

Does vertical velocity influence entrainment in moist thermals?

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Background and motivation

- Previous observational and numerical modeling studies have found an inverse correlation/scaling between in-cloud vertical velocity (w) and fractional entrainment rate (ϵ)
 - *Tian and Kuang (2016; GRL); Zhang et al. (2016; CLIM DYN); Kirshbaum and Lamer (2021; JAS)*
- Numerous cumulus parameterization schemes invoke the proposed inverse relationship between w and ϵ
 - *Lin (1999; JAS); Gregory (2001; QJRMS); Neggers et al. (2002; JAS); de Roode et al. (2012; MWR); de Rooy and Siebesma (2010; QRJMS); Tian and Kuang (2016; GRL); Tan et al. (2018; JAMES)*
- “Chicken & egg argument”:
 - Does w affect ϵ ?
 - Does ϵ affect w ?
 - Or do both affect each other in some way?



Scientific question and hypothesis

Scientific question:

Do moist thermals with stronger w have smaller ϵ -driven dilution than moist thermals with weaker w ? (for deep convection)

Hypothesis:

Moist thermals with stronger w have less time to be exposed to the free troposphere, and thus, have smaller ϵ -driven dilution.

- Akin to “core-exposure effect” as discussed by *Drueke et al. (2021; ACP)*

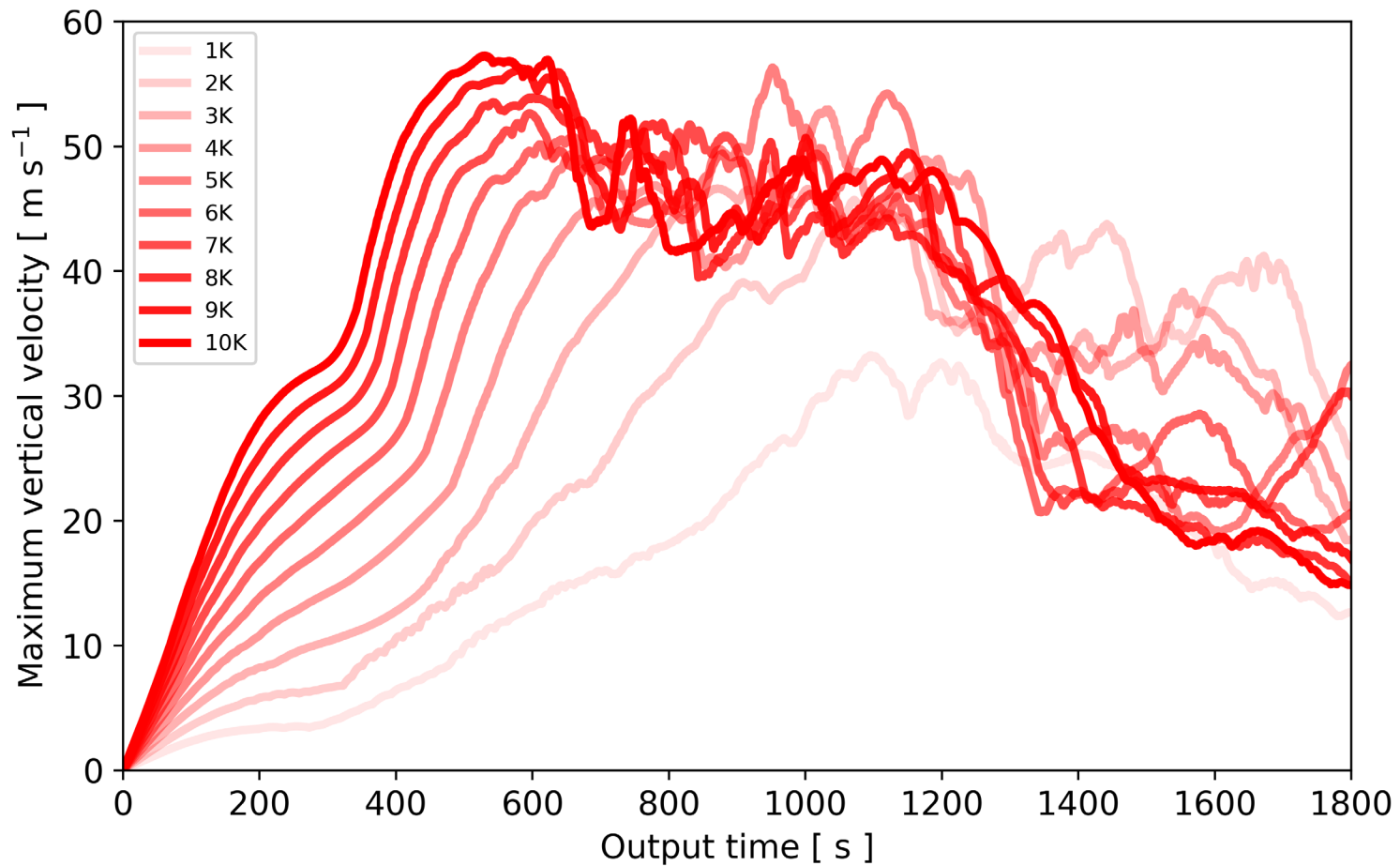


Numerical modeling framework

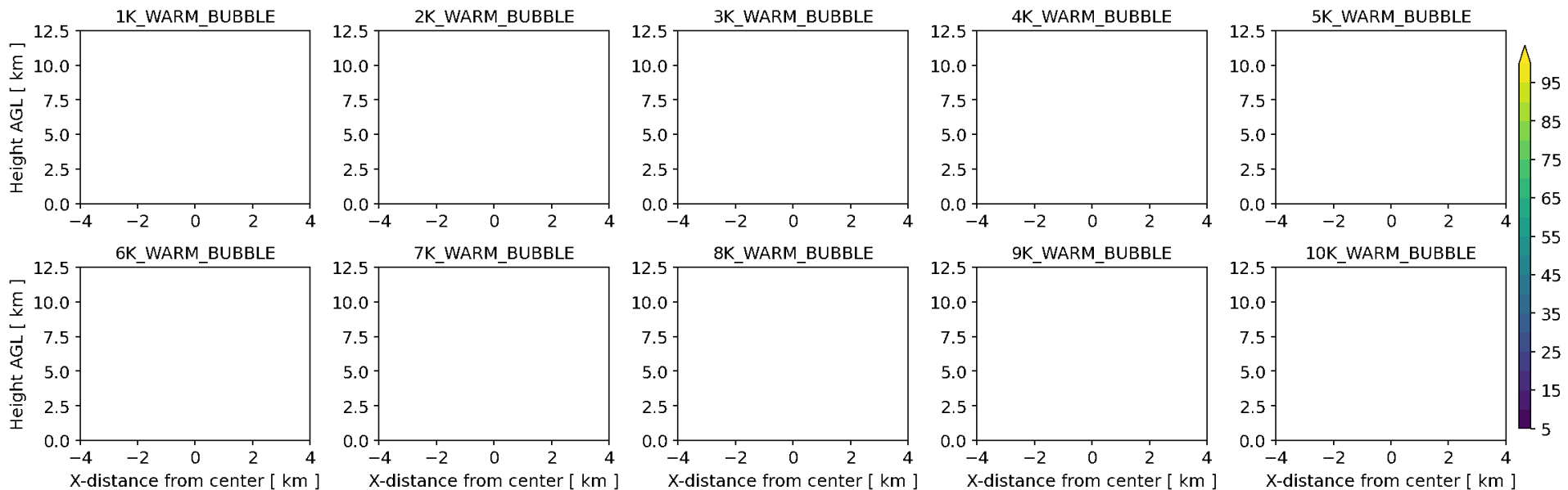
- Idealized “single cloud” simulations using Cloud Model 1 (CM1)
- No winds; [Weisman and Klemp \(1982; MWR\)](#) thermodynamic profile
- “Warm bubble” convection initiation technique
 - Systematically alter potential temperature perturbation to vary w
 - Horiz. radius = 2 km; Vert. radius 1.5 km; Centered at 0.5 km AGL
- 100 m grid spacing in all directions (25 x 25 x 25 km³ domain)
- Morrison two-moment microphysics scheme (“ihail” = graupel)
- Smagorinsky turbulence scheme
- Initial +/- 0.5 K potential temperature perturbations domain-wide
 - “Spin-up” turbulence spectrum similar to [Peters et al. \(2019; JAS\)](#)
- 30-min simulations; 30-s output
- Passive tracer (PT) within 0-1 km AGL high-CAPE/low-CIN layer



Warmer warm bubble = stronger w ?



Stronger w = smaller ϵ -driven dilution?

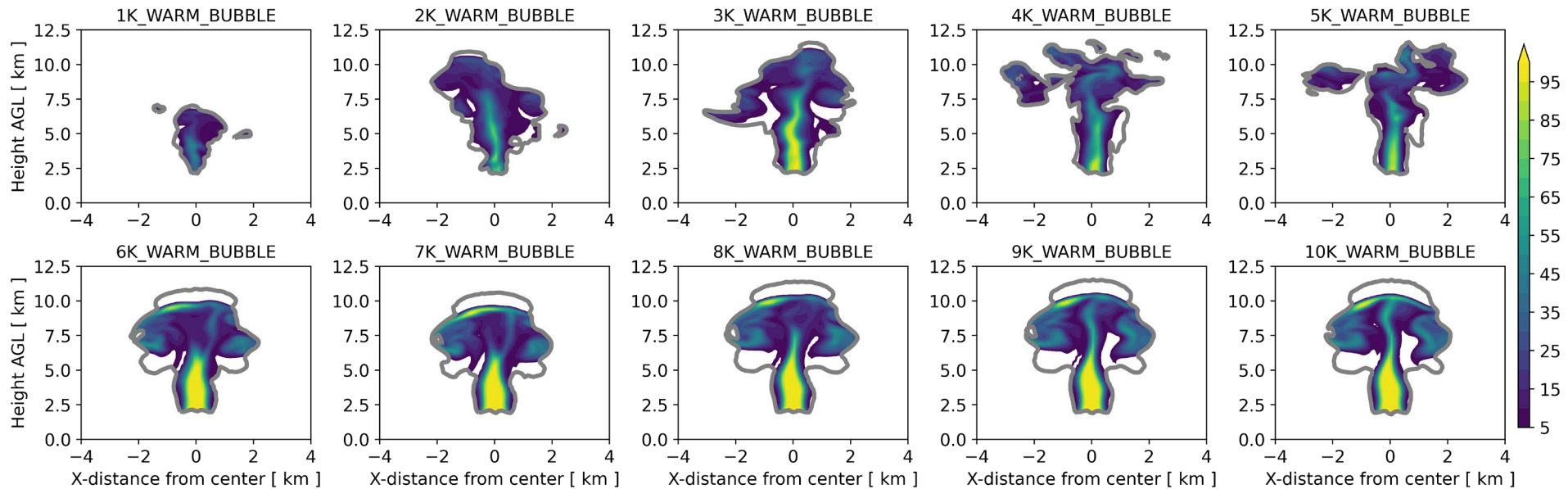


Grey contour = 20 dBZ
Shading = Passive tracer conc. [%]



Stronger w = smaller ϵ -driven *dilution?*

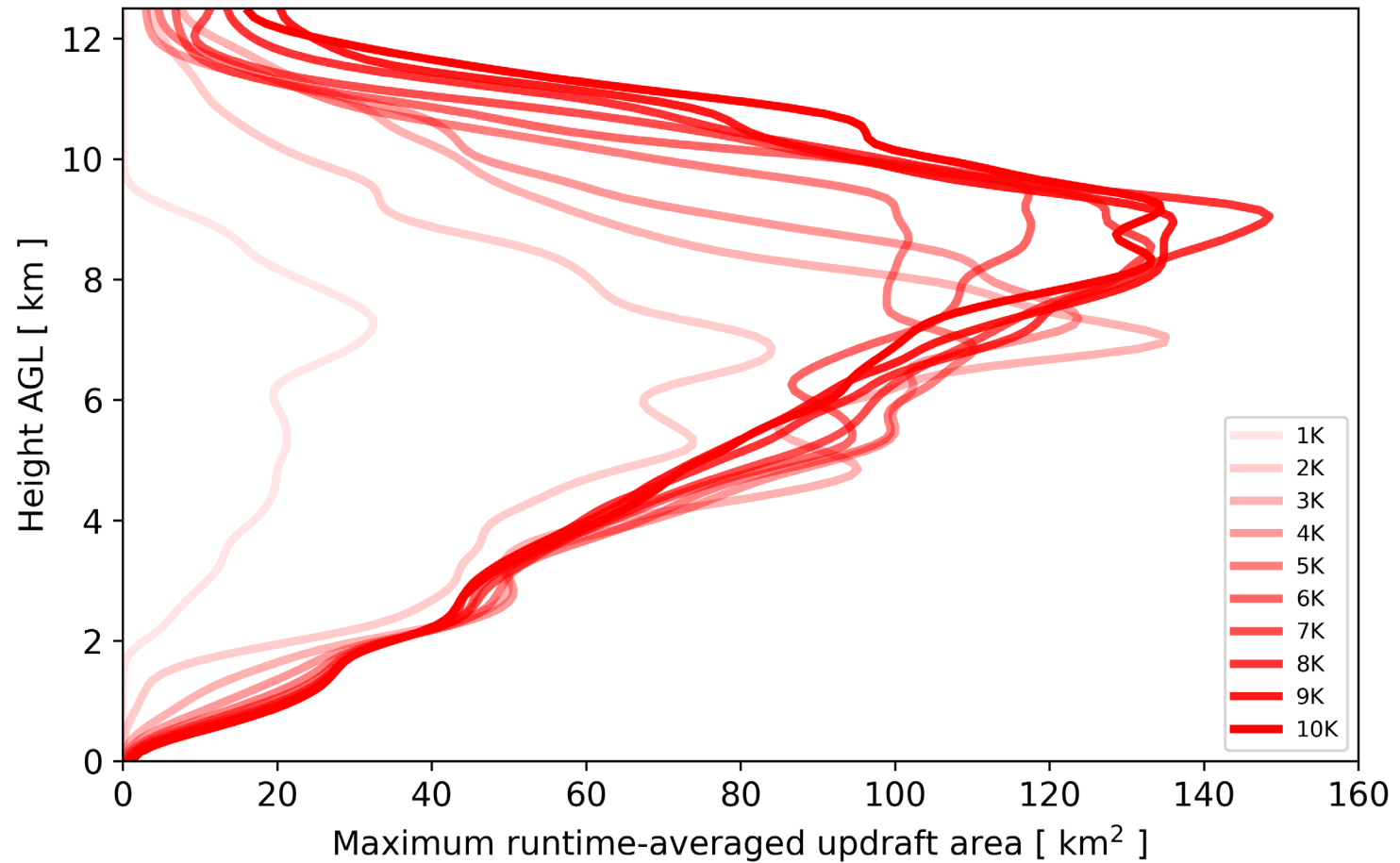
@ time of w_{MAX}



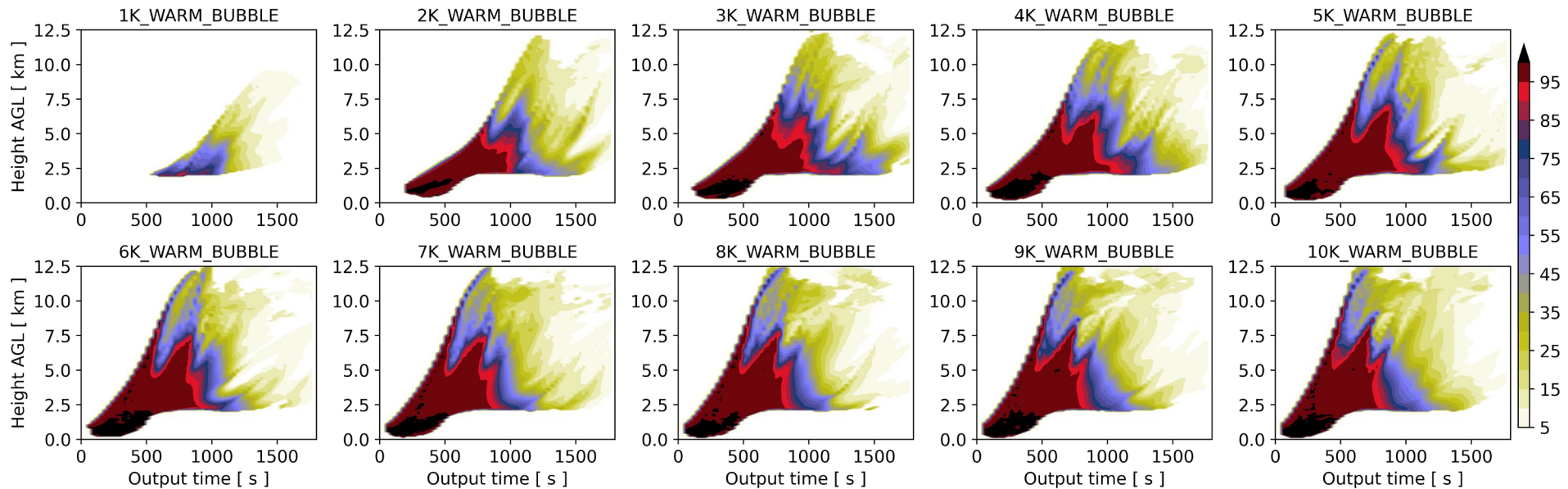
Grey contour = 20 dBZ
Shading = Passive tracer conc. [%]



Does updraft width change?



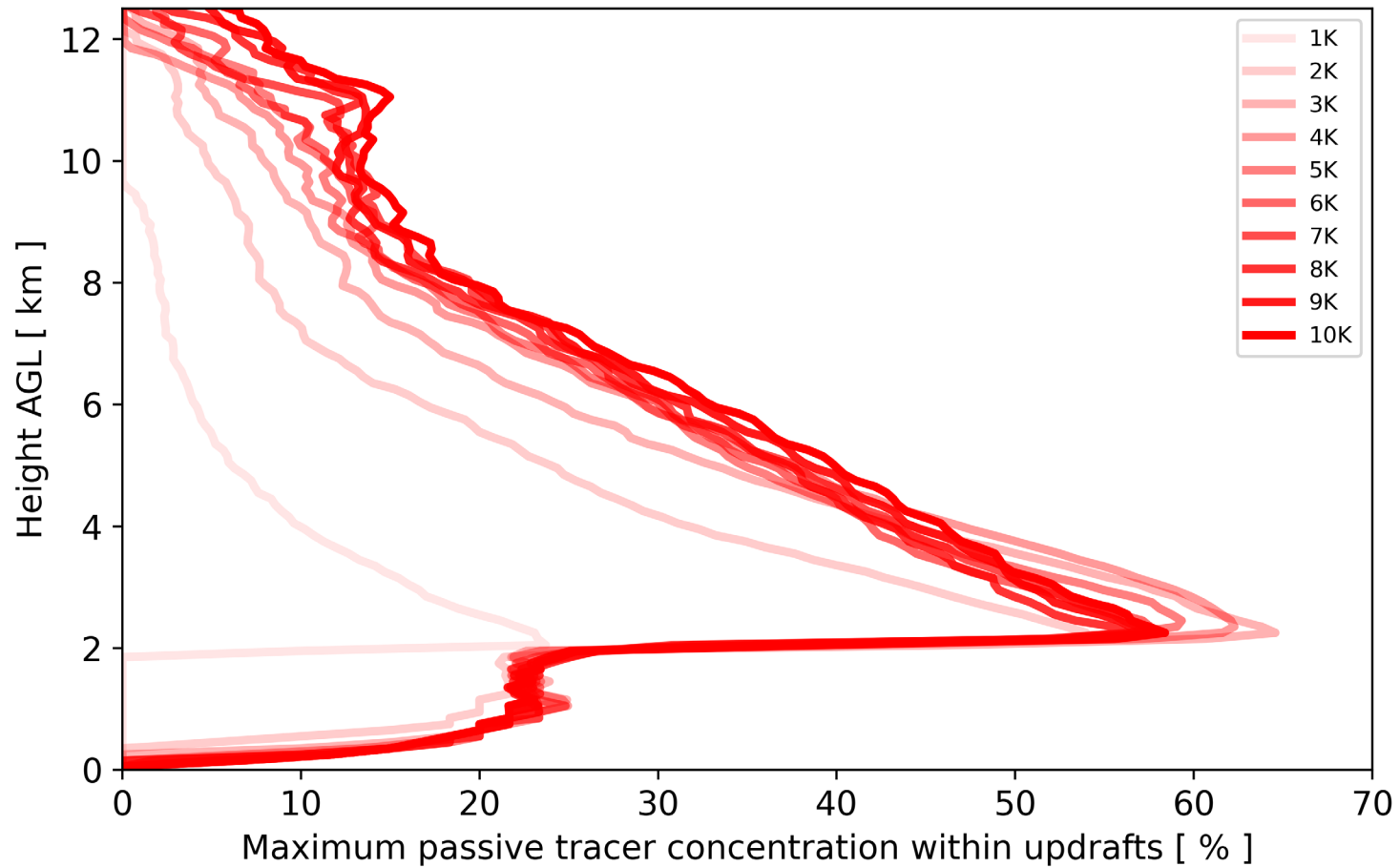
Stronger w = smaller ϵ -driven dilution?



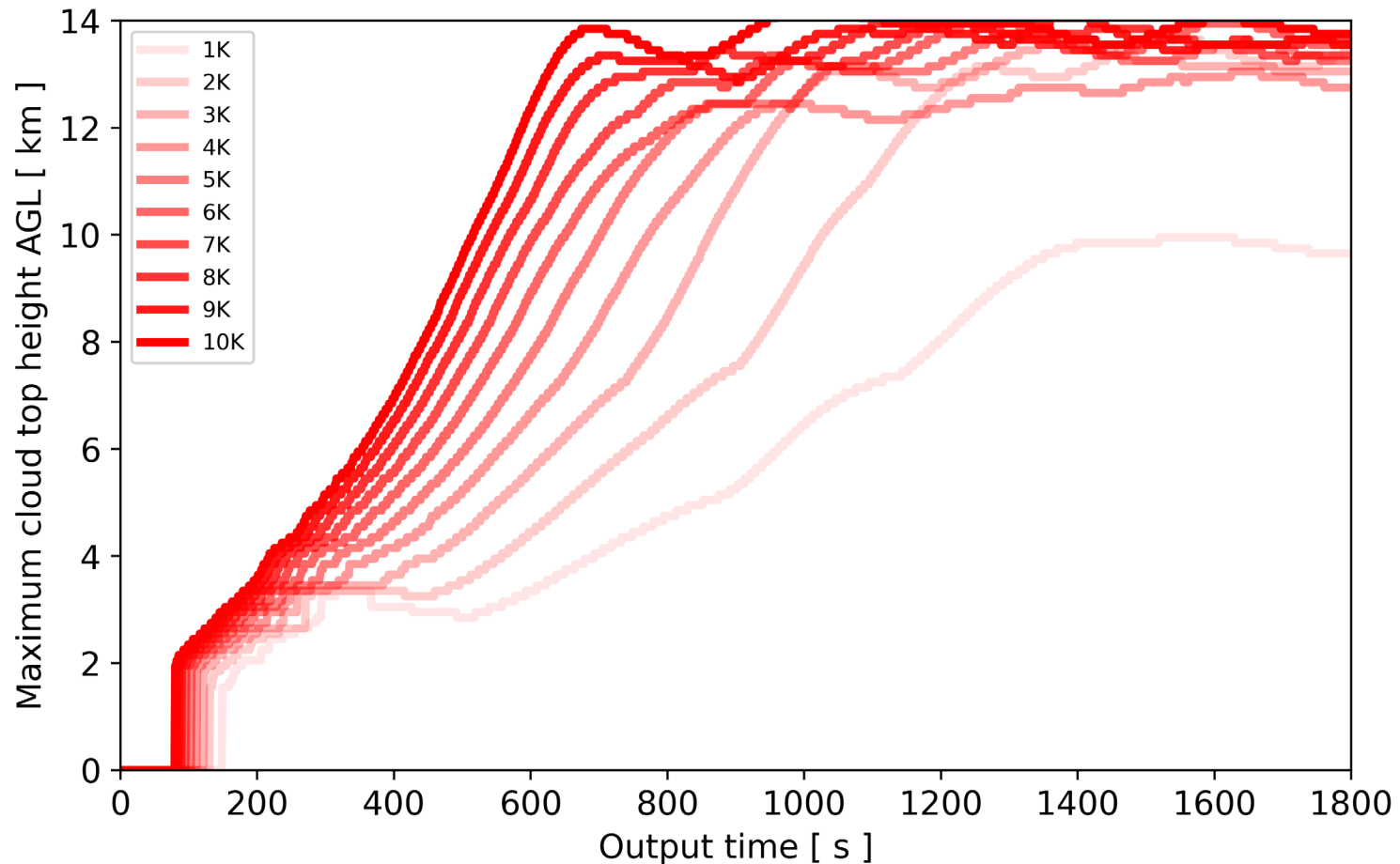
Shading = Max. passive tracer conc. [%]



Stronger w = smaller ϵ -driven dilution?



Smaller ϵ -driven dilution = taller cloud tops?



Conclusions and discussion

- Apparent inverse scaling between w and ϵ holds true at weaker w magnitudes
 - However, updraft width also increases \rightarrow may also explain decreasing ϵ
- At stronger w magnitudes, inverse scaling between w and ϵ vanishes
 - Updraft width approximately constant
- LOTS of caveats here...
 - Enough time for turbulence to “spin up” in these short-fused simulations?
 - Warm bubble convection initiation technique realistic?
 - Only one thermodynamic environment with no winds



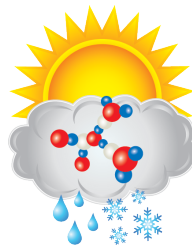
Future work

- Examine inverse scaling between w and ϵ with various convection initiation techniques to build statistical robustness:
 - Updraft nudging (*Naylor and Gilmore 2012; MWR*)
 - Forced convergence (*Loftus et al. 2008; MWR*)
 - Momentum forcing (*Morrison et al. 2015; MWR*)
 - Surface fluxes (*Morrison et al. 2022; JAS*)
- Attempt to tease out impact of updraft width on the inverse scaling between w and ϵ
- Additional simulations where...
 - Condensate loading term in the microphysics scheme is altered to vary w “less artificially”
 - Precipitation formation and ice processes removed to simplify things
 - Stability of free troposphere is altered



Thank you for your attention!

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And the toroidal circulations?

