## Cross-Scale Land-Atmosphere Experiment (CSLAEX) PI: Pierre Gentine – Student: Yu Cheng Science Columbia University









#### Main objective: Understanding the space-time nature of land-atmosphere interactions

Site: Department Of Energy **Southern Great Plains** 

Main strategy: Using optical fibers and Distributed Temperature Sensing (DTS) Span 1 mile transect





## Why this setup?





## **First step** Initial smaller setup 200m - 2016 - MOISST site



Site in OK, near OSU Similar landscape but smaller and with existing DTS for soil moisture measurements

Campaign lasted for two weeks and installations stayed there over the summer (May to August)

Started acquiring and using data

Several issues: Tensions, breaks, animals



#### Second step Full installation SGP site – Summer 2017

#### Initial envisioned setup





#### Second step Full installation SGP site – Summer 2017



#### **Final setup**





#### Second step Full installation SGP site - Summer 2017

## Depth and heights changed a bit Also installation had to adjust (no pulleys - too much tension - and a lot of issues...)







## Second step Full installation SGP site - Summer 2017 Spans two land cover types (alpha alpha and prairie)





## How does it work? Span fiber linked to DTS (in a shelter), and need some reference temperature bath for gradient reconstruction Based on Ramann scattering



#### Can know T(x,t)





#### How does it work?





## 1. Taylor frozen hypothesis



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1. Taylor frozen hypothesis Is turbulence really frozen, i.e. U does not depend on eddy size Or is varying?

Use DTS to compute U(k) with k wavenumber, based on phase difference

$$U(k_1) = \frac{\Delta \varphi}{k_1 \times \Delta t},$$

**Convective velocity (wavenumber dependent)** 

Usually assumed to be **constant** 

Cheng et al. 2017 GRL



### 1. Taylor frozen hypothesis Results from DTS



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## Across stability far from constant!



#### 1. Taylor frozen hypothesis Could it be an artefact from sluggish DTS measurements



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## Checked with LES A real signal



#### 1. Taylor frozen hypothesis Then develop a theory based on eddy diffusion at small scale



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# $10^{1}$ 10

## Departure is eddy size dependent.

# Diffusion of coherence



## 2. Stable boundary layer Challenging to model – too many scales Plateau in energy spectrum



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## 2. Stable boundary layer Challenging to model – too many scales Plateau in energy spectrum



Eddy covariance (time using Taylor hypothesis)

#### Plateau confirmed by DTS

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2. Stable boundary layer Theory: two main time scales Ozmidov scale (stratification and dissipation)  $L_0$ Buoyancy scale (stratification and mean wind) L<sub>B</sub> Can develop a theory based on Weinstock (1978) explaining why we have this spectrum



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2. Stable boundary layer Theory: two main time scales Implication: Failure of Monin-Obukhov Similarity theory for stable (and especially very stable boundary layers)



 $\frac{\partial X}{\partial z} = f(\frac{z}{L})$ With *L* the Obukhov length

Missing an independent variable:  $L_{R}$ and slope of plateau (As  $L_0$  scales with L – see Li et al. 2016)

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**Current work:** 1. Correcting Monin-Obukhov Similarity Theory (MOST) in unstable conditions, Large eddies are very important and need to be included AS they lead to departure from z/L scaling



Current work:

2. Investigating heterogeneity in ground heat flux (G) and soil moisture at ~0.1m resolution and implication for surface energy balance





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## 2. Investigating heterogeneity in ground heat flux (G) and soil moisture at ~0.1m resolution and implication for surface energy balance









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Svstem Research







#### Data is part of the DOE ARM SGP site

Publicly available https://adc.arm.gov

Any question please contact me pg2328@Columbia.edu

We hope people use the data

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Atmospheric System Research

