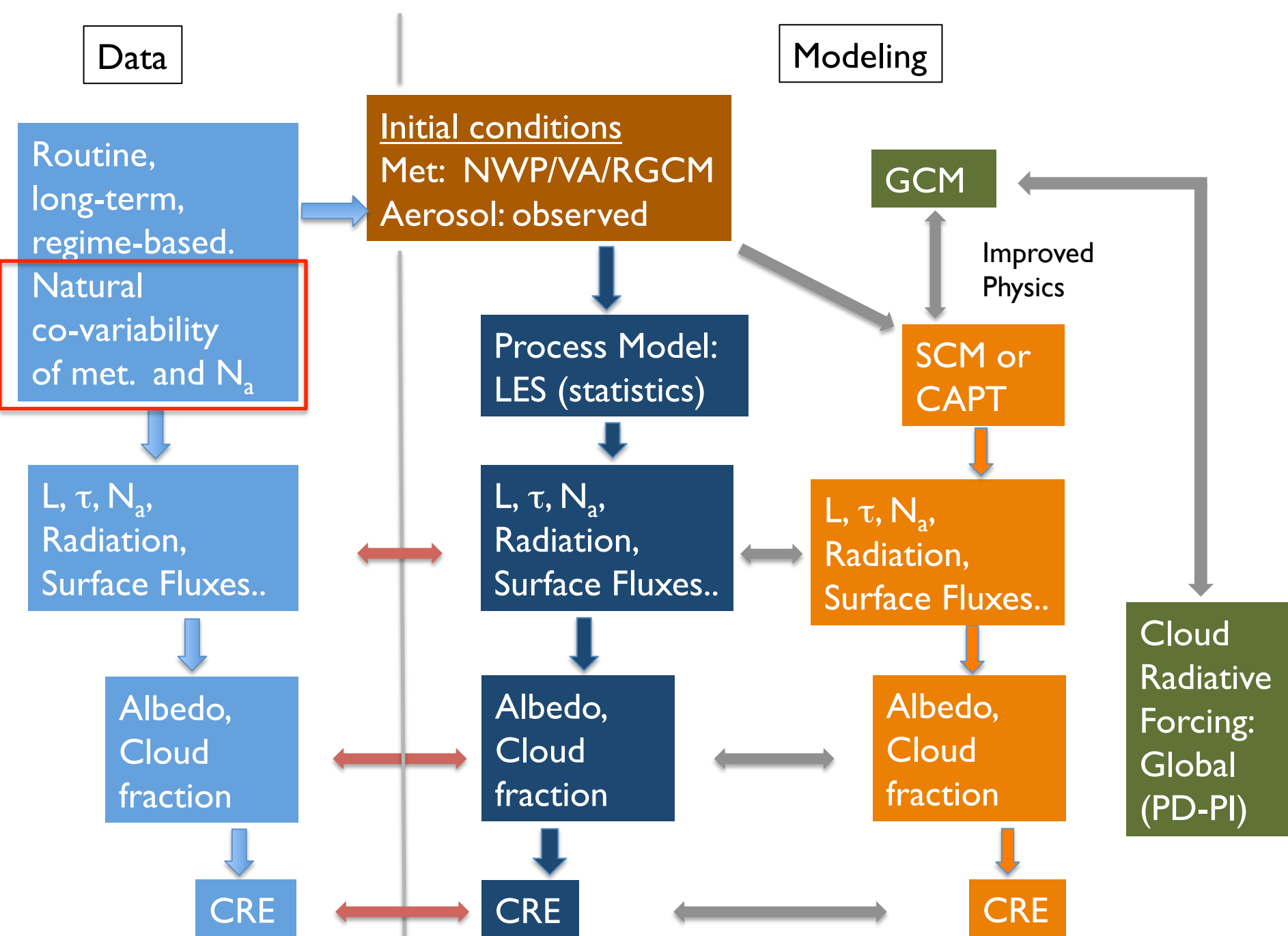
Process Model:
LES (statistics)

$L, \tau, N_a,$
Radiation,
Surface Fluxes..

Albedo,
Cloud
fraction

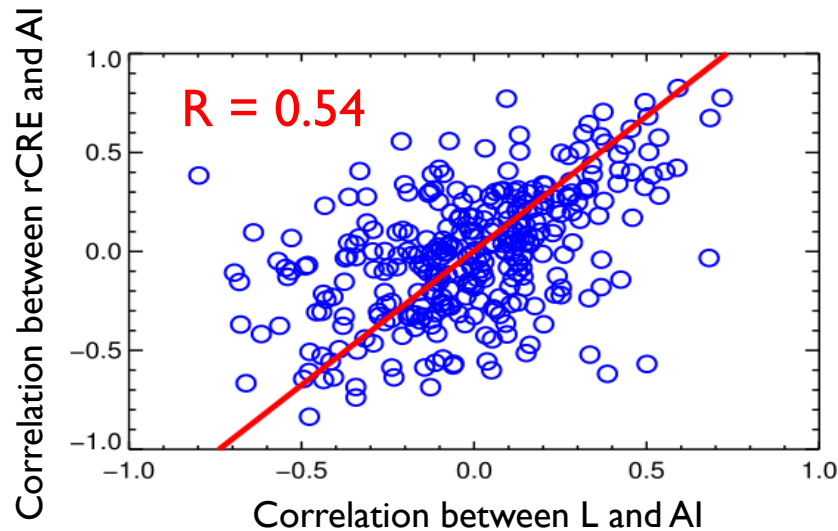
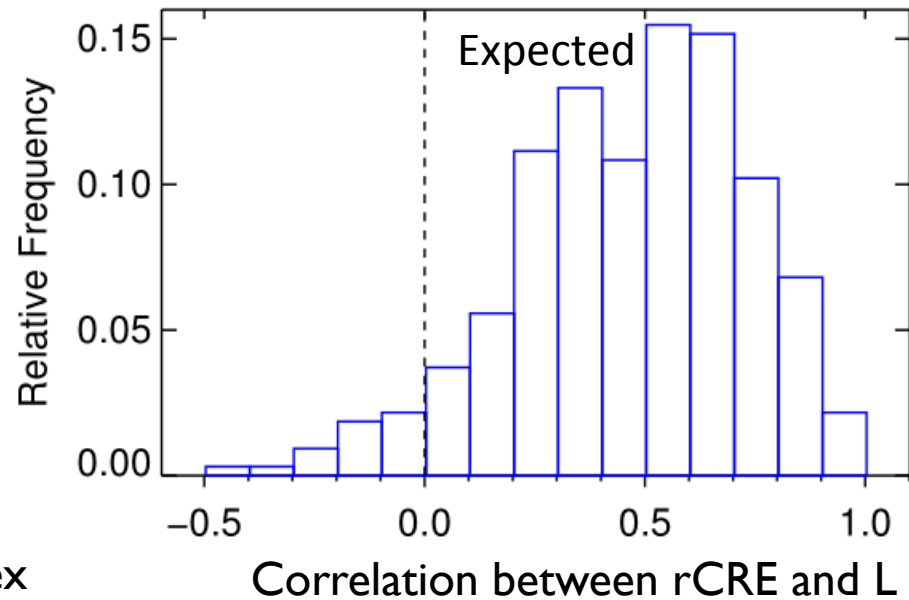
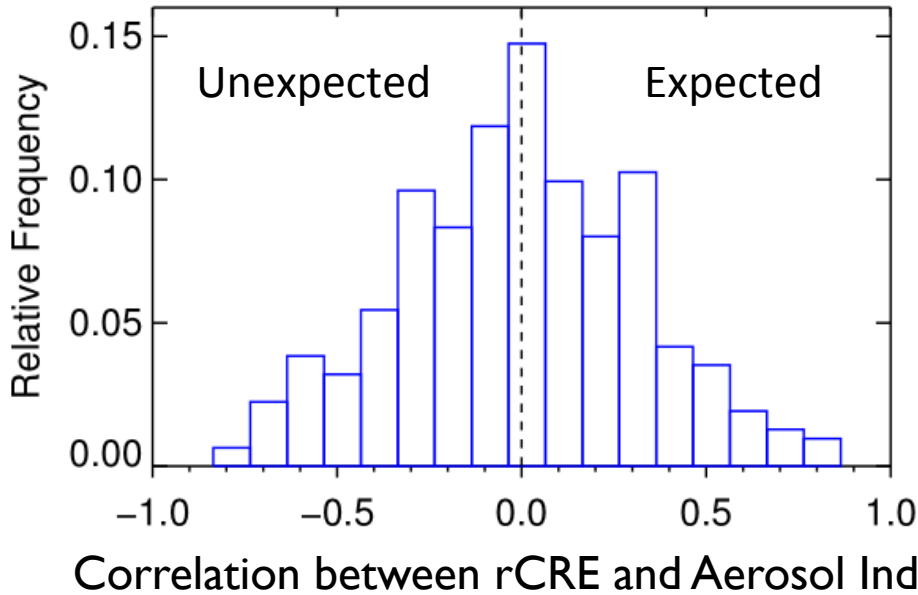
CRE





e.g., LASSO; Neggers 2012, Schalkwijk 2015; Cabauw

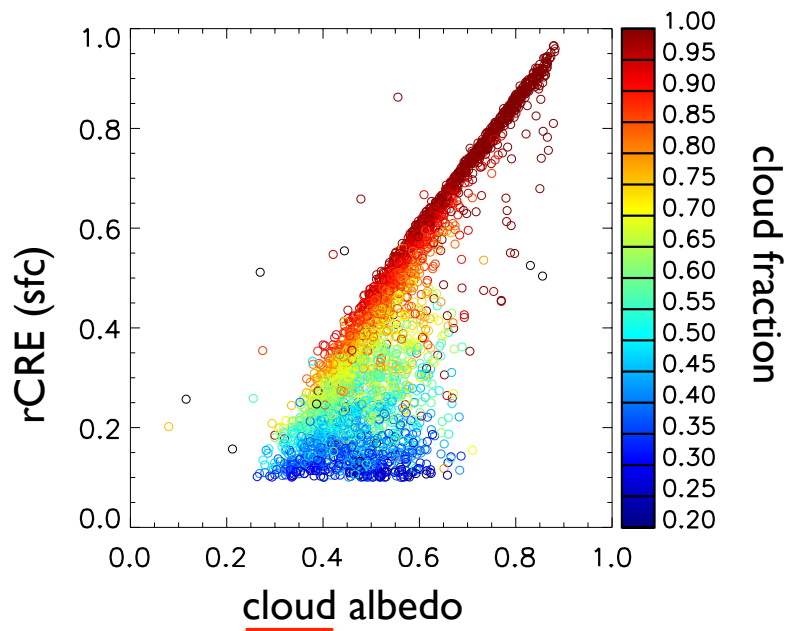
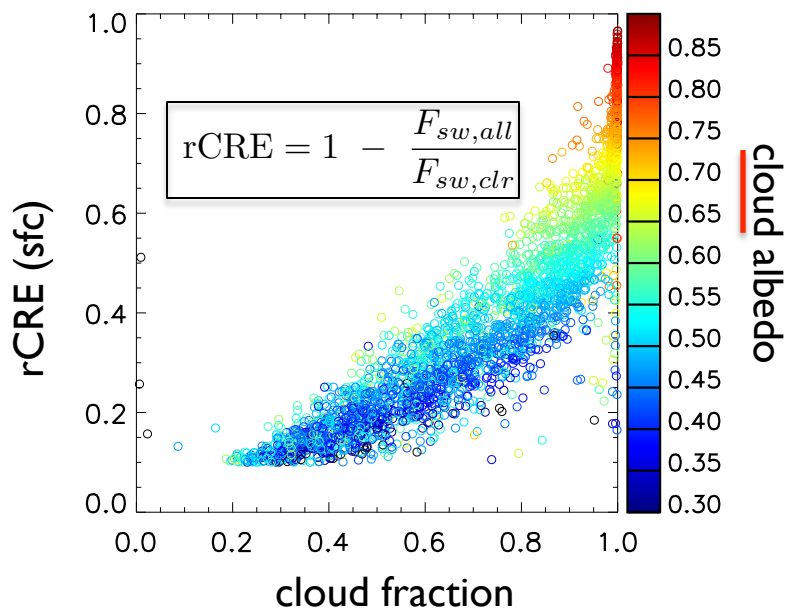
Observations: 14 years of warm clouds at SGP Continental US



$$r_{\text{CRE}} = 1 - \frac{F_{sw,all}}{F_{sw,clr}}$$

F_{sw} = downwelling
shortwave flux

Observations: 12 years of warm clouds at SGP
Continental US (McComiskey and Long, to be submitted)

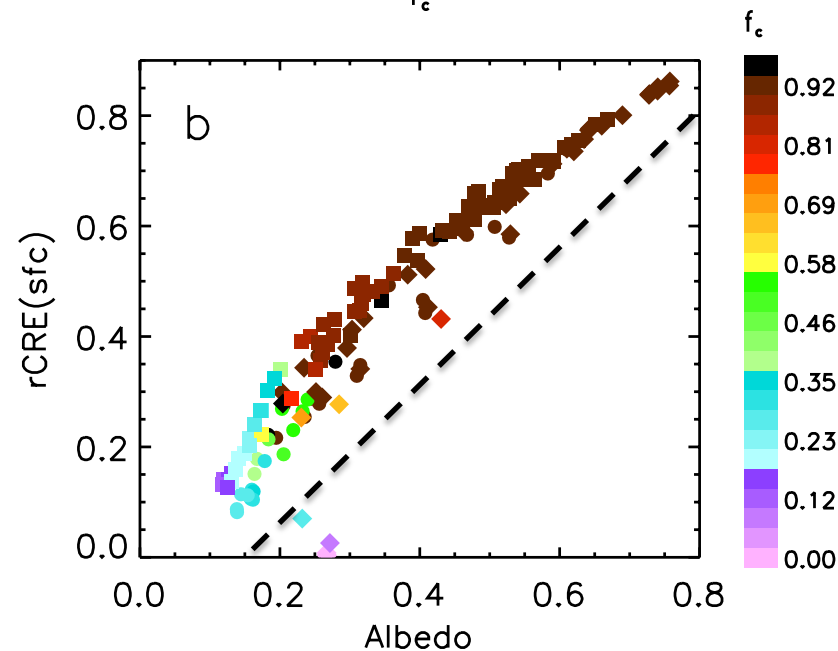
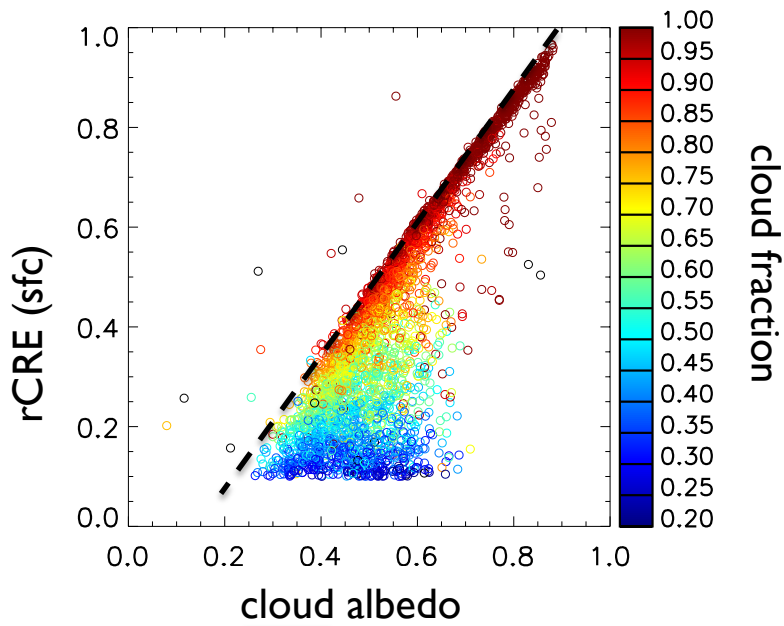
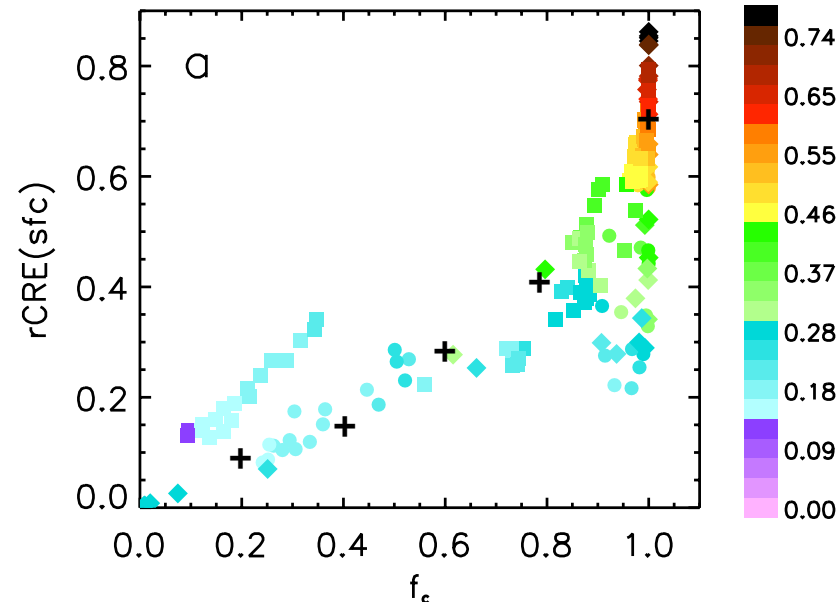
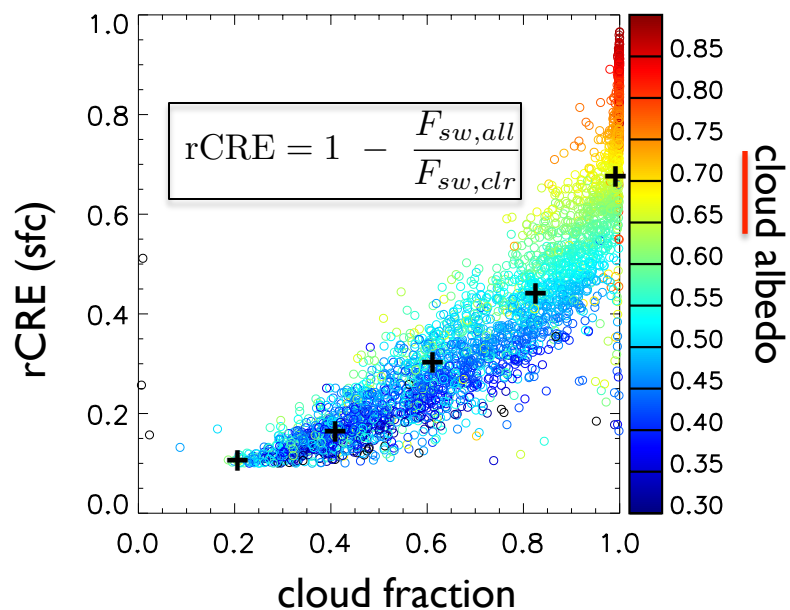


Observations: 12 years of warm clouds at SGP
 Continental US (McComiskey and Long, to be submitted)

LES/CRM Modeling

Ocean

Albedo

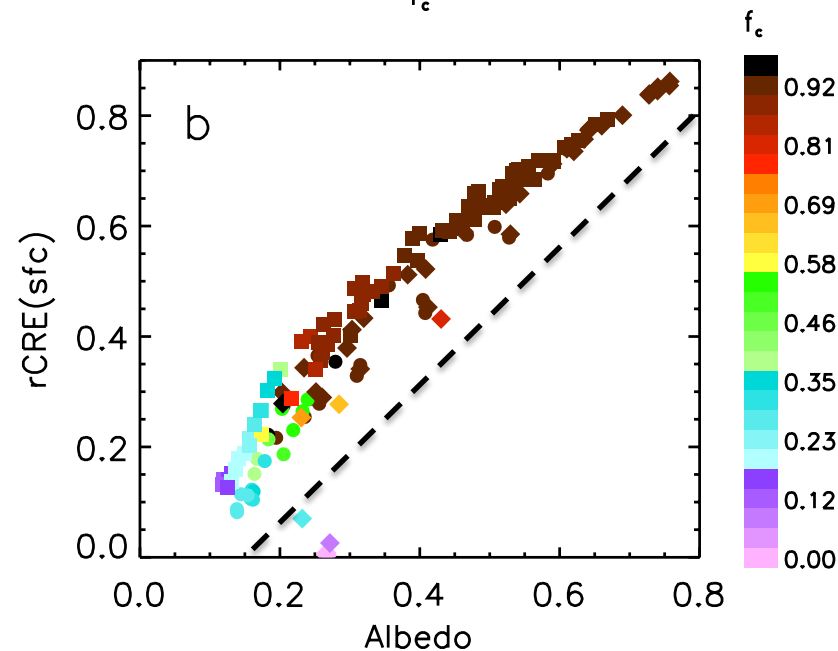
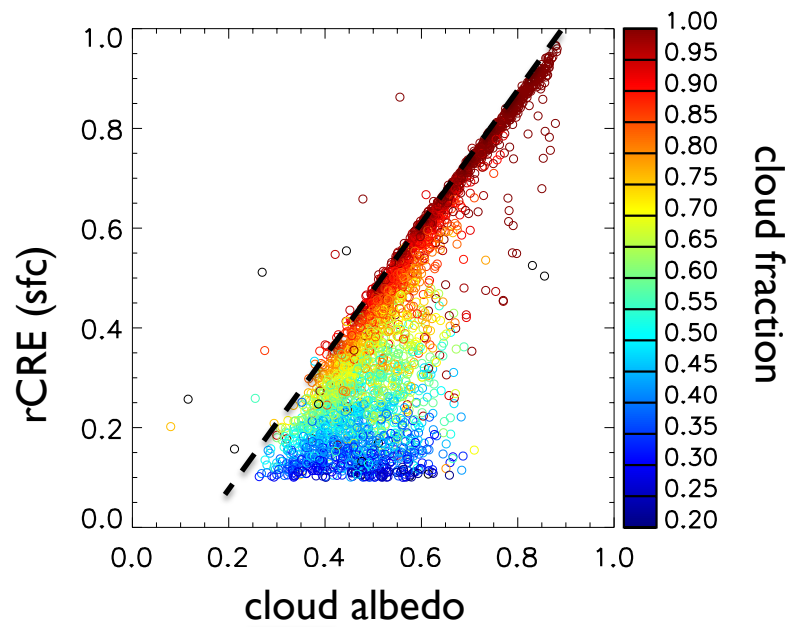
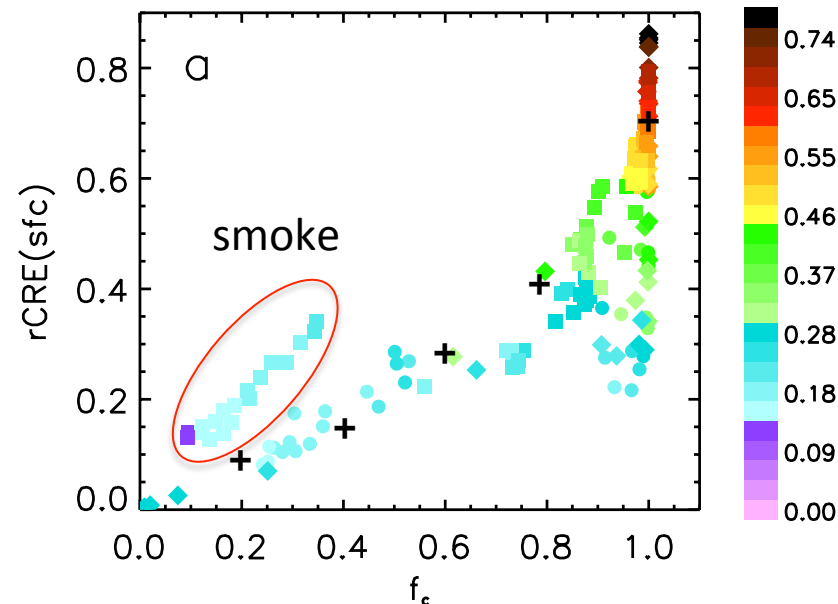
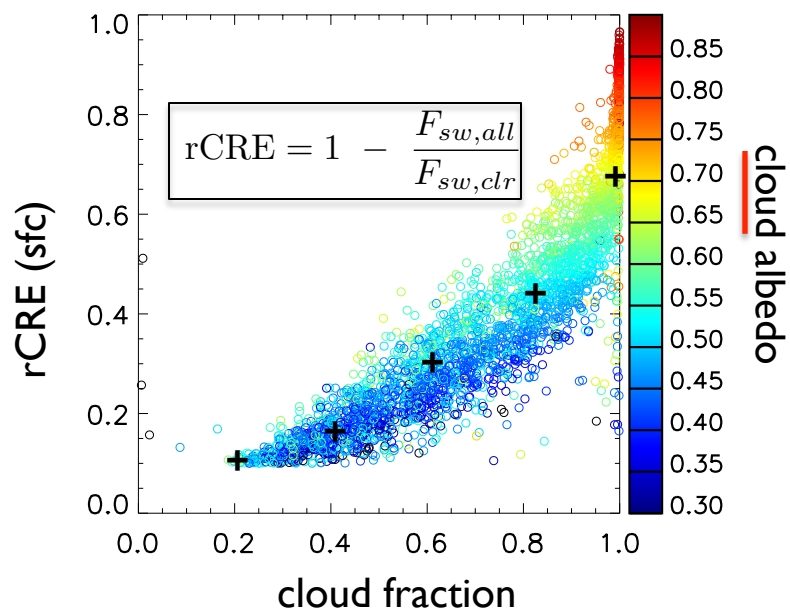


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LES/CRM Modeling

Ocean

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Constraining dL/dN

Can S_o or S_{pop} be used to constrain dL/dN

Large Eddy Simulation (Stratocumulus)

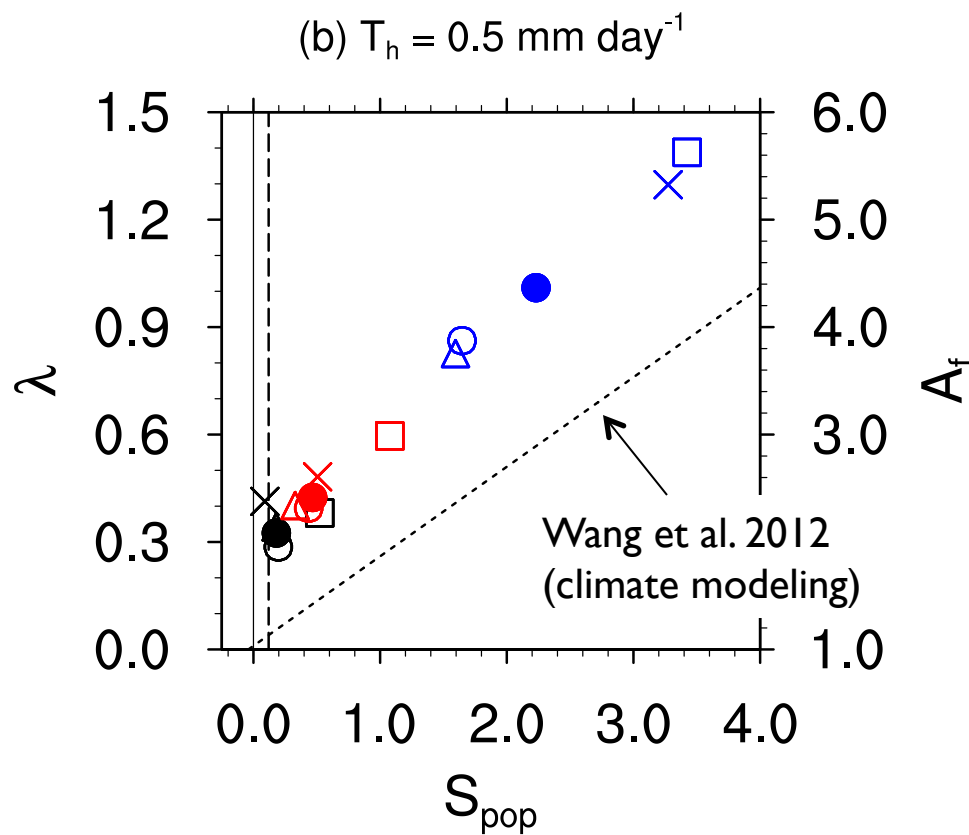
λ is positive; albedo is enhanced

Quantitative differences from GCM

Dependence on R threshold
(not shown)

$$\lambda = \frac{d \ln L}{d \ln N}$$

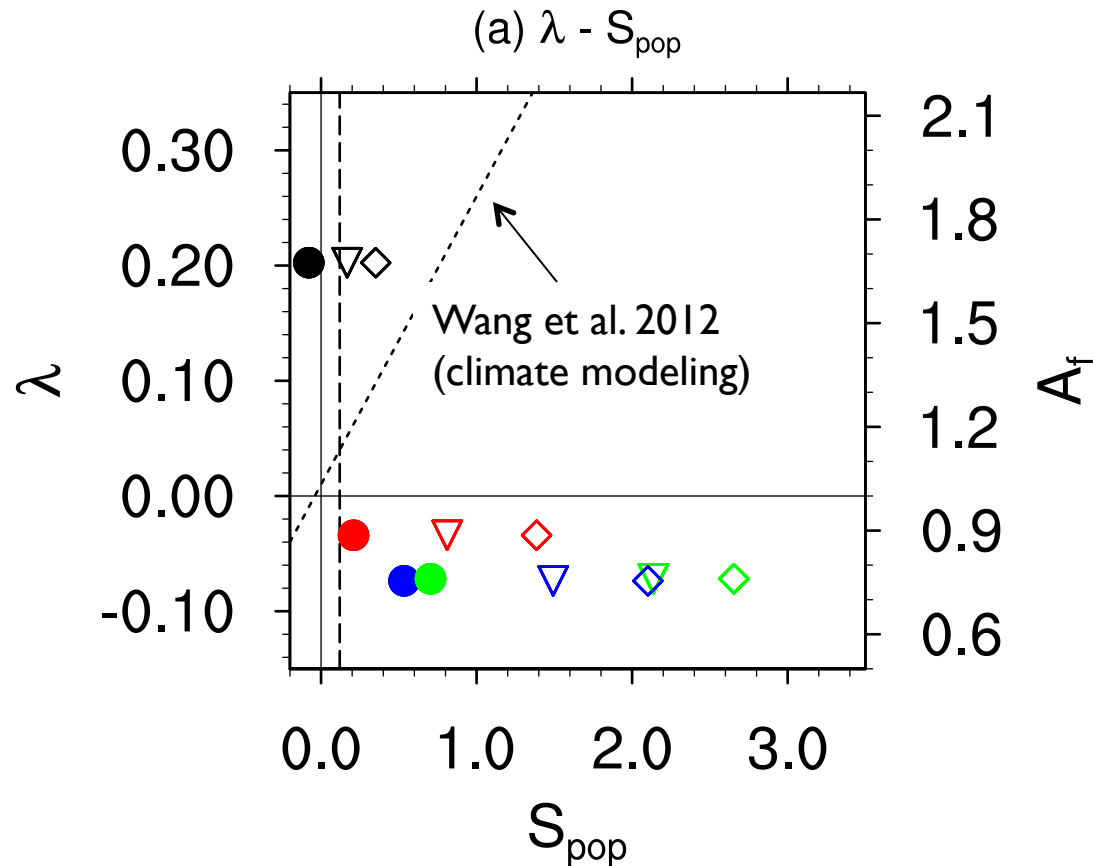
A_f = Albedo enhancement
factor due to λ
(enhancement over Twomey albedo
effect)



Constraining dL/dN

Large Eddy Simulation (Cumulus)

λ becomes negative for large enough aerosol perturbations



Symbol types represent
different R thresholds for POP

Summary

Observational constraints on Cloud Radiative Effect (CRE)

- Top-down vs. bottom-up approach
 - More focus on macroscale parameters tightly linked to radiation (CRE, A , f_c , L)
- Link CRE, A , f_c plots to:
 - microphysics, cloud field properties, scale, place
 - co-variability in controlling parameters

Summary

Observational constraints on Cloud Radiative Effect (CRE)

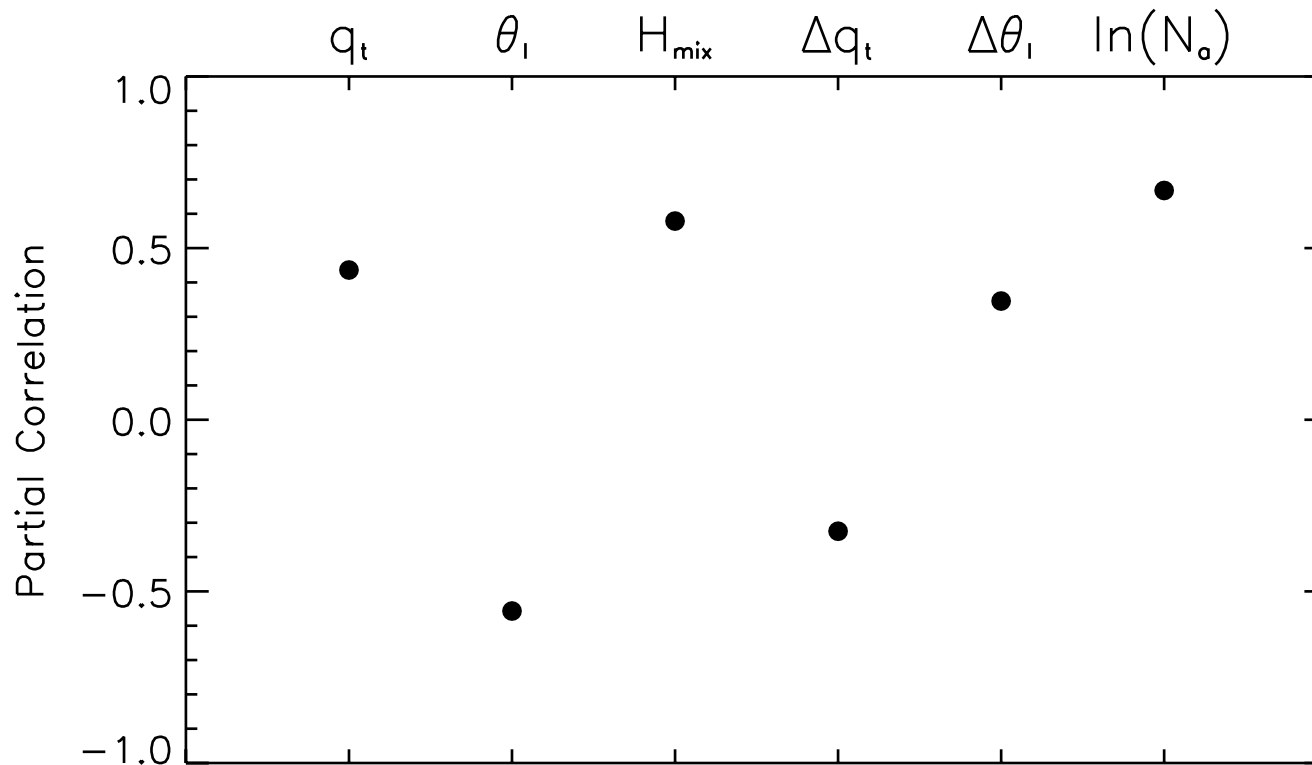
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Co-variability between aerosol and meteorology affects detectability of ACI

- Embrace the co-variability!
- Routine process modeling and observations will capture this co-variability (LASSO; Cabauw, SGP, HDCP2)

Cumulative Multivariate Correlation: Set I

$$\text{Albedo} = f(q_t, \theta_l, H_{\text{mix}}, \Delta q_t, \Delta \theta_l, \ln(N_o))$$



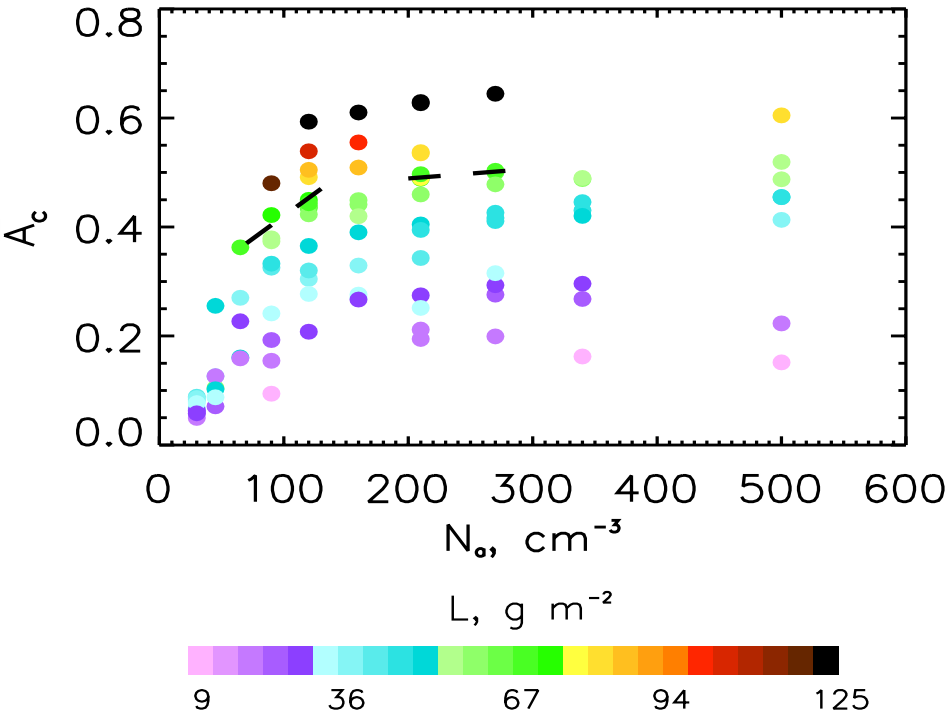
All variables contribute to the Albedo

Albedo Susceptibility

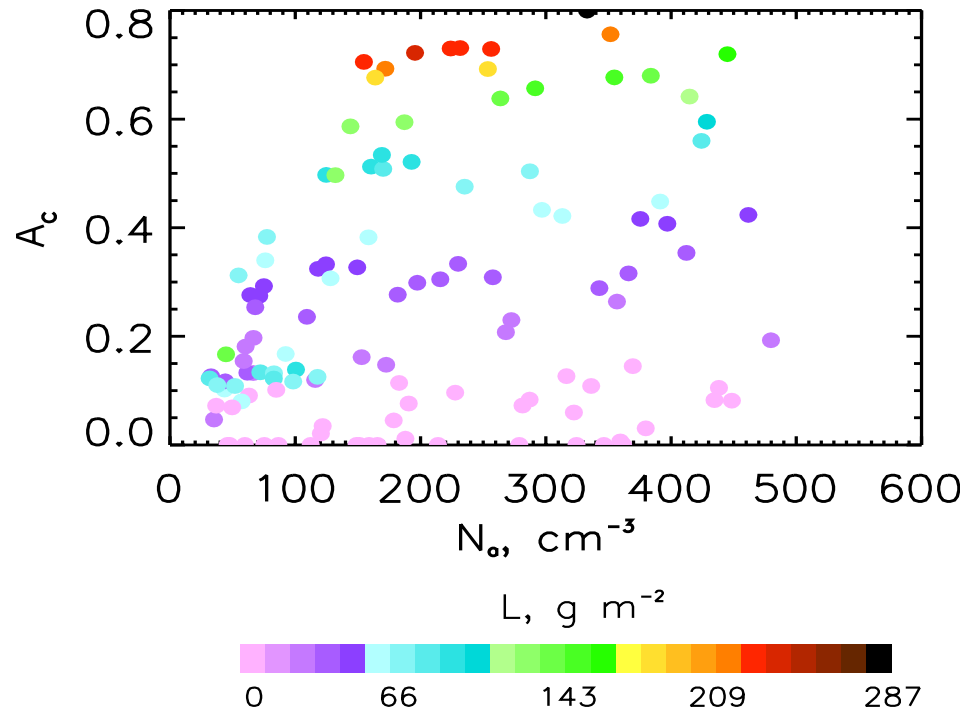
$$S_a = \left. \frac{\partial A_c}{\partial N_d} \right|_L = \frac{A_c(1 - A_c)}{3N_d}$$

Both Sets show expected steeper slopes at small N and $A_c \sim 0.5$

Set 1



Set 2



A_c calculated based on plane parallel assumptions

$$A = A_c f_c + A_s(1 - f_c)$$

Set I

$N_o, \text{ cm}^{-3}$

