

# Effects of decoupling boundary layer on the change of phase partitioning in the mixed-phase stratiform clouds

**Fan Yang**<sup>1</sup>, Mikhail Ovchinnikov<sup>2</sup>, Damao Zhang<sup>1</sup>, Edward Luke<sup>1</sup>,  
Mariko Oue<sup>3</sup>, Dan Lubin<sup>4</sup>, Pavlos Kollias<sup>1,3</sup>, Andrew Vogelmann<sup>1</sup>



1 Brookhaven National Laboratory,  
Upton, New York, USA

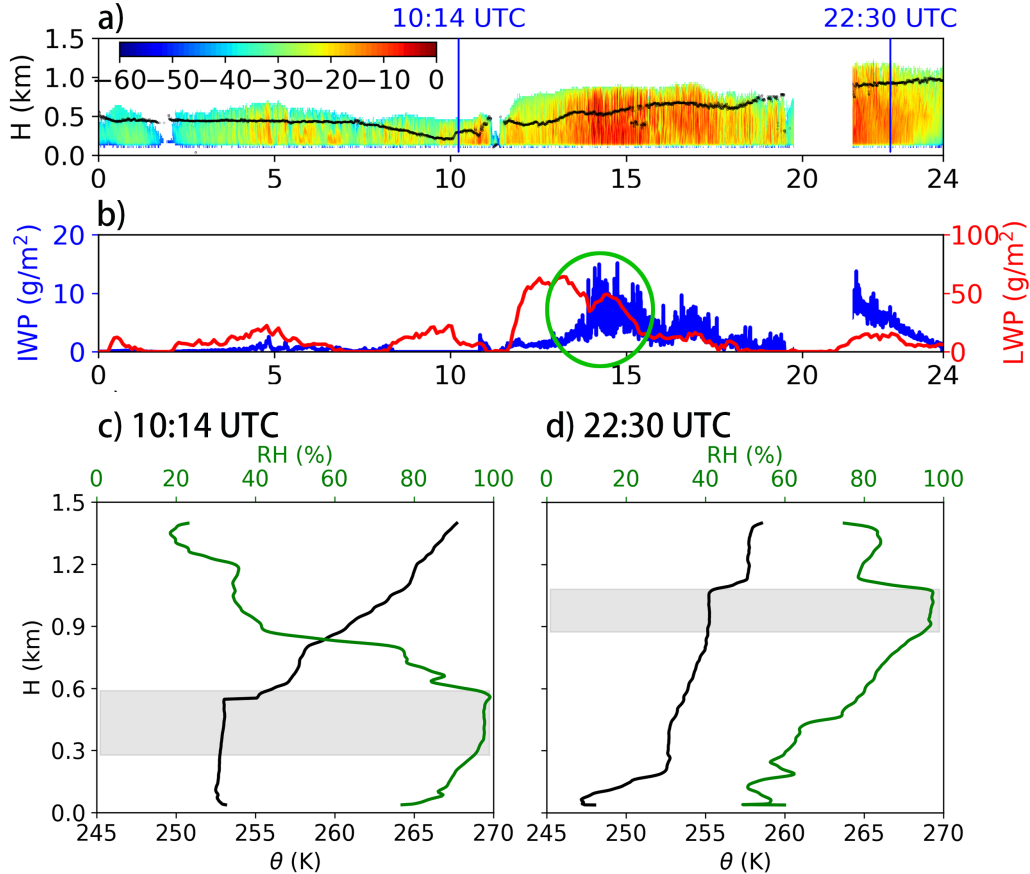
2 Pacific Northwest National Laboratory,  
Richland, Washington, USA

3 School of Marine and Atmospheric  
Sciences, Stony Brook University,  
Stony Brook, New York, USA

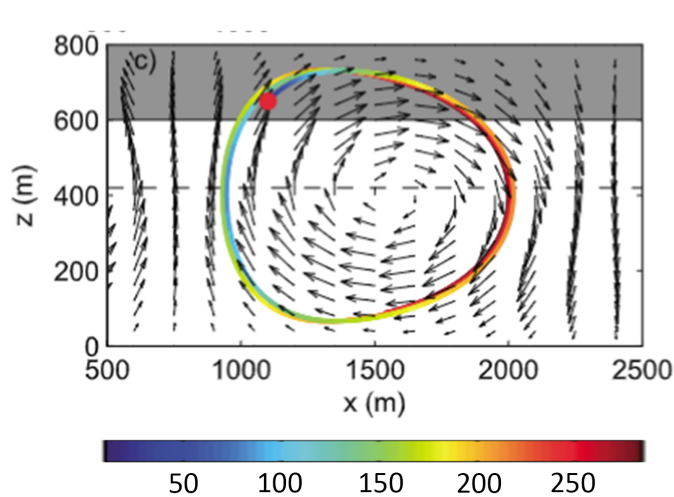
4 Scripps Institution of Oceanography,  
University of California San Diego,  
California, USA

June 12, 2019

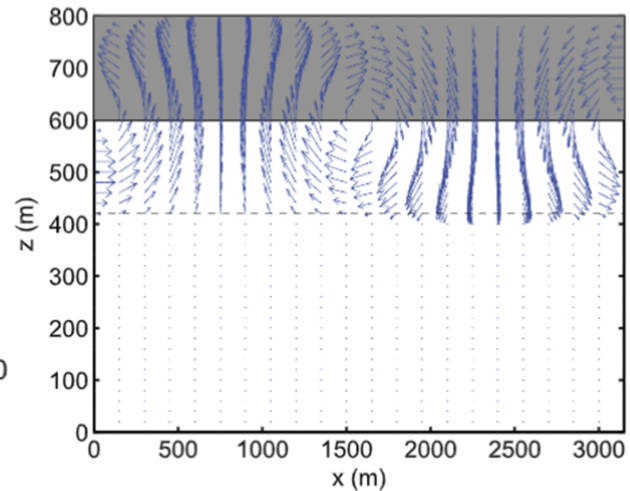
# March 31, 2016, AWARE



# Effect of decoupled PBL on phase partitioning



coupled  
low IWP

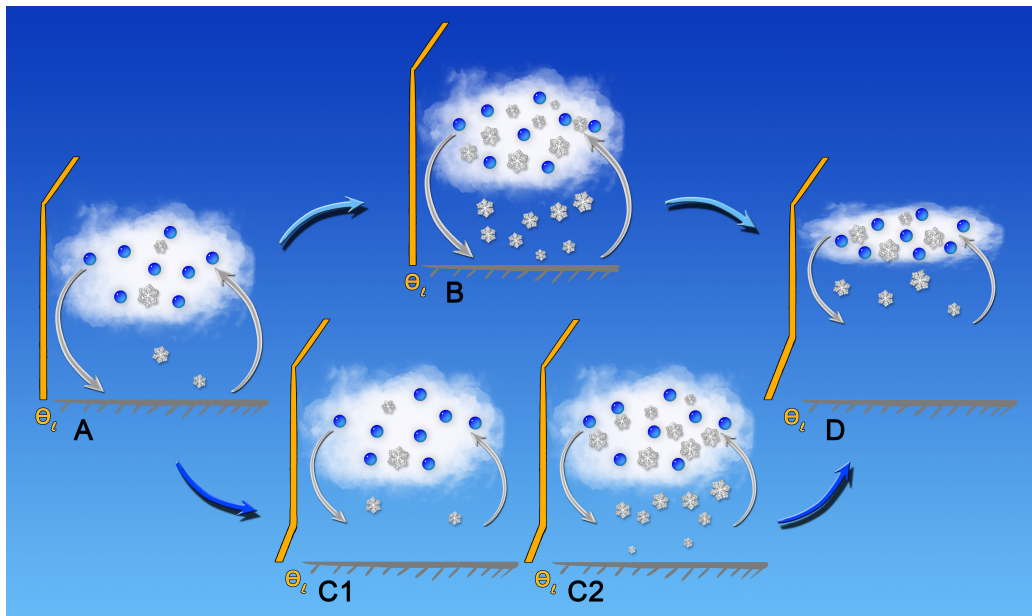


decoupled  
high IWP

Yang et al. (2015) suggested that IWP in a decoupled field is larger than that in a coupled field with the same mixed-phase cloud thickness and ice nucleation rate<sup>1</sup>.

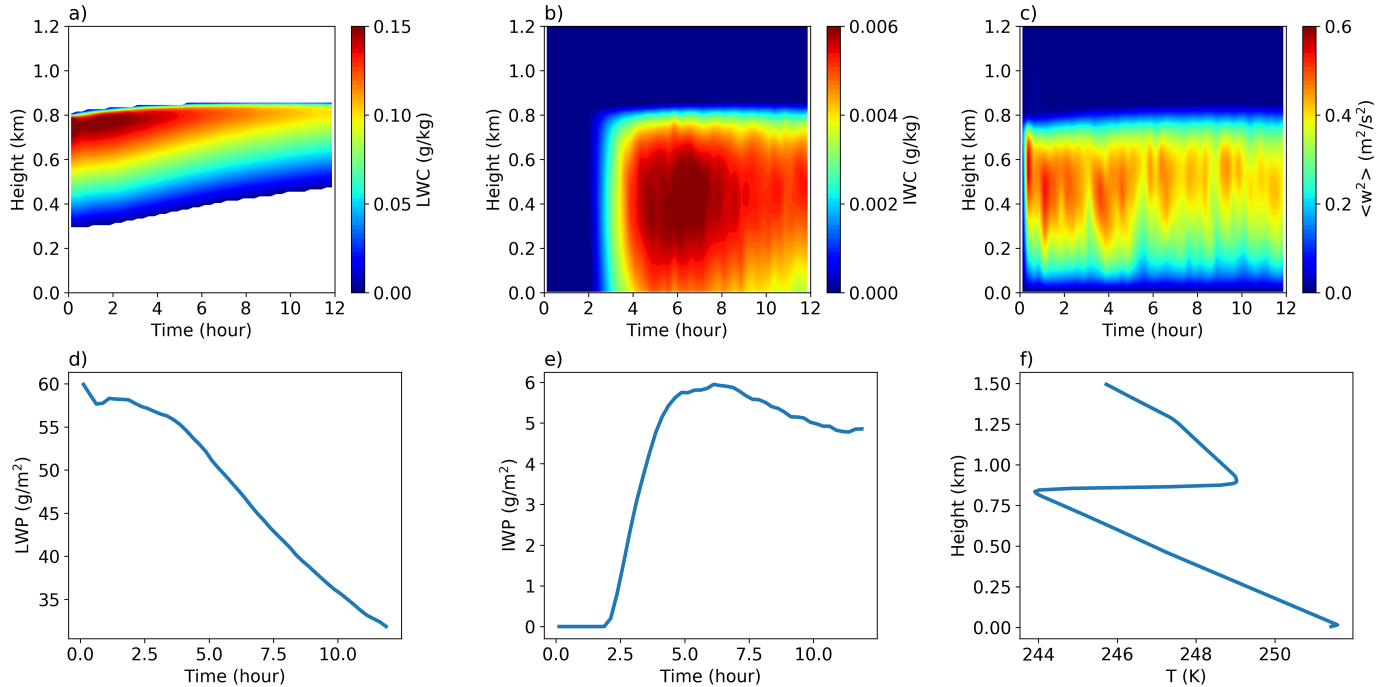
1. Yang et al., JGR, 2015

# Working hypothesis



Is the coupled-to-decoupled transition of the atmospheric boundary layer the **cause** or the **result** of the fast change of phase partitioning in the mixed-phase stratiform clouds?

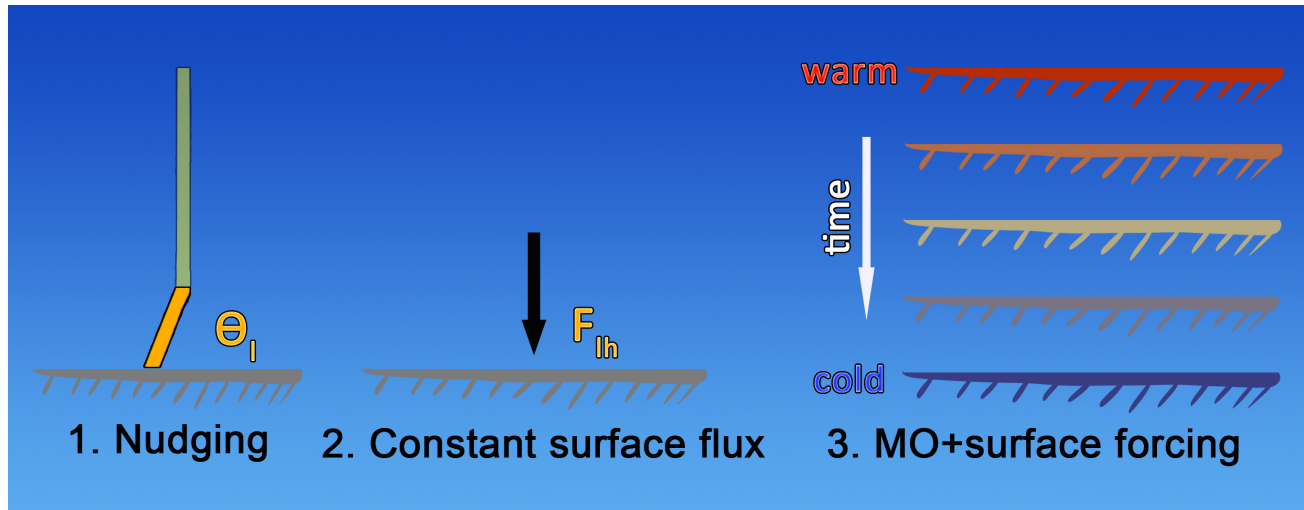
# Control run



This simulation is similar to the ISDAC<sup>1</sup> case.

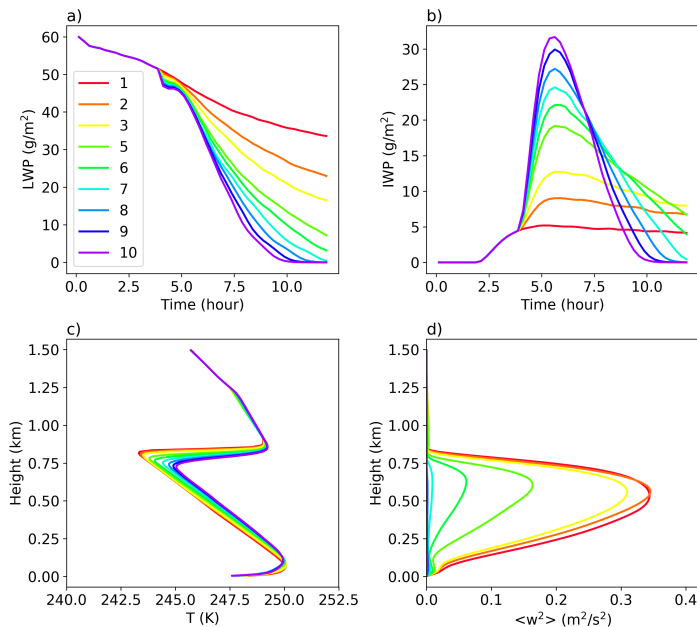
1. Ovchinnikov et al., JAMES, 2014

# Ways to generate surface inversion



**Conclusion I:** Decoupled PBL has **minor** effect in *LWP* and *IWP*. For details, please come to see our poster **B2-99**.

# Important effect of ice number concentration



**Conclusion II:** The “only” way that I can mimic the observation is considering both **the change of ice number concentration** (fast change of phase partitioning) and **land-atmosphere interaction** (surface inversion).

Lines in the figure represent different values of ice number concentration ( $L^{-1}$ ) in the mixed-phase clouds.

# Conclusion and discussion

- ▶ Changing ice number concentration significantly alter the *LWP* and *IWP* time evolution.
- ▶ Addition of surface flux/inversion reduces the intensity of BL turbulence ( $\langle w^2 \rangle$ ) and introduces **minor** quantitative changes in *LWP* and *IWP*.
- ▶ The coupled-to-decoupled transition of the atmospheric boundary layer is **unlikely** to be the main cause of the observed fast change of phase partitioning.

Please come to see our poster **B2-99**.



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# Backup Slides

# Model setup

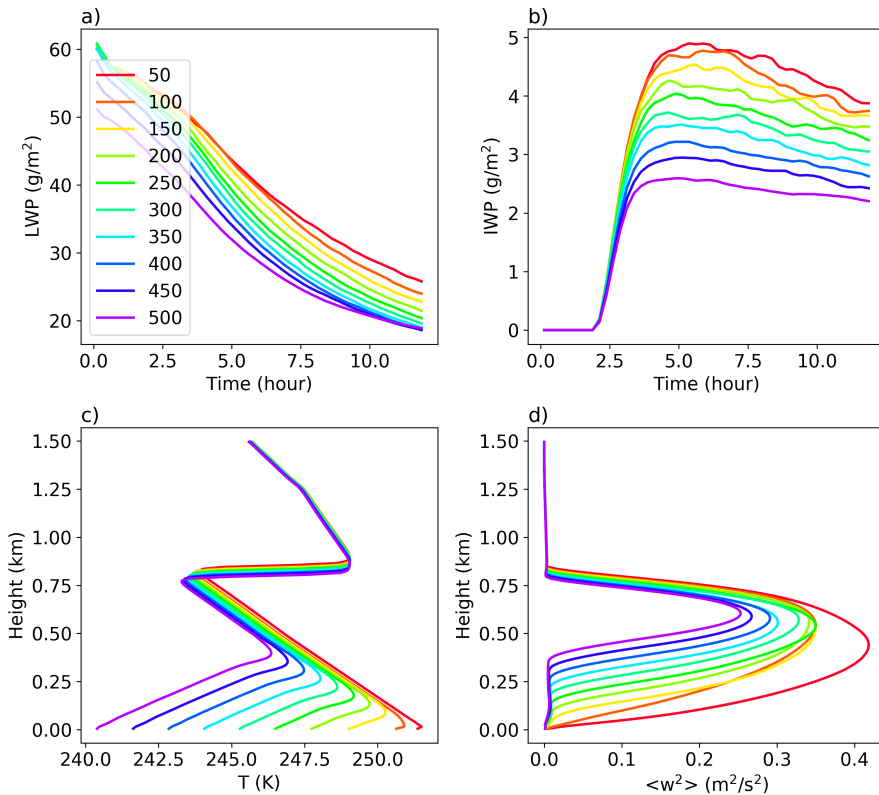
The simulation is similar to ISDAC<sup>1</sup> case with some modifications.

Model	:	System for Atmospheric Modeling <sup>2</sup> (SAM 6.11.2)
Resolution	:	50 m × 50 m × 10 m
Domain	:	3.2 km × 3.2 km × 1.5 km
Total Time	:	12 hours
Profiles	:	Sounding at 10:14 UTC
Radiation	:	longwave radiation from NCAR CAM3 model
Microphysics	:	Morrison (2009) two-moment $\mu$ physical scheme <sup>3</sup>
Forcing	:	$w_{ls}$
Nudging	:	$u, v, \theta_l, q_t$
Surface	:	no flux, MO+surface forcing, constant flux

1. Ovchinnikov et al., JAMES, 2014      2. SAM

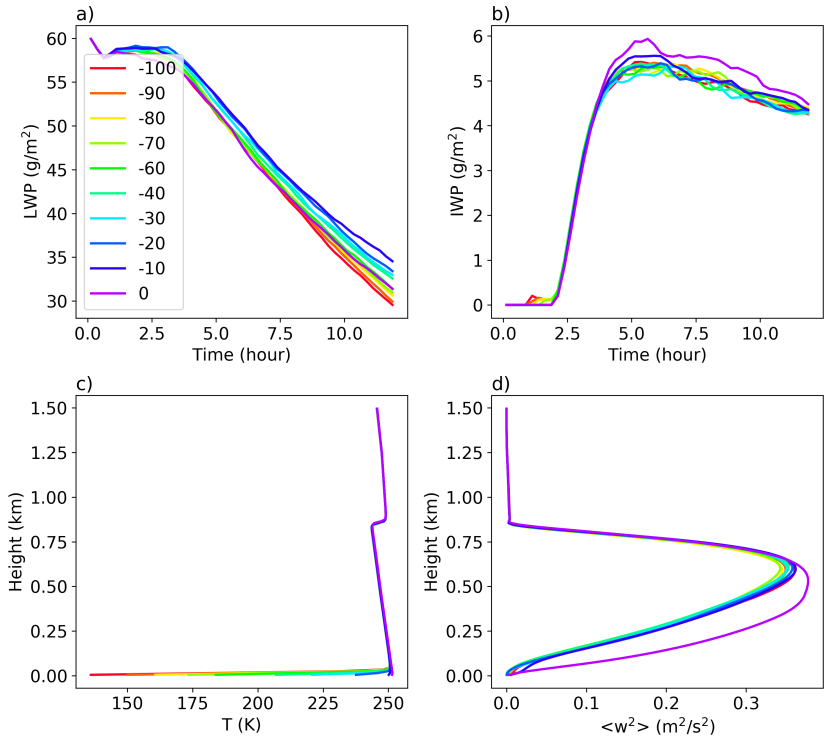
3. Morrison et al., MWR, 2009

# (1) Nudging-induced inversion layer



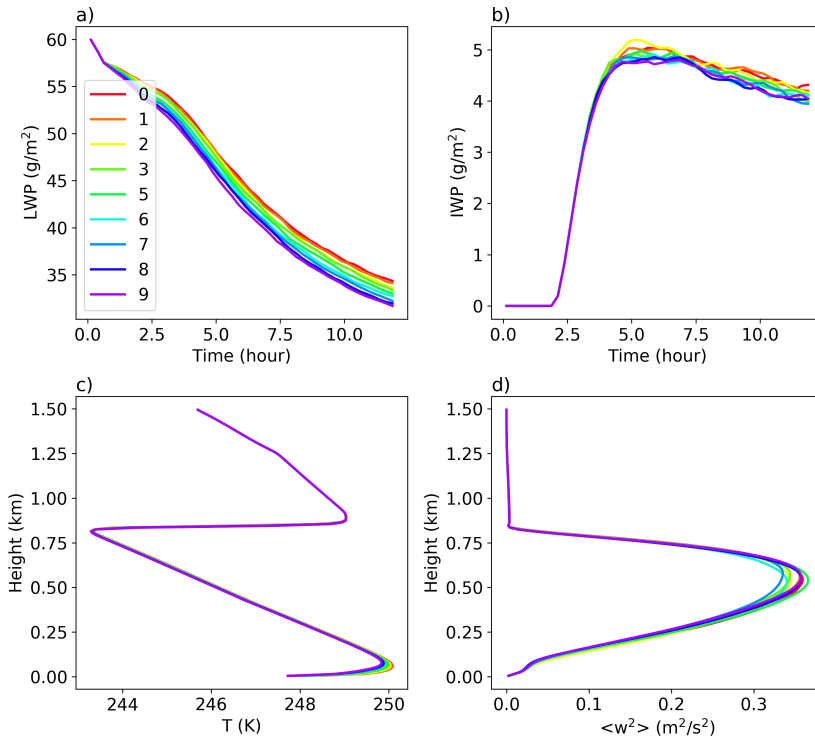
Lines in the figure represent different thicknesses of the nudged-inversion layers above the surface.

## (2) $F_{lh}$ -induced inversion layer



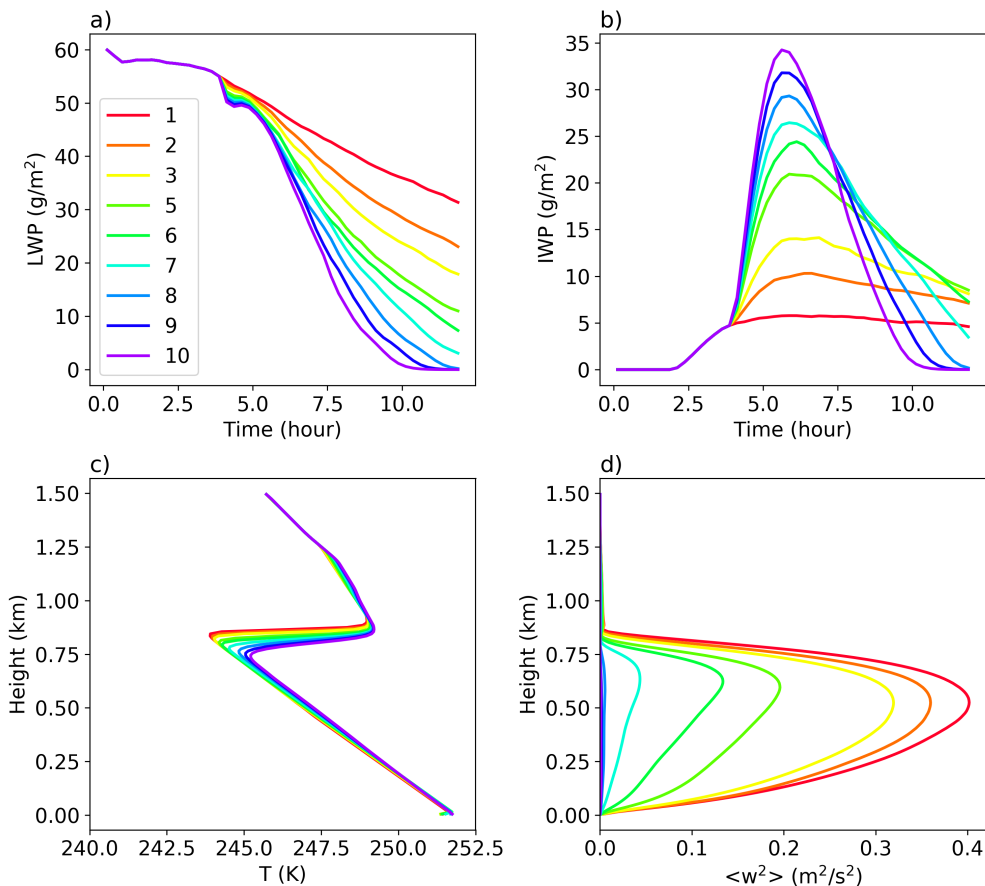
Lines in the figure represent different values of  $F_{lh}$  at the surface.

# (3) $T_s$ -induced inversion layer



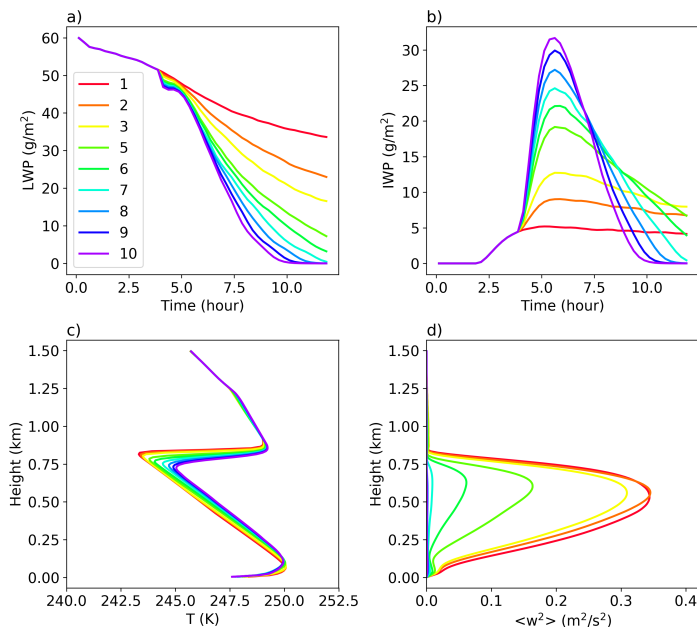
Lines in the figure represent different values of  $z_0 = 2^n \times 10^{-4}$  at the surface.

# Effect of change of ice number concentration



Lines in the figure represent different values of ice number concentration ( $L^{-1}$ ) in the mixed-phase clouds.

# Important effect of ice number concentration



**Conclusion II:** The “only” way that I can mimic the observation is considering both **the change of ice number concentration (fast change of phase partitioning)** and **land-atmosphere interaction (surface inversion)**.

Lines in the figure represent different values of ice number concentration ( $L^{-1}$ ) in the mixed-phase clouds.